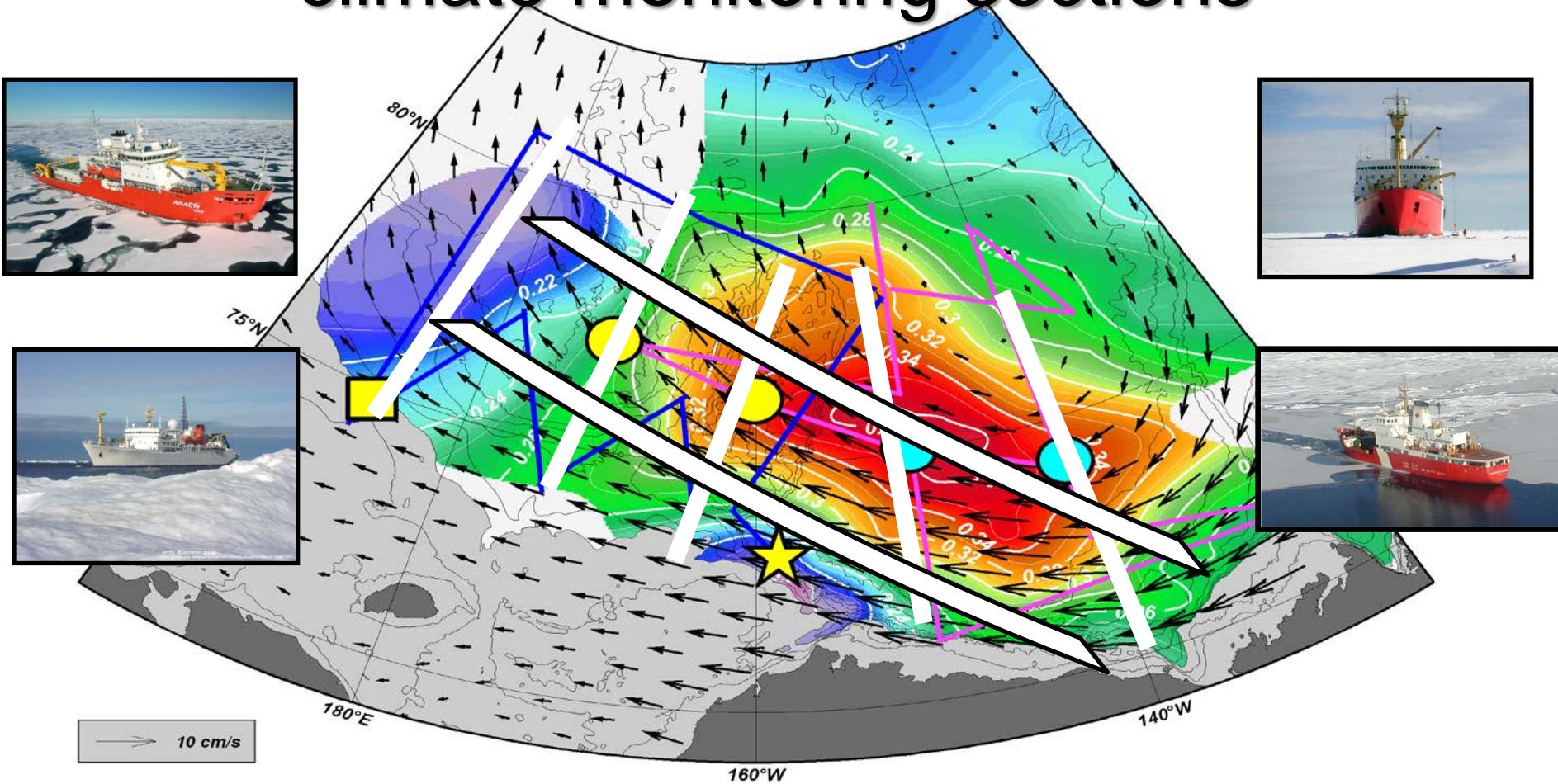


Presentation items for PAG meeting in KOPRI

Koji Shimada

Oct. 28-29

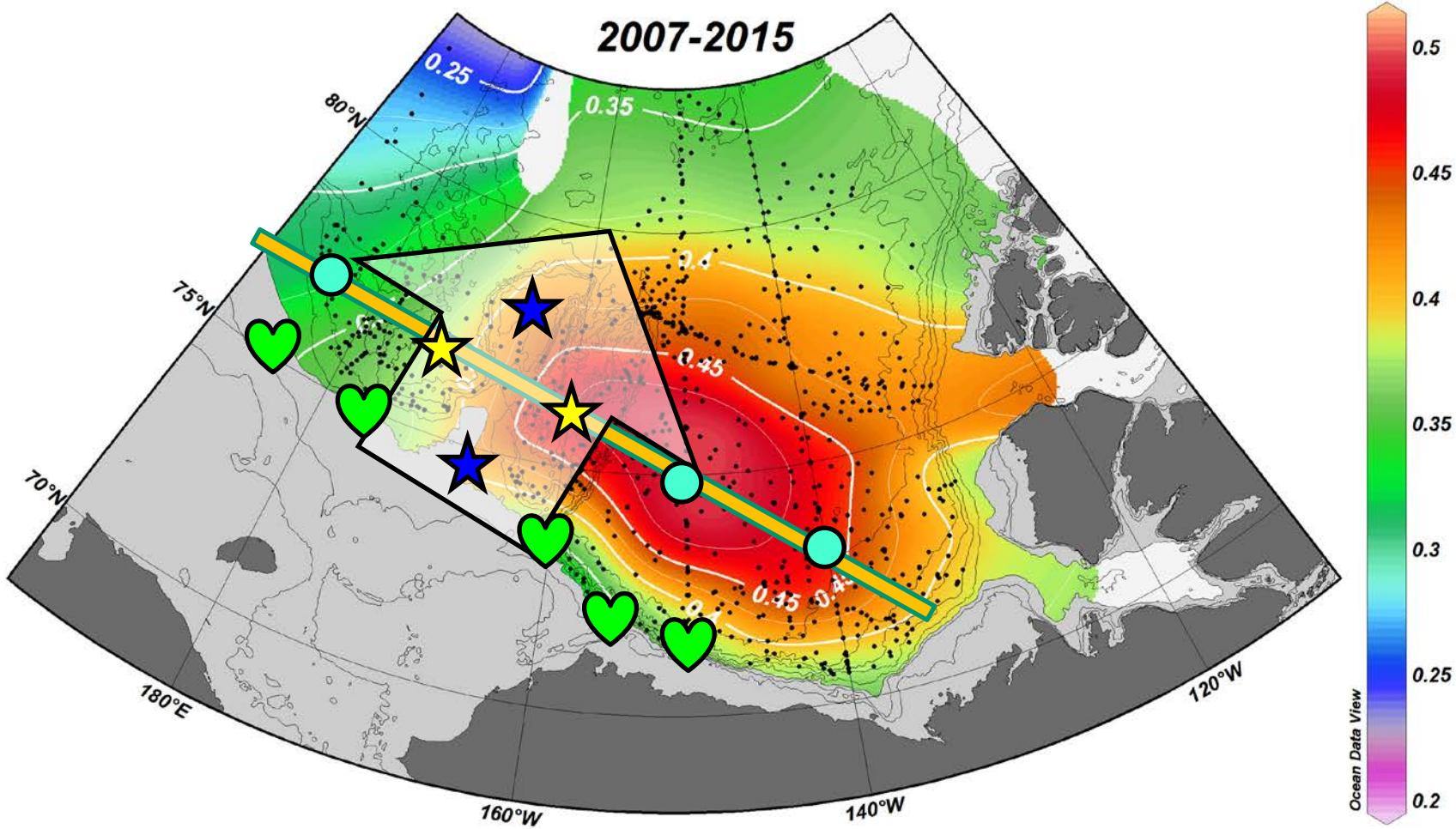
Proposed international Pacific Arctic climate monitoring sections

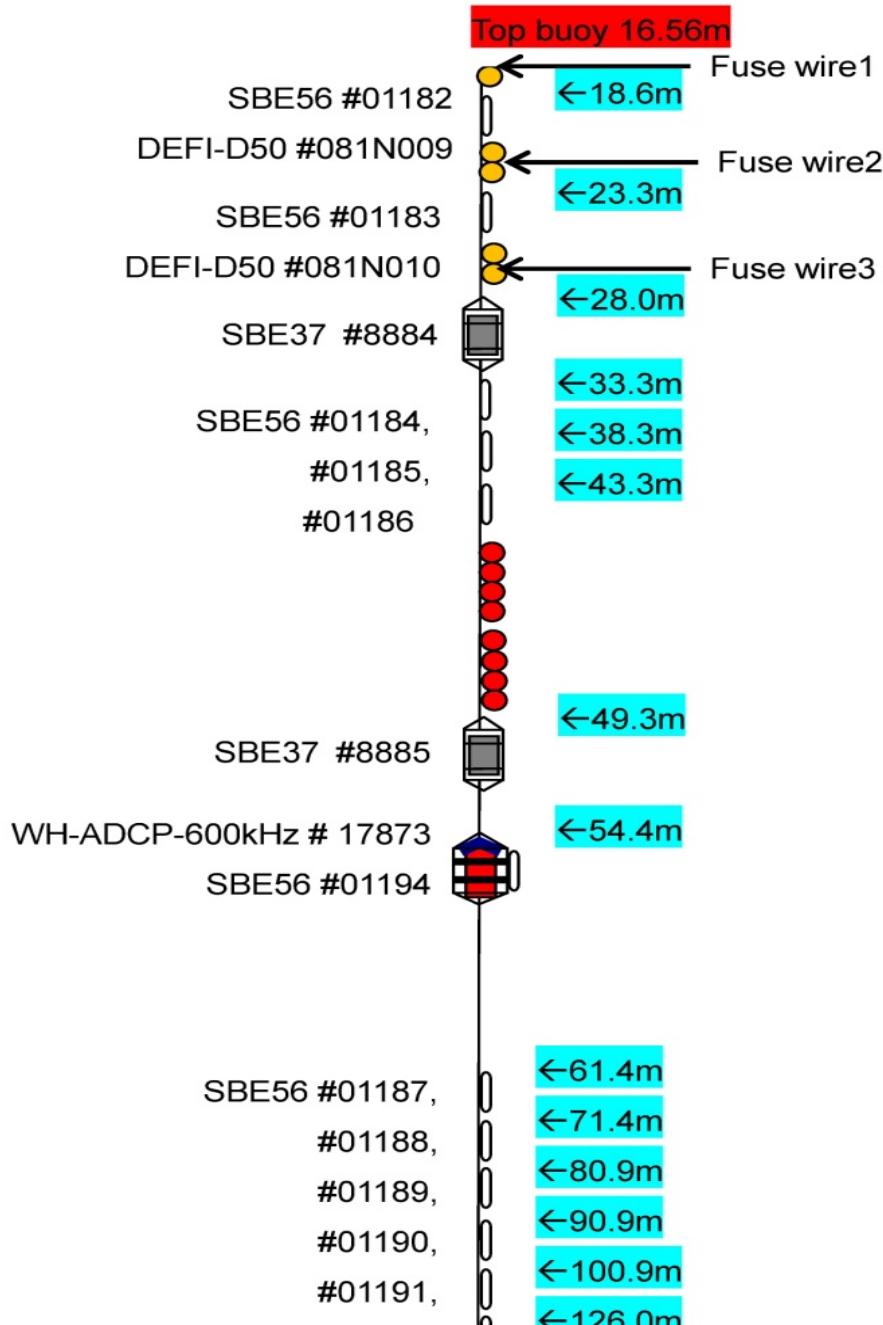


Background color: dynamic height at 100dbar relative to 800dbar from Mirai and Louis S. St-Laurent 2008 cruises (Oceanic Beaufort Gyre)

Black vectors: average sea ice motion vectors for Nov. 2007- Apr. 2008 (Sea Ice Beaufort Gyre)

Symbols: Mooring array in 2012-2013 (TUMSAT/KOPRI/NIPR & WHOI)





Mooring CP13

Position:

(Anchor dropped position near the A frame)

77°28.331'N, 164°07.178'W

(Triangulation result)

77°28.3353'N, 164°07.0789'W

[77.47255°N, 164.11798°W]

Deployment time (Anchor):

August 26, 2013 23:42 (UTC)

Bottom depth:

277m
(272m[Multi beam depth] + 5m)

CTDs:

SBE37-SM S/No. 8884, 8885

Temperature loggers:

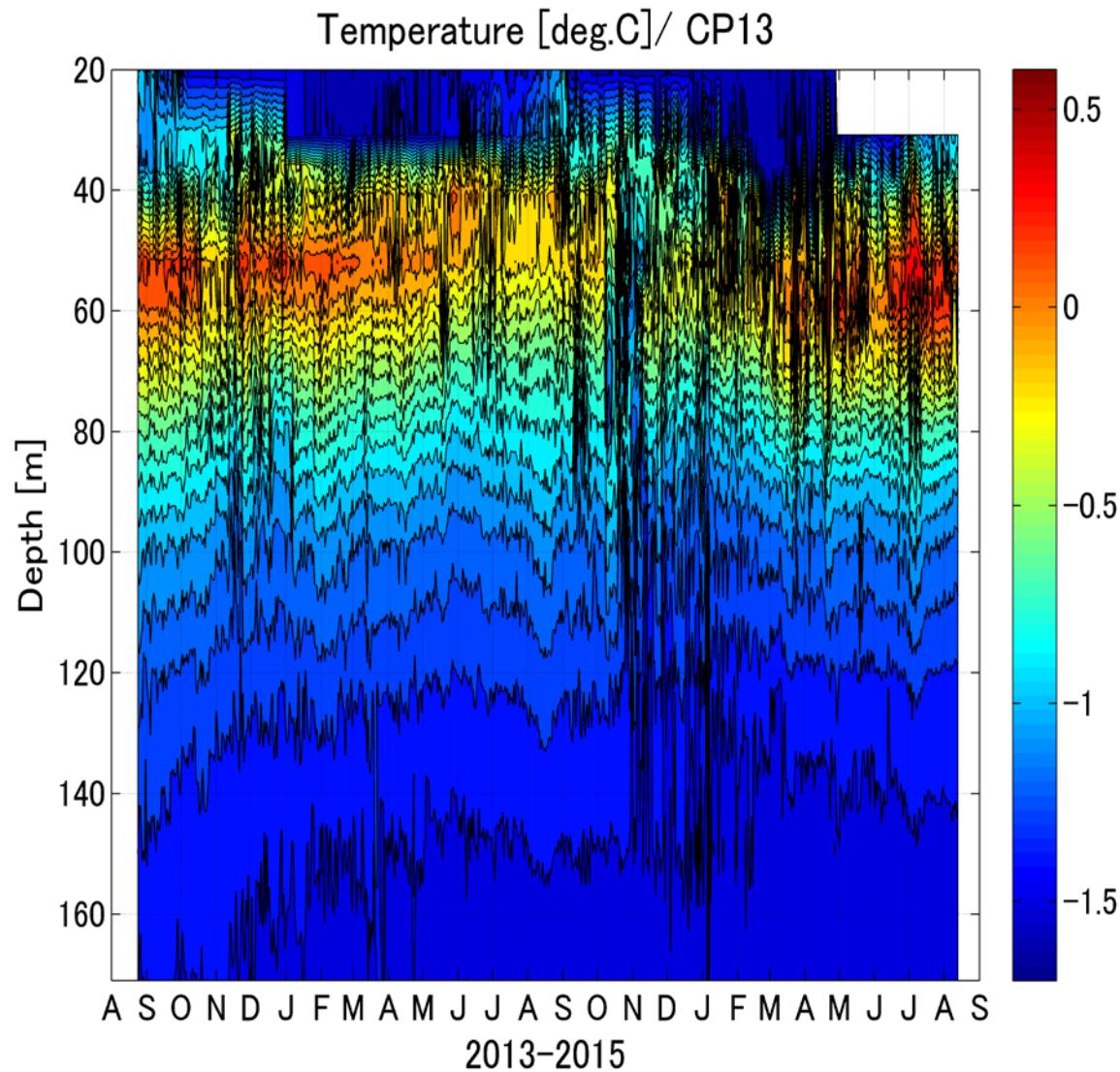
SBE56 S/No. 01182, 01183, 01184, 01185, 01186, 01187, 01188, 01189, 01190, 01191, 01192, 01193, 01194, 01195, 01196

Pressure sensors:

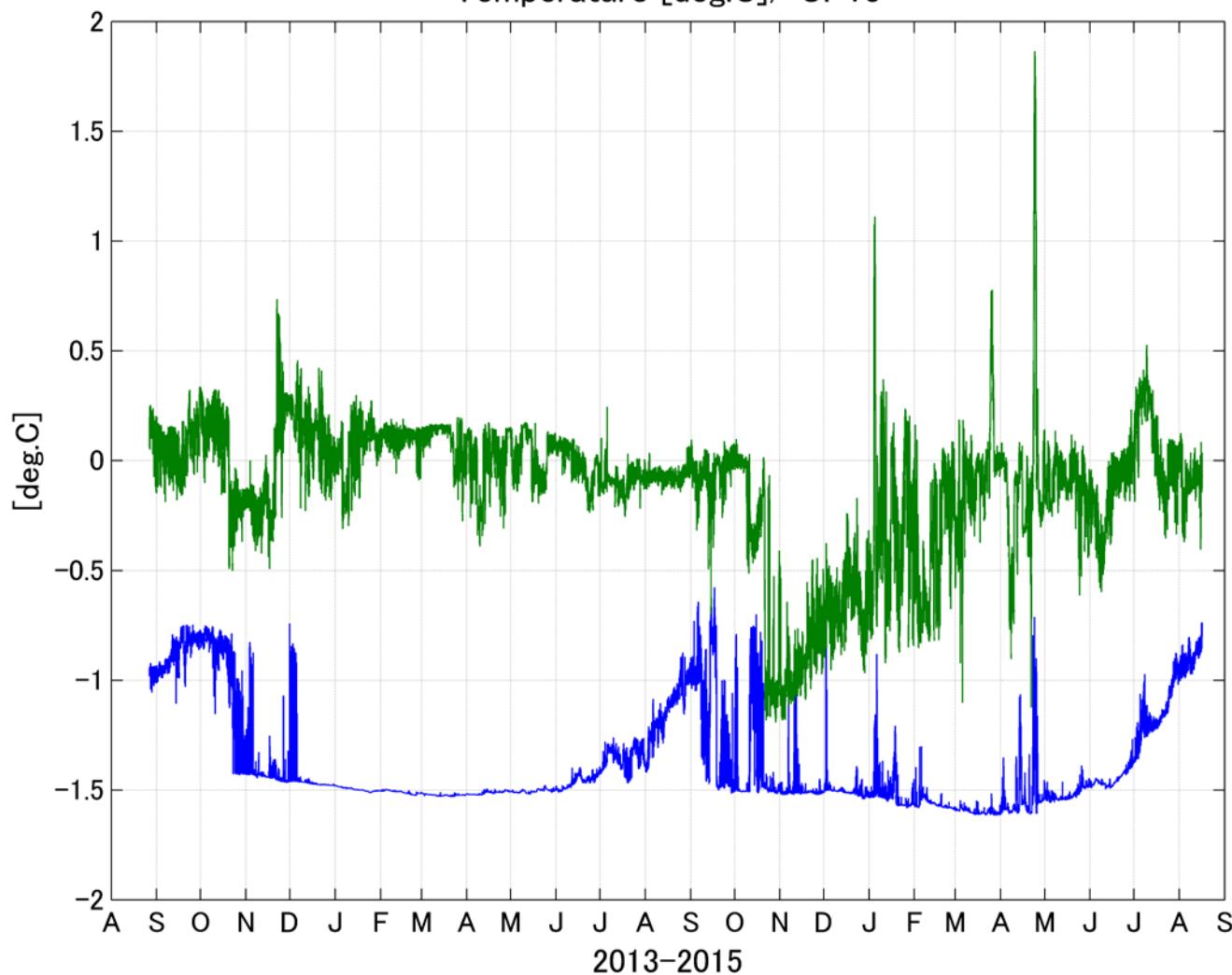
DEFI-D50 S/No. 081N009, 081N010

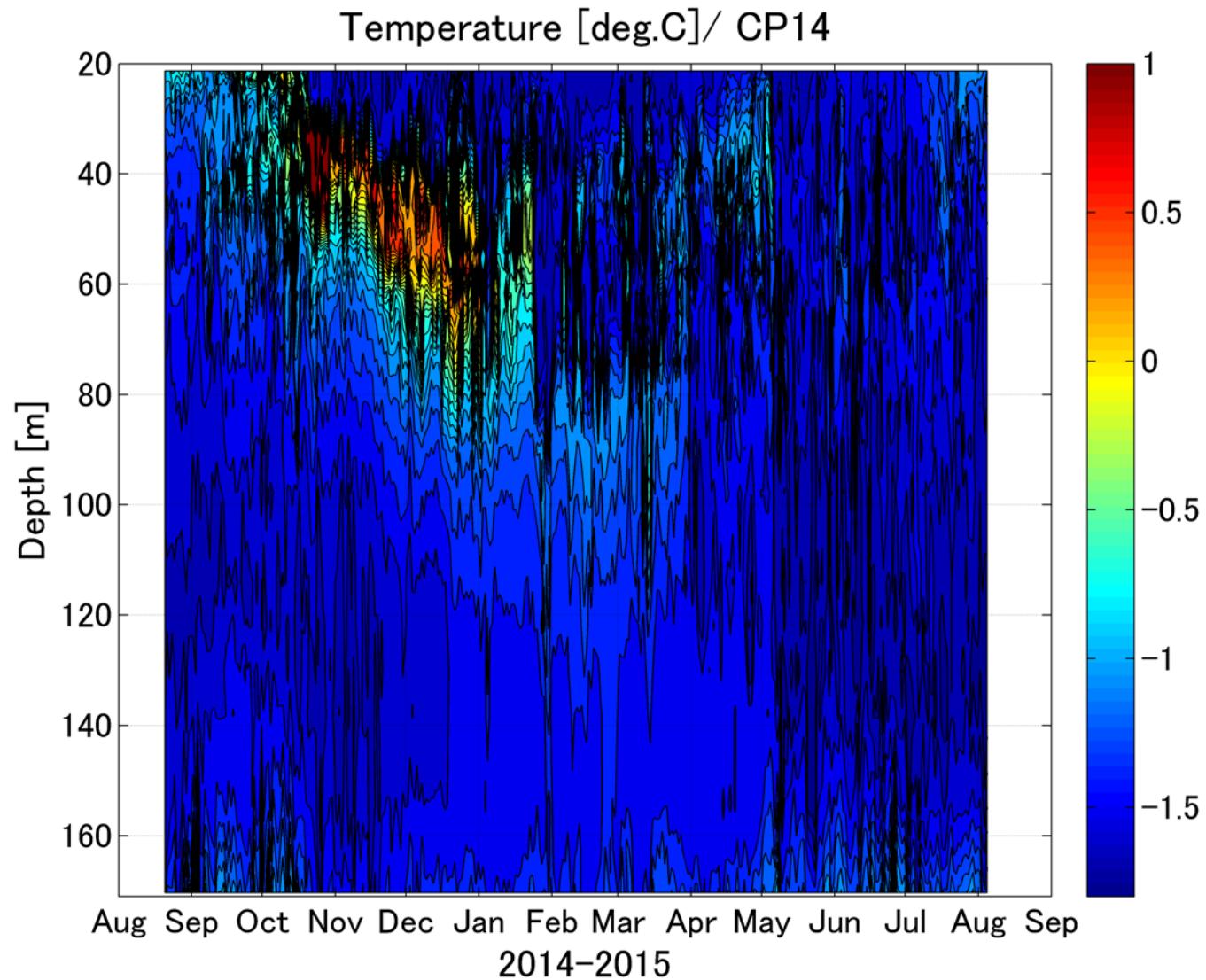
Current Meter:

WH-ADCP-600kHz Sentinel S/No. 17873



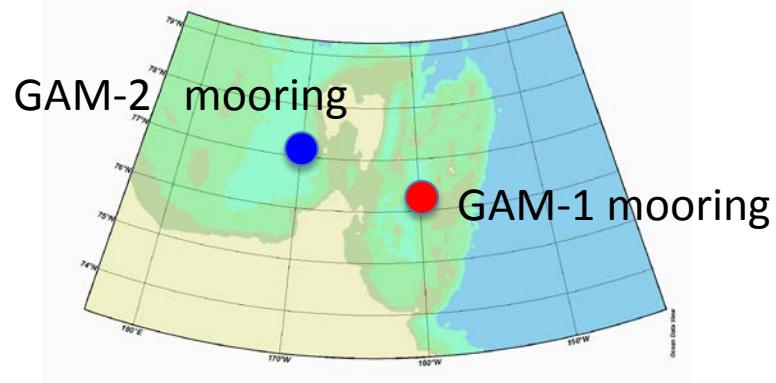
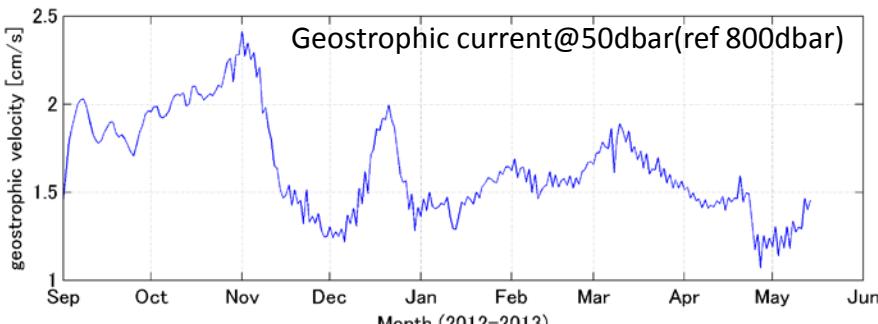
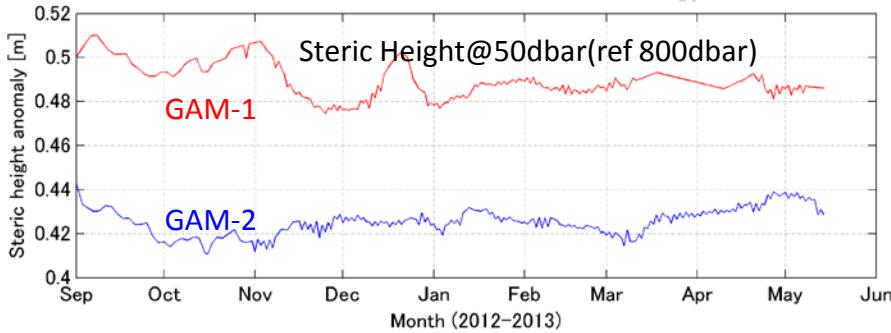
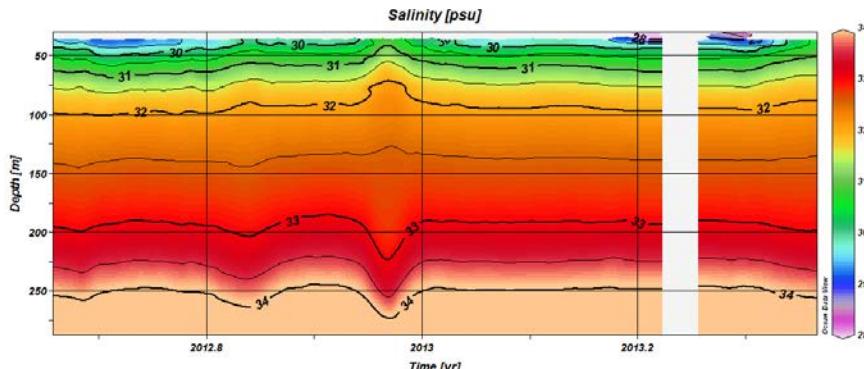
Temperature [deg.C]/ CP13





Strategy of monitoring for climate oriented studies

Variations of upper ocean circulation



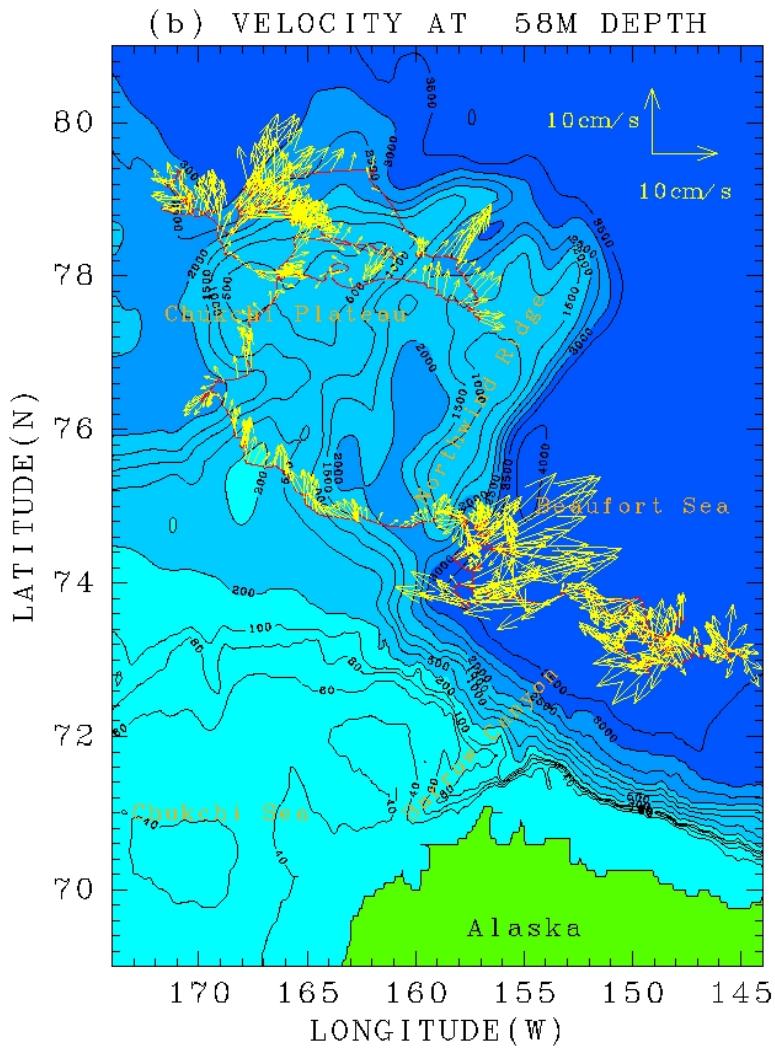
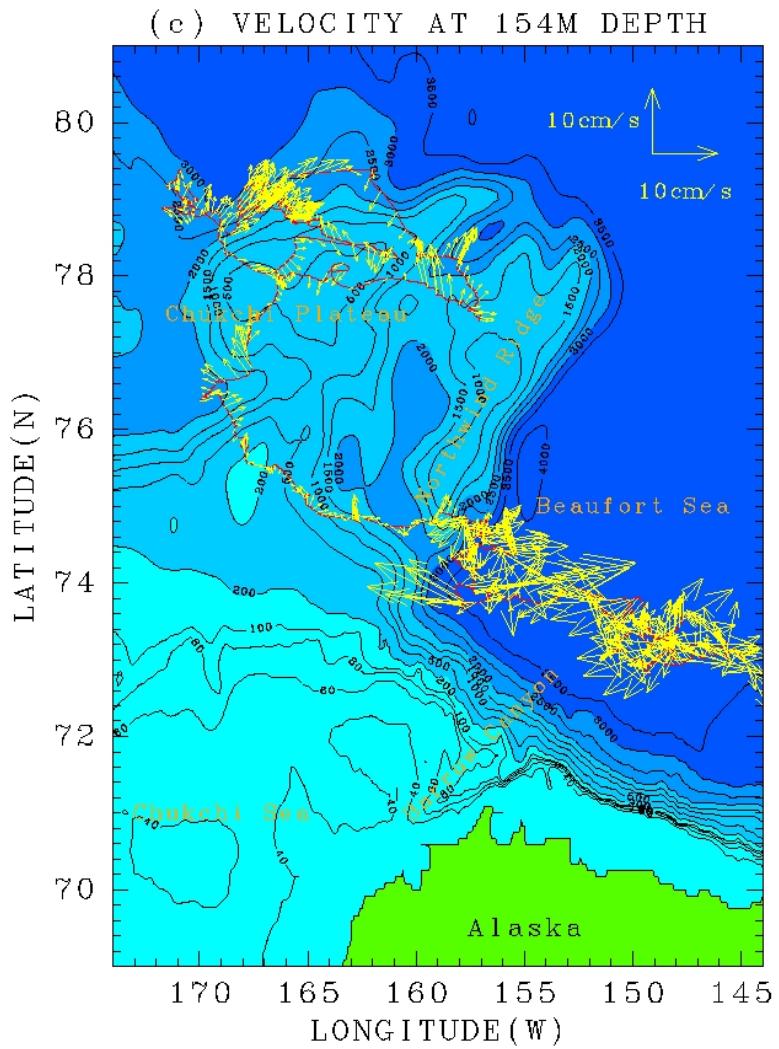
Mean northward current: 1.64cm/s
(500km/year)

STD of the mean current: 0.28cm/s
(including eddies)

In the CBL area west of the Northwind Ridge, almost no seasonal variations, just inter-annual variations

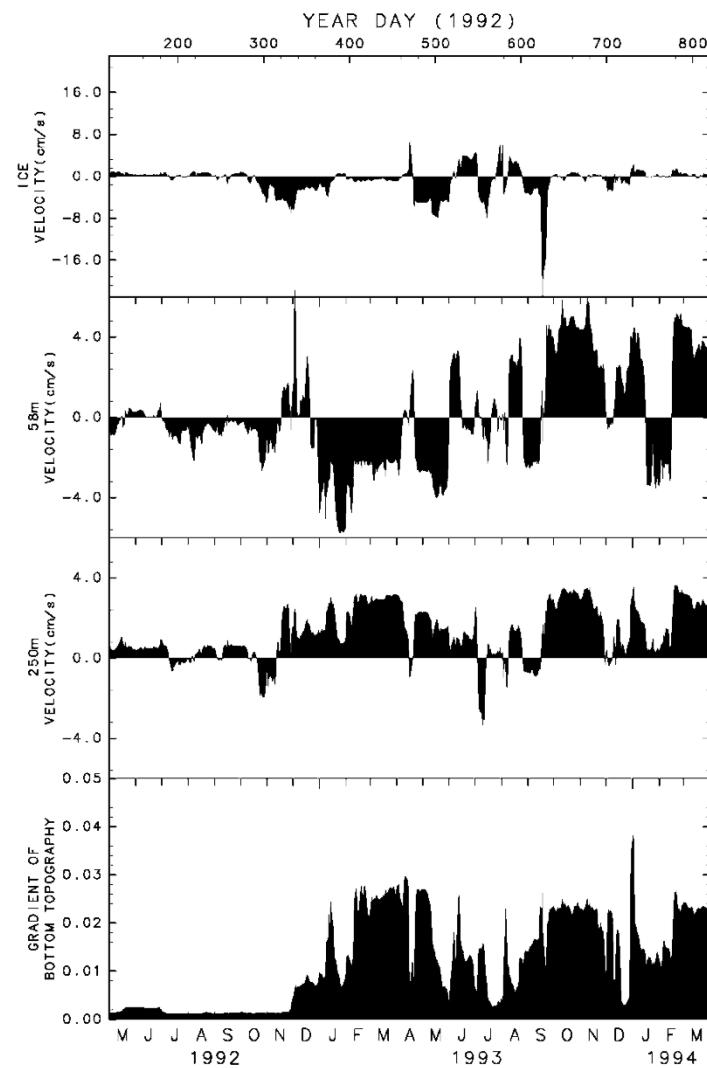
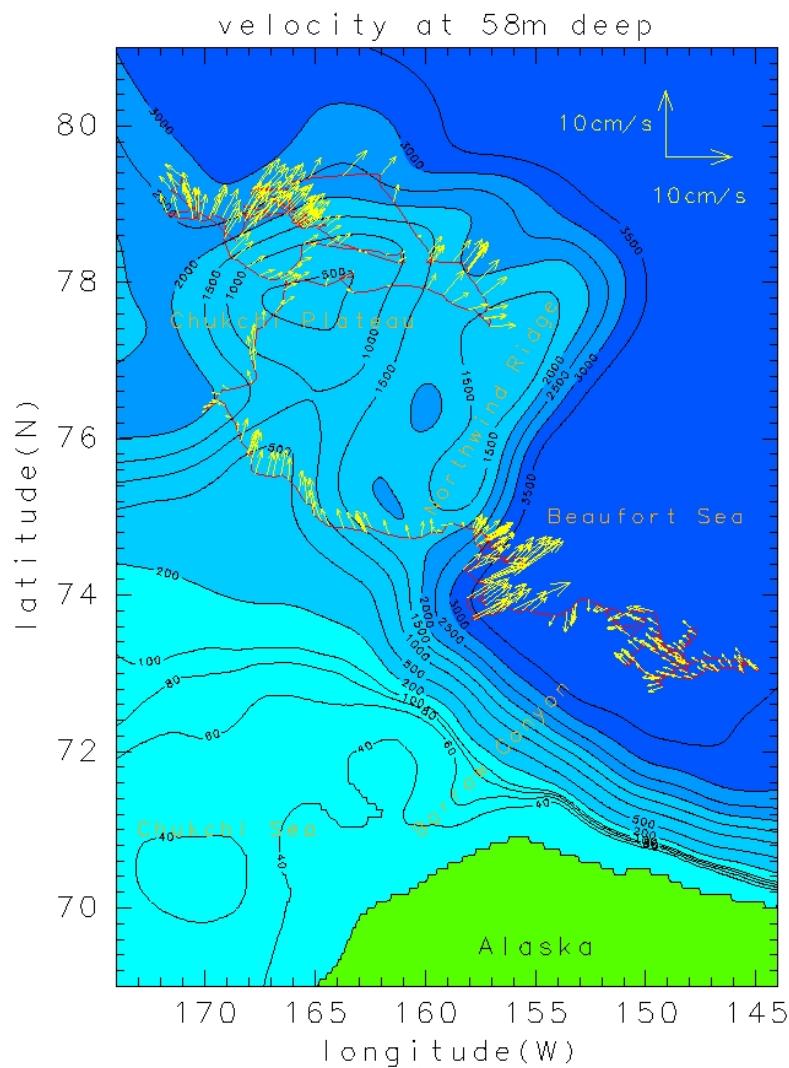


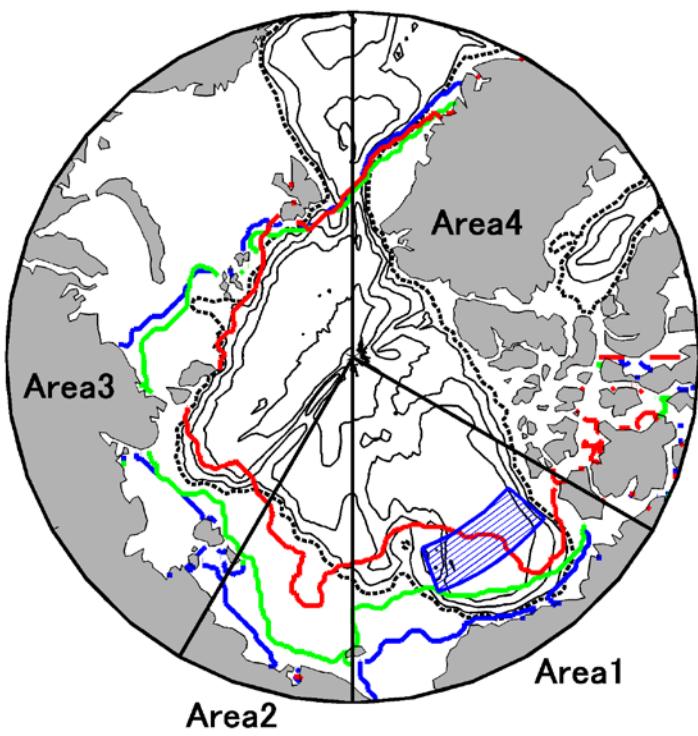
In the western Canada Basin, one section per year is enough to detect the change of large-scale ocean circulation



ノースウインド海嶺の重要性

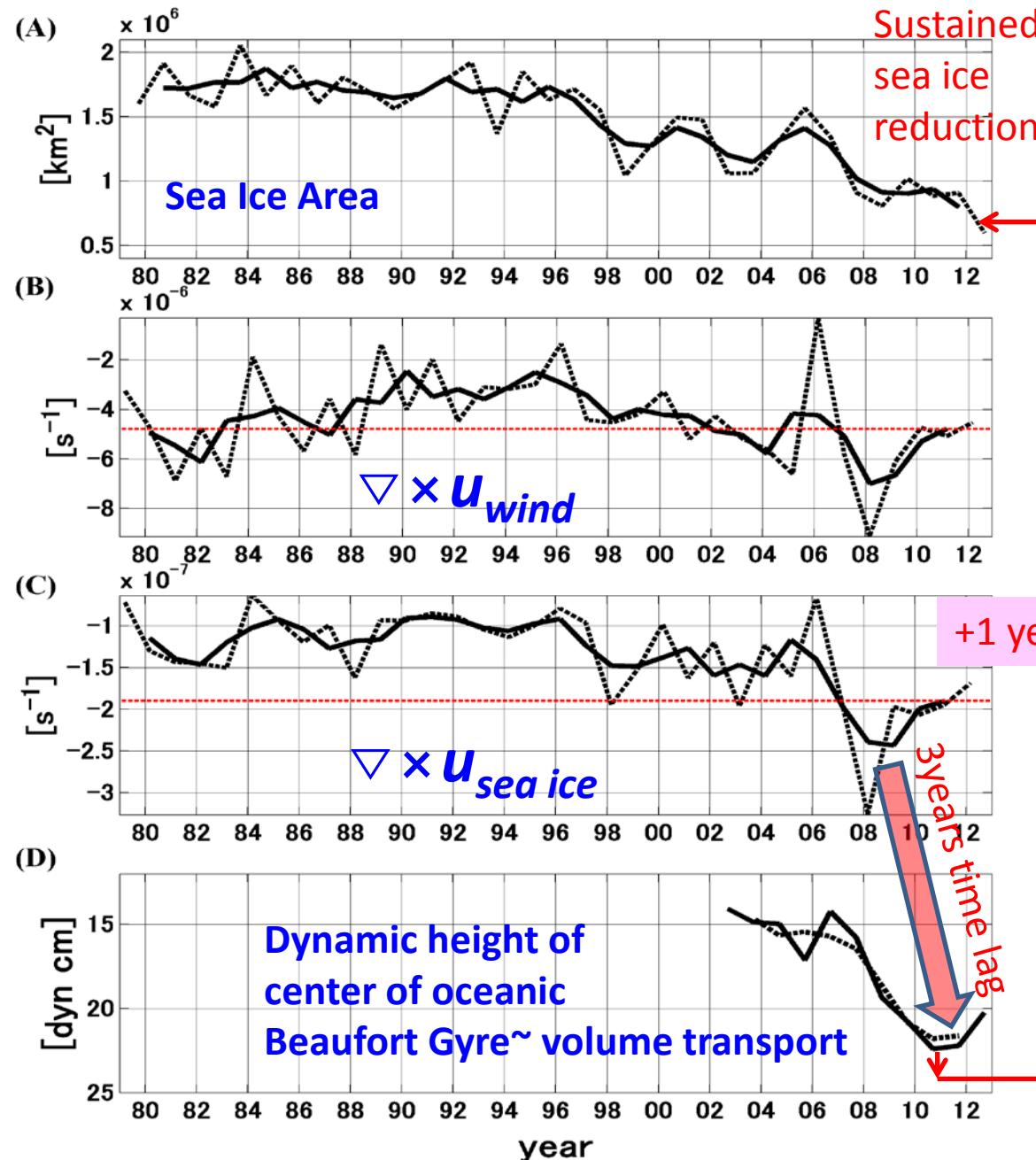
UT AND GRADIENT OF BOTTOM TOPOGRAPHY ($R=50\text{KM}$)

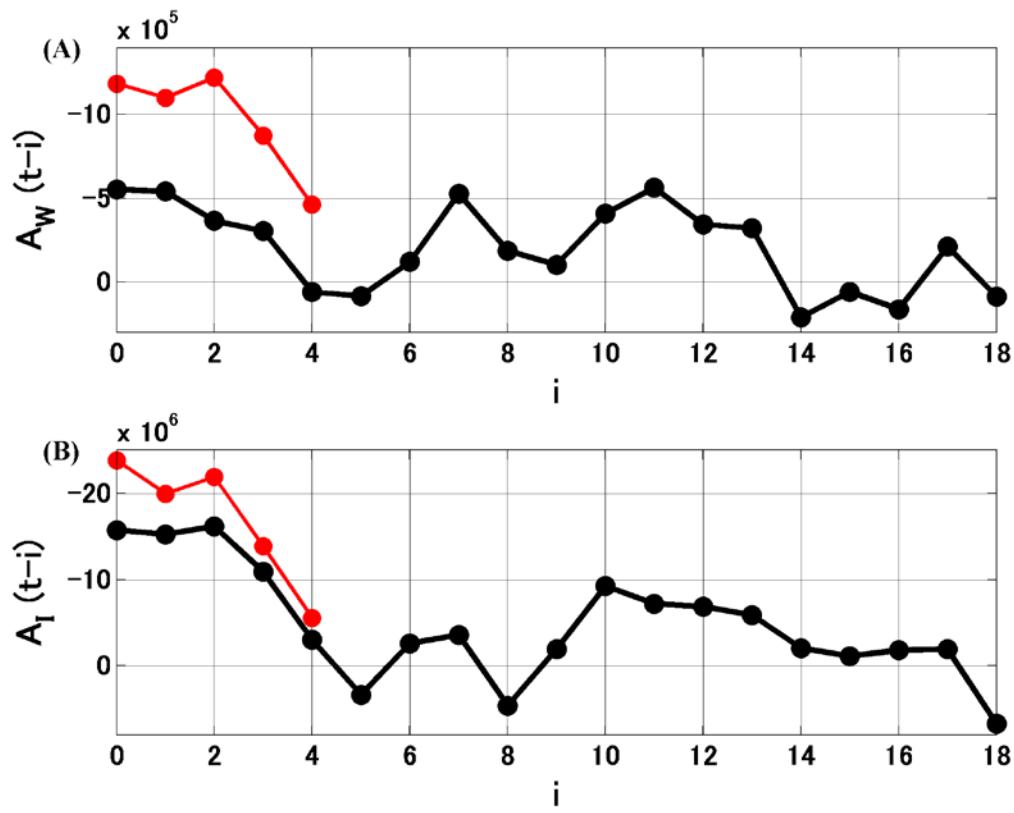




Upper ocean response
delayed about 3 years
relative to the surface forcings
(wind or sea ice motion).

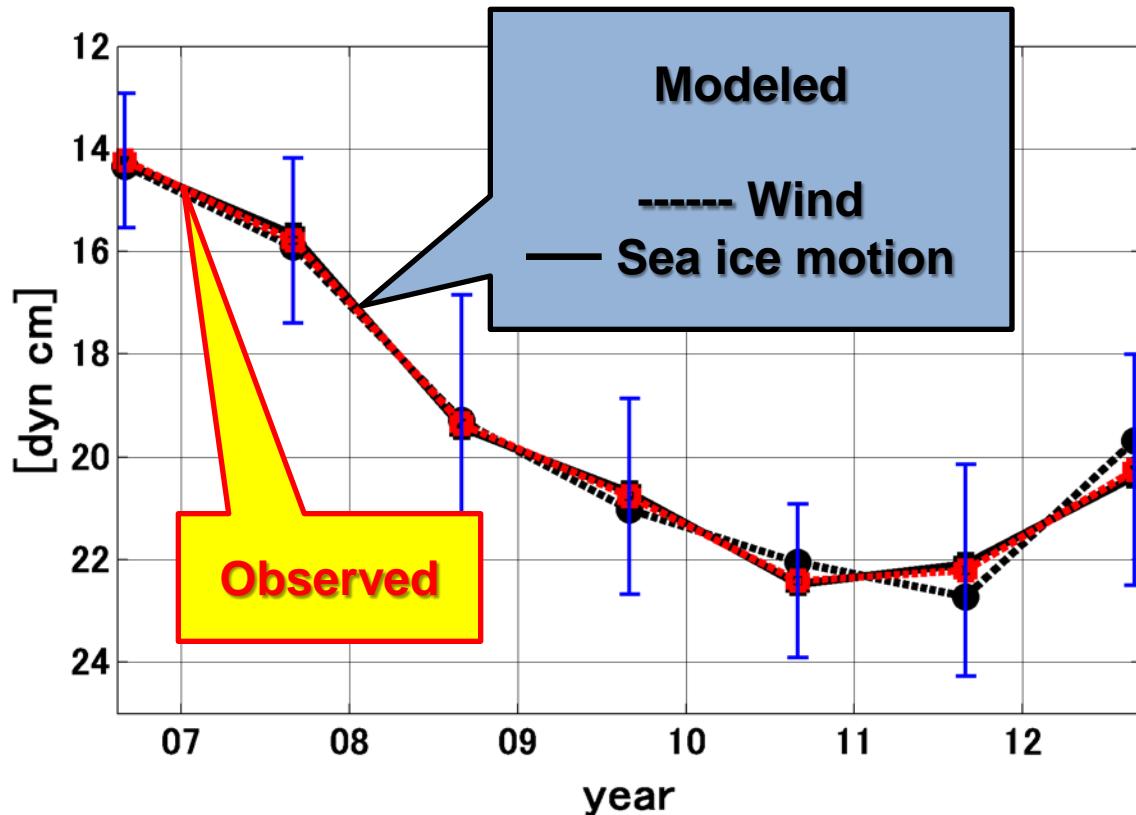
Yoshizawa et al., (2015)





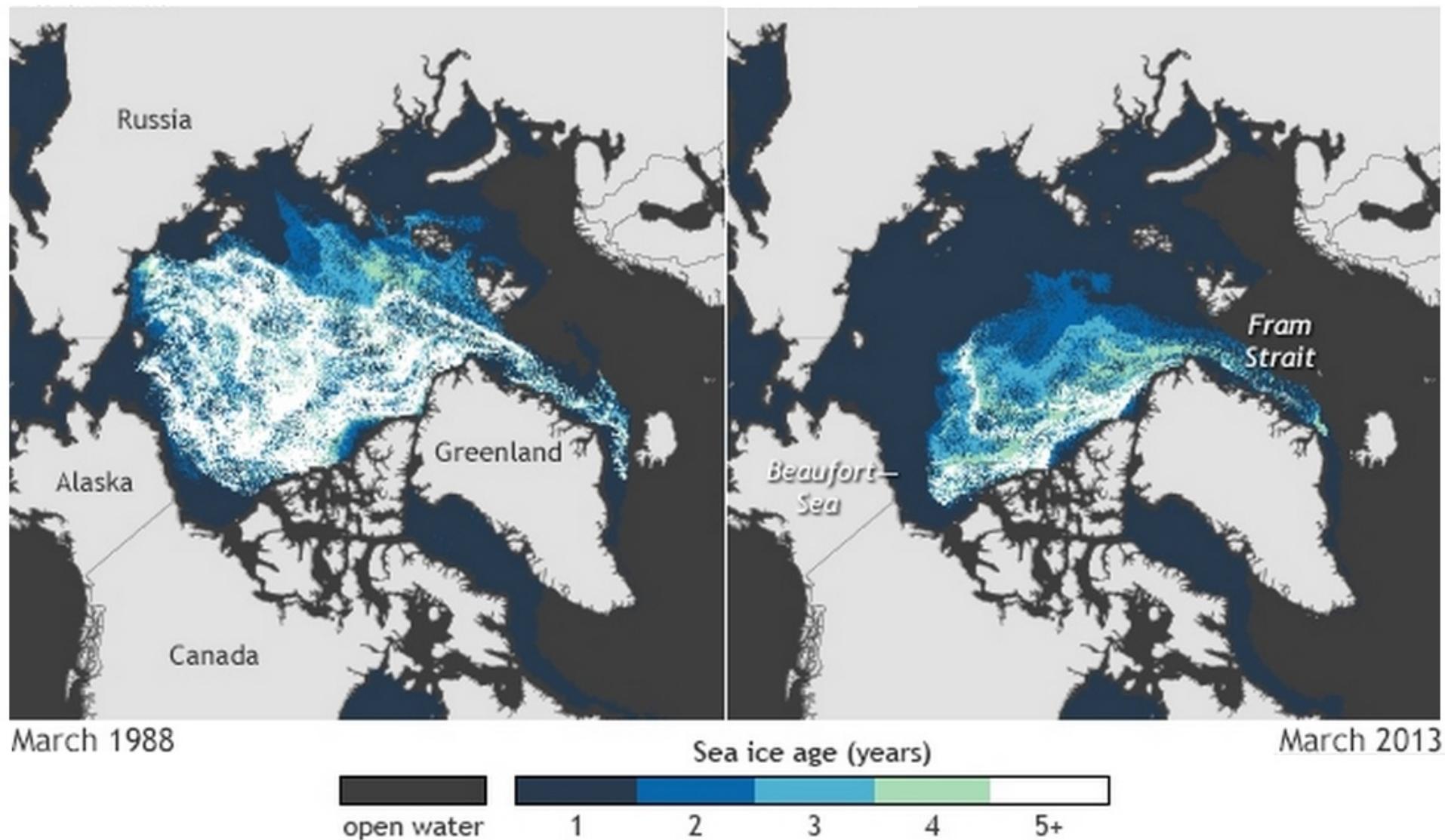
$$D_{wind\ model}(t) = \sum_{i=0}^{\infty} [A_W(t-i) \cdot \nabla \times \vec{u}_W(t-i)] + B_W$$

$$D_{sea\ ice\ model}(t) = \sum_{i=0}^{\infty} [A_I(t-i) \cdot \nabla \times \vec{u}_I(t-i)] + B_I$$



$$D_{wind\ model}(t) = \sum_{i=0}^4 [A_W(t-i) \cdot \nabla \times \vec{u}_W(t-i)] + B_W$$

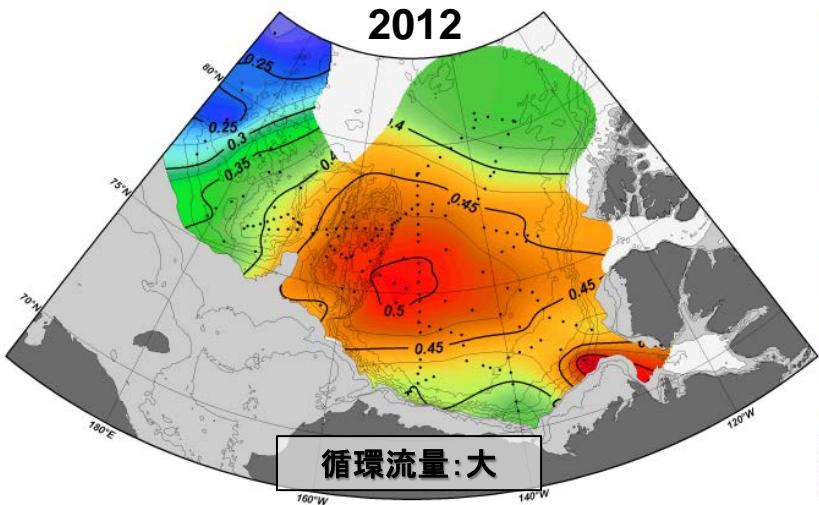
$$D_{sea\ ice\ model}(t) = \sum_{i=0}^4 [A_I(t-i) \cdot \nabla \times \vec{u}_I(t-i)] + B_I$$



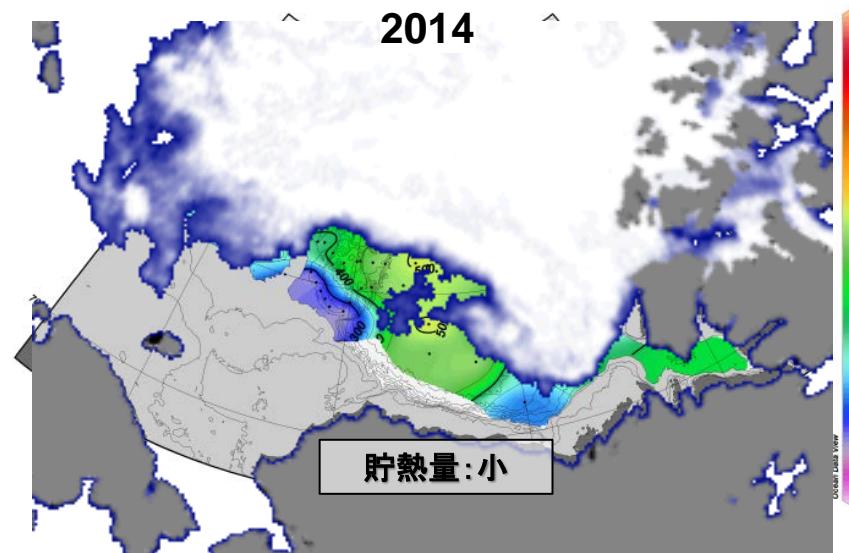
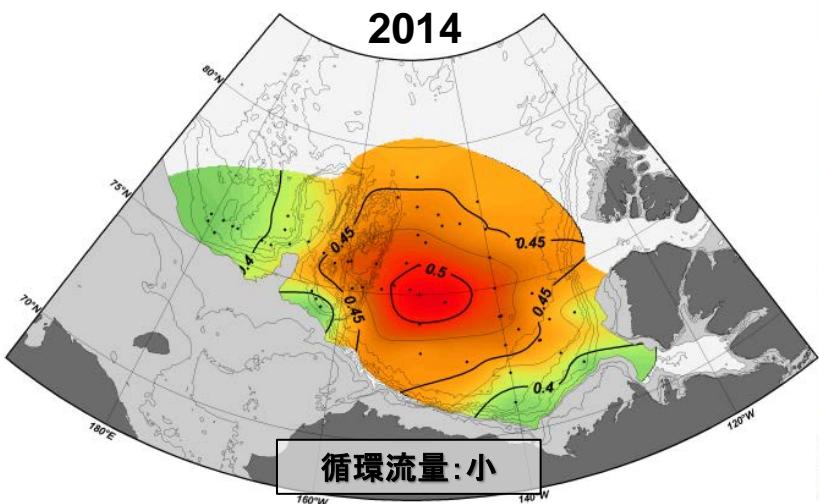
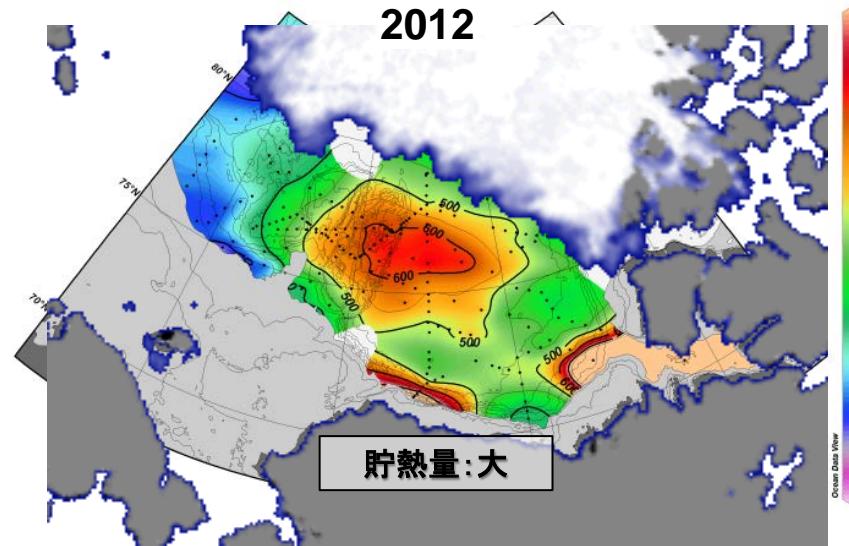
by Mark Tschudi, University of Colorado

海洋循環・構造変動を海水変動

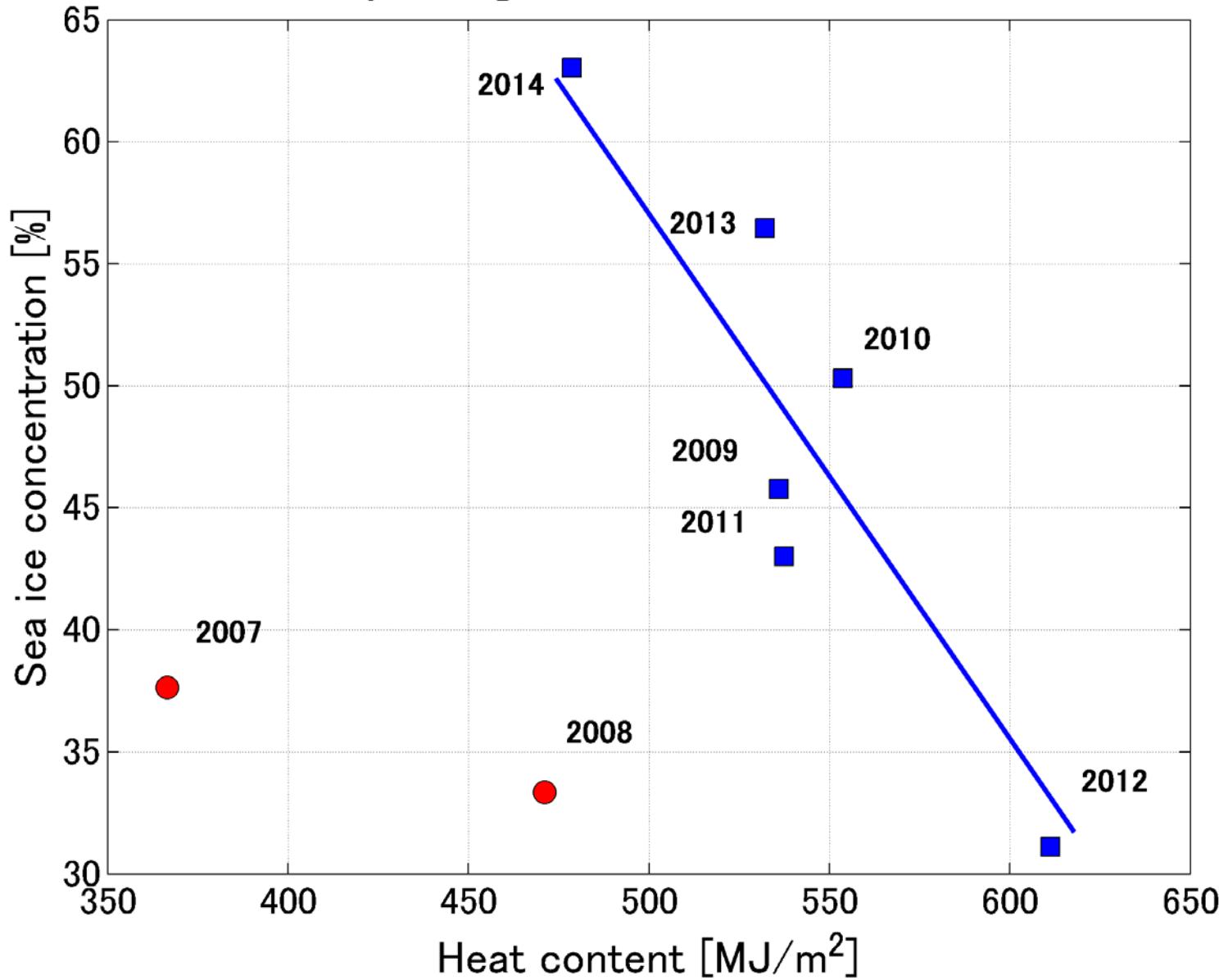
海洋循環
(800dbar基準50dbar面の力学高度偏差)



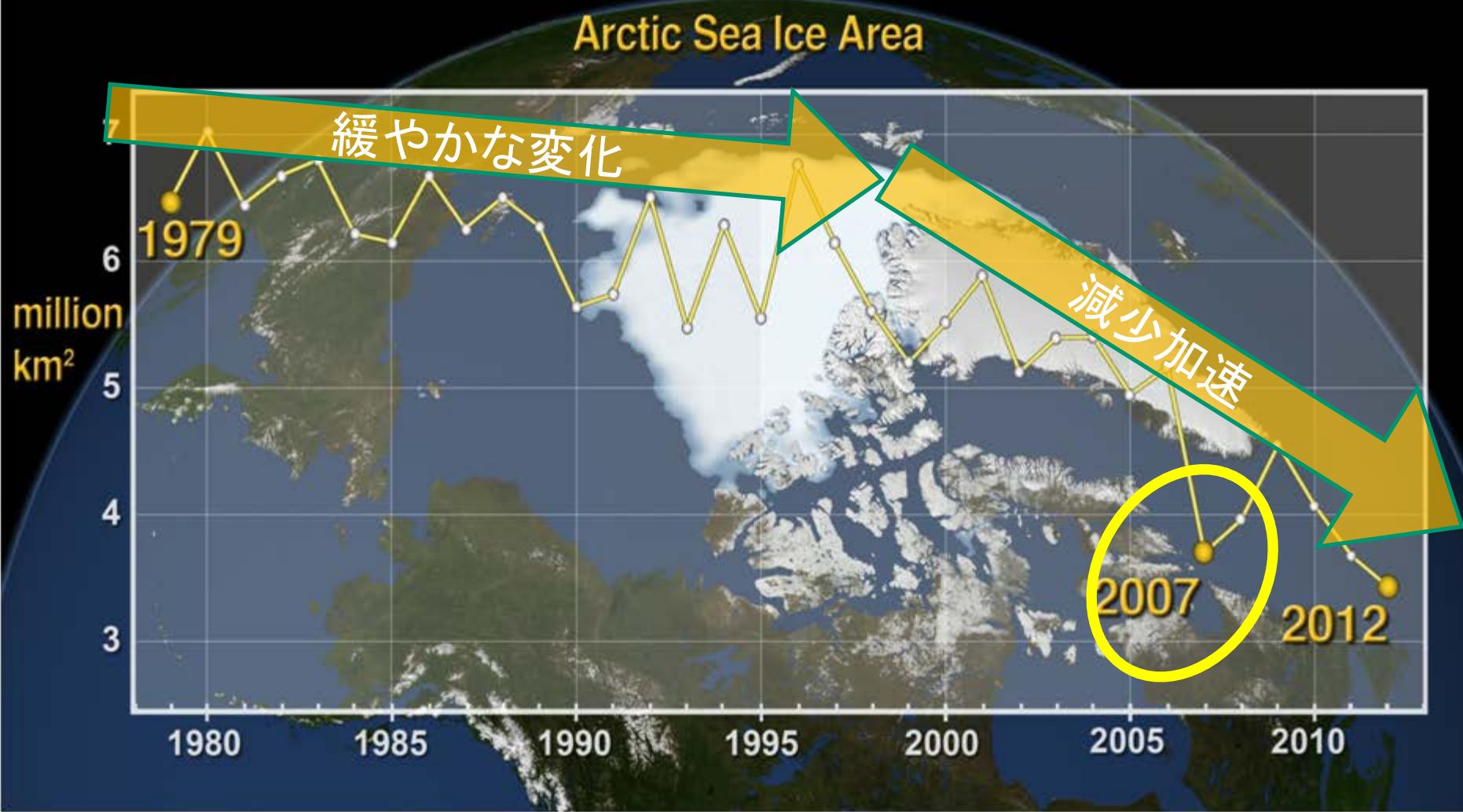
貯熱量
太平洋水層(20-150m)



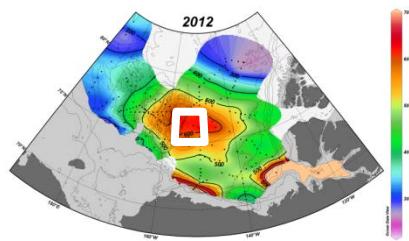
July & August, 74–78°N, 150–180°W



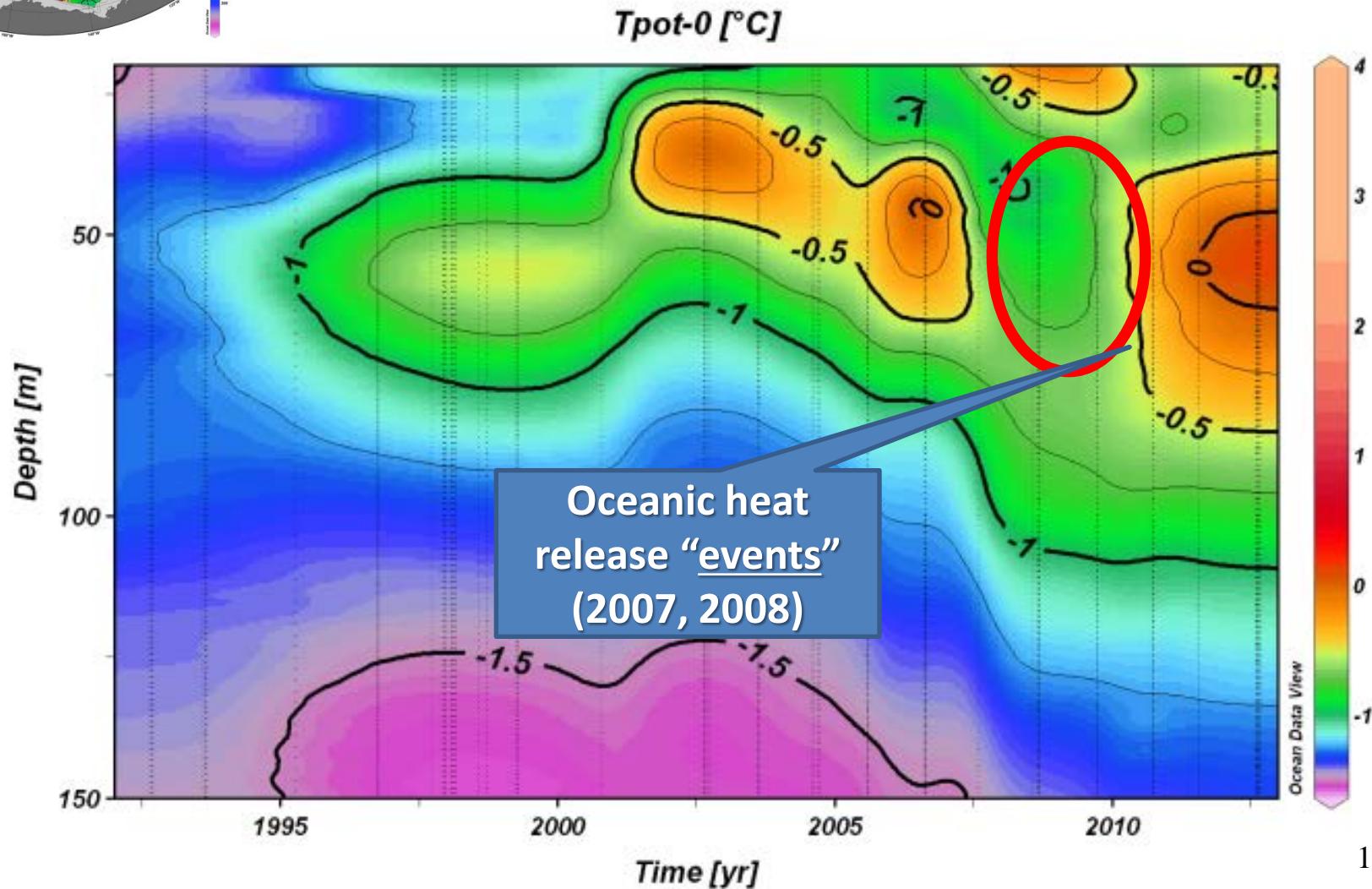
海氷面積の変化1979～2012



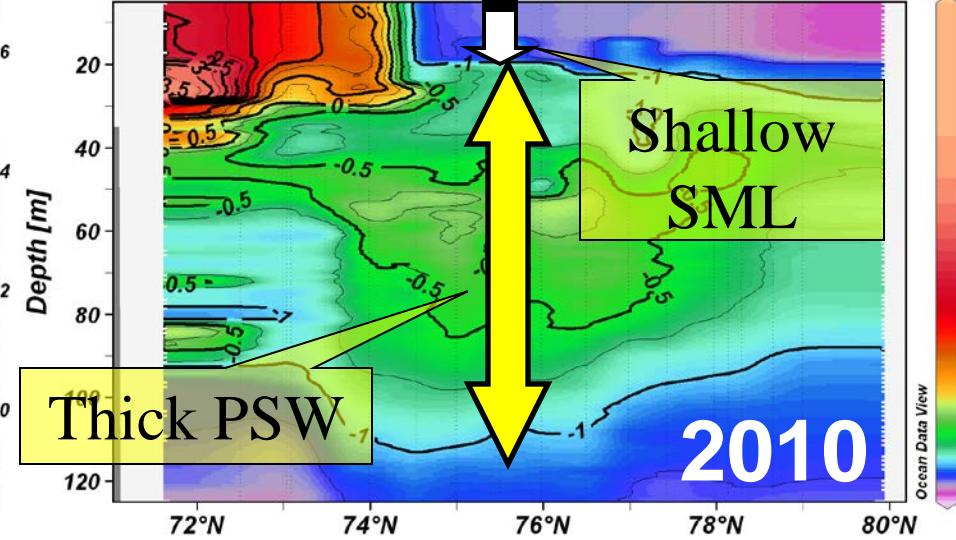
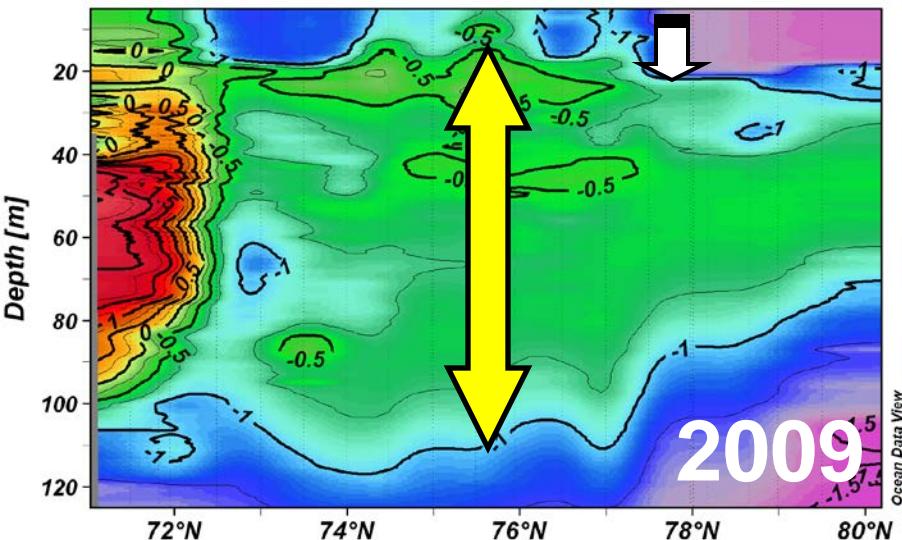
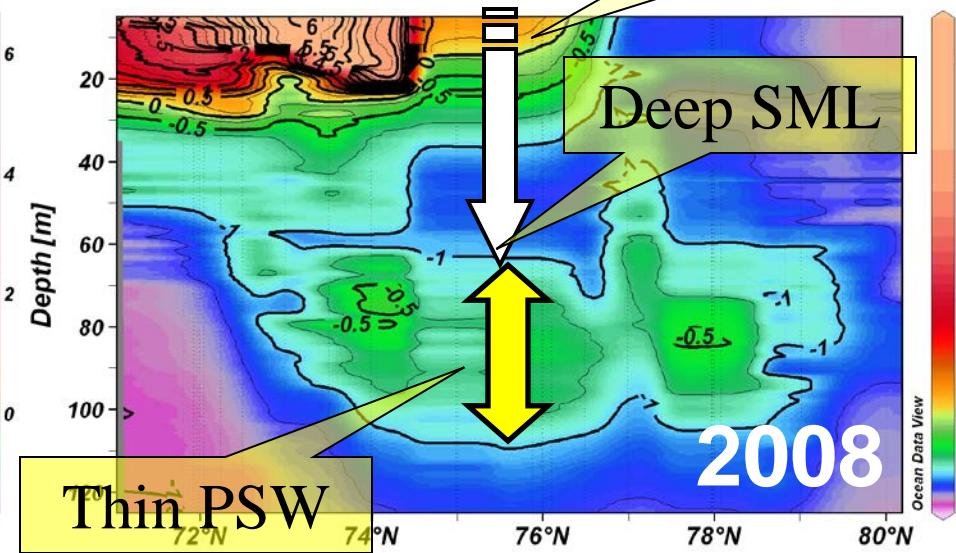
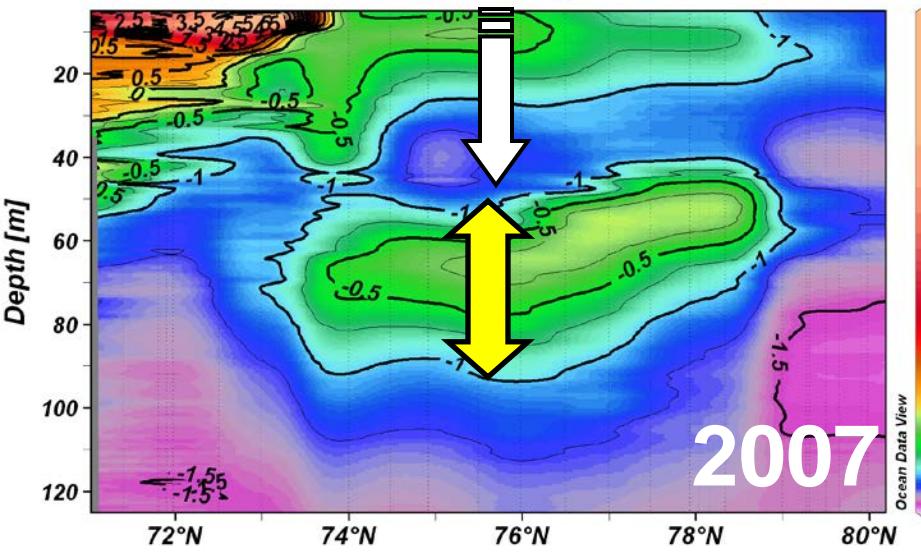
Time series of temperature on the Northwindridge



Heat release “event”.



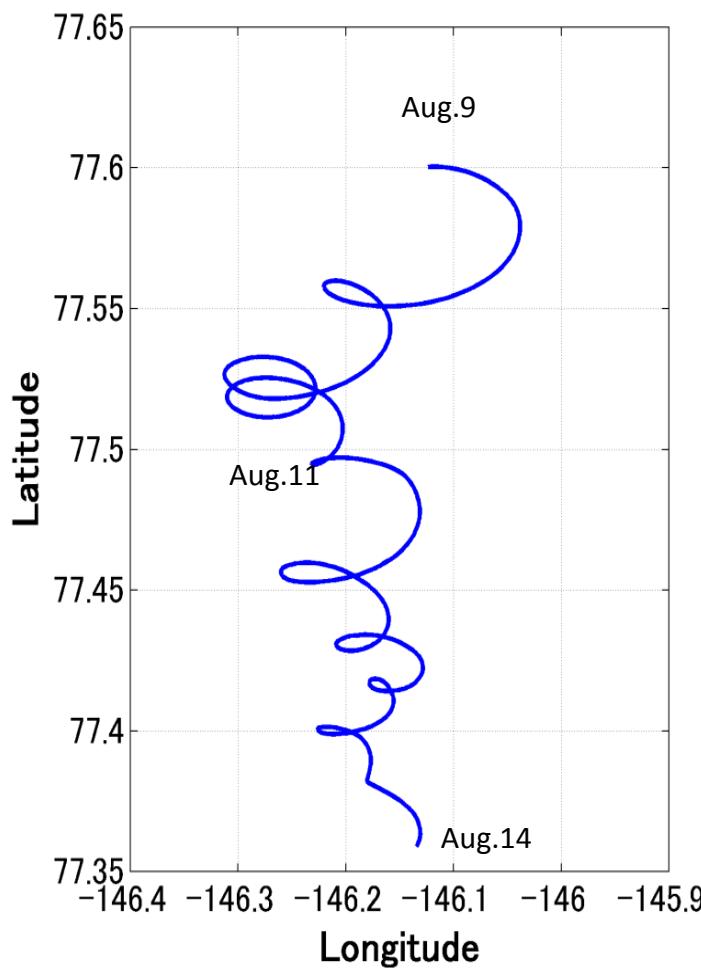
Changes in temperature along 150W



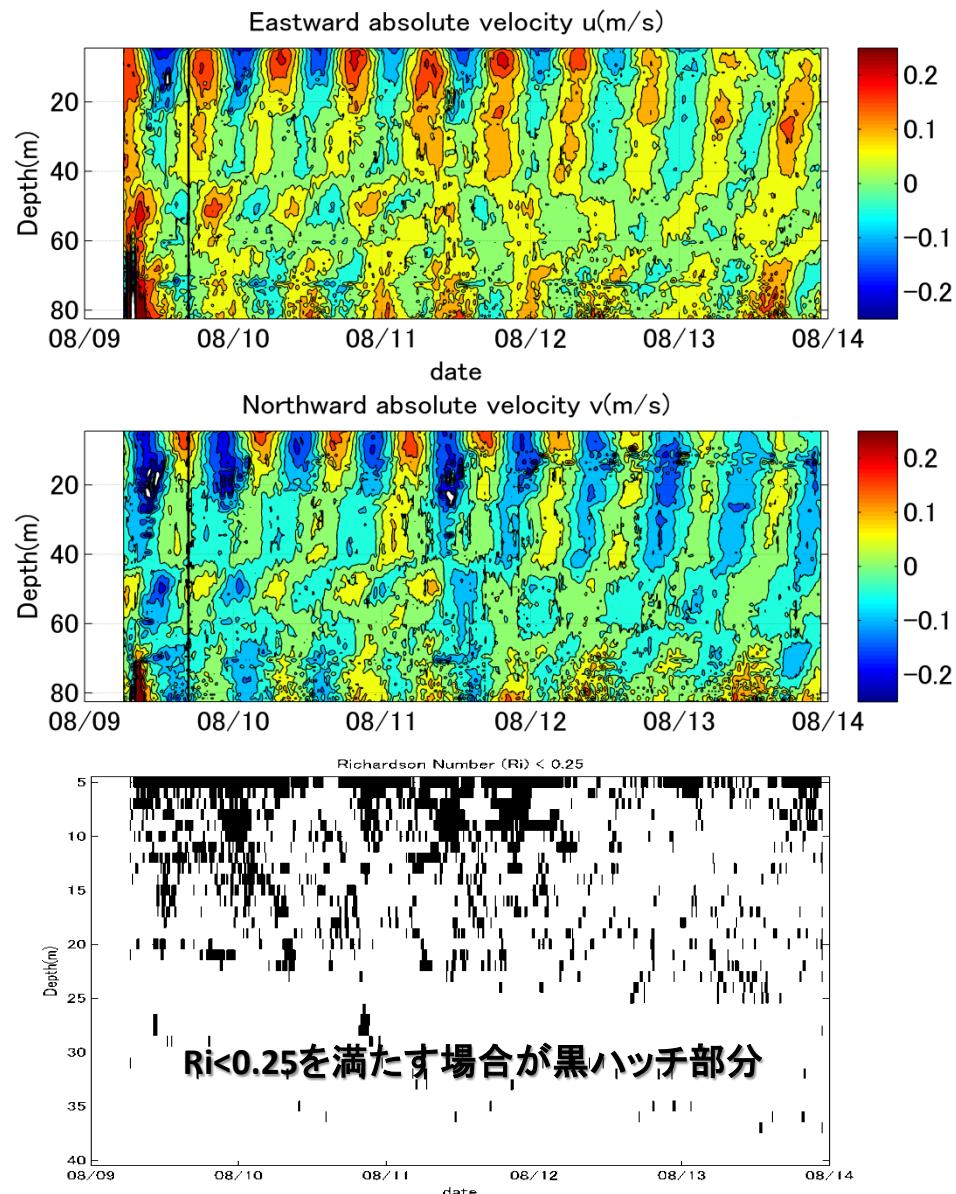
Solar heat

2.研究成果

戦略研究目標に対して何を達成したのか、テーマごとに図表・数値を用いて分かりやすく示して下さい。
特に注目に値するチャンピオンデータや代表的研究成果については、冒頭に特出しで明記して下さい。

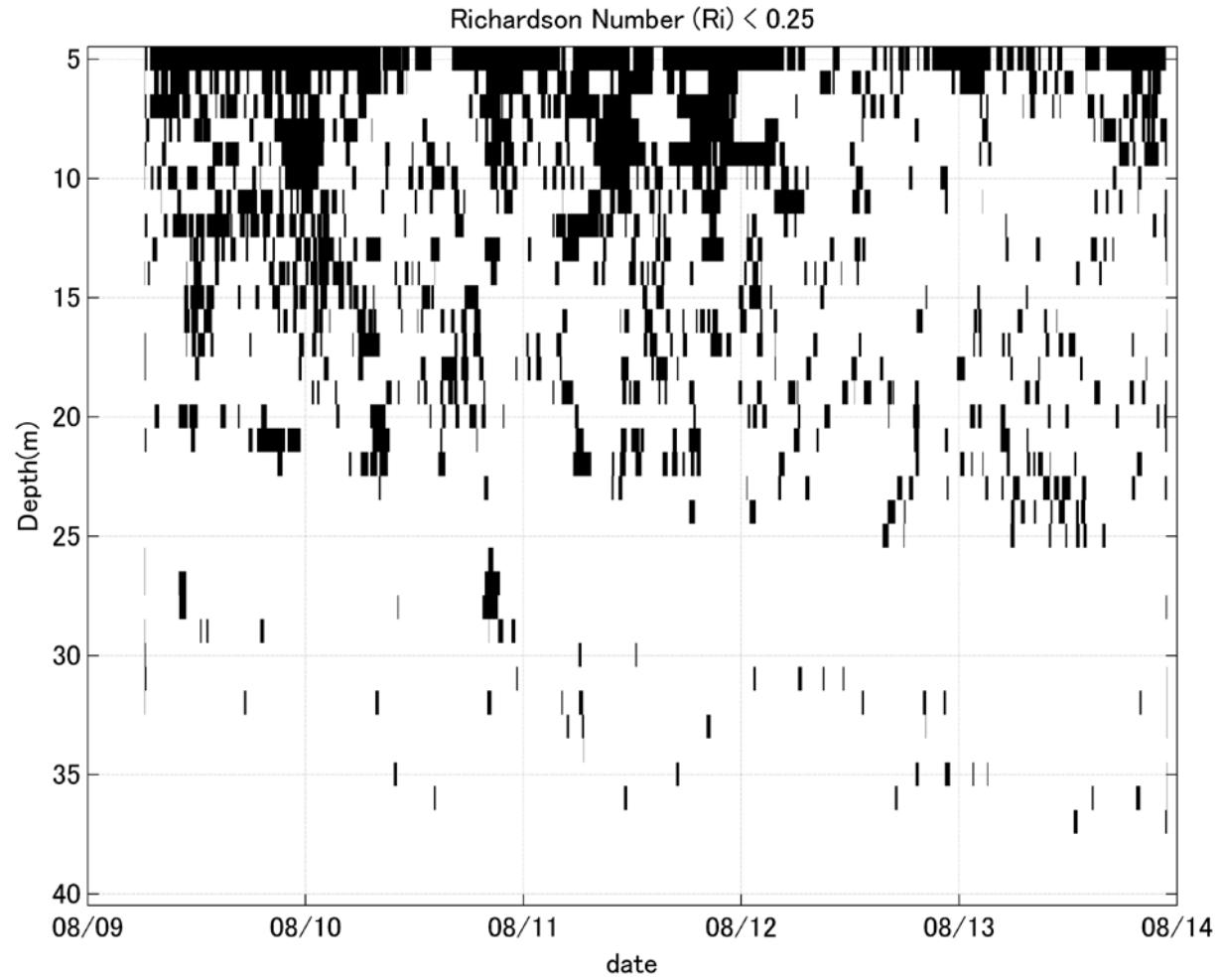
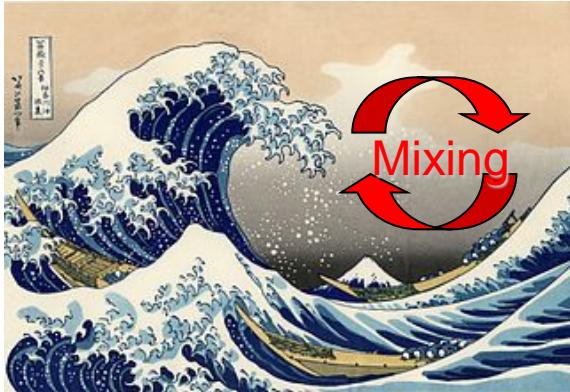


水上観測ステーションの漂流軌跡

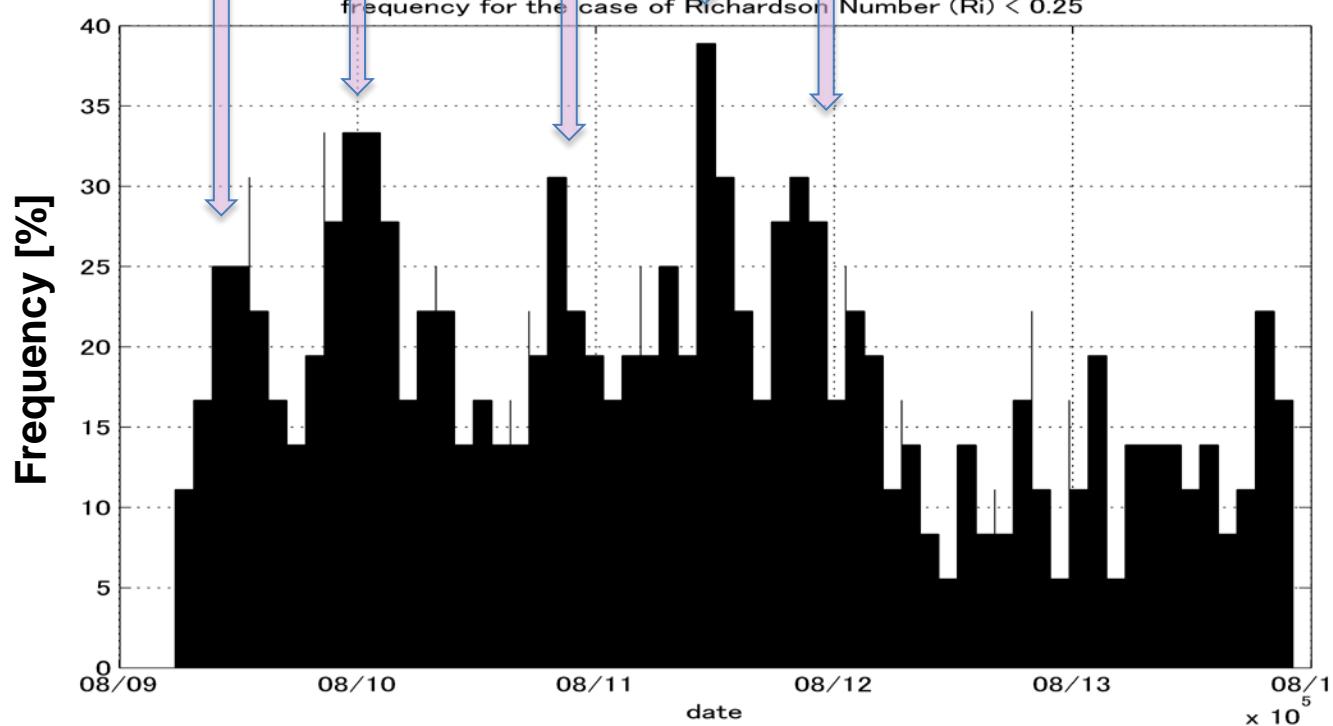
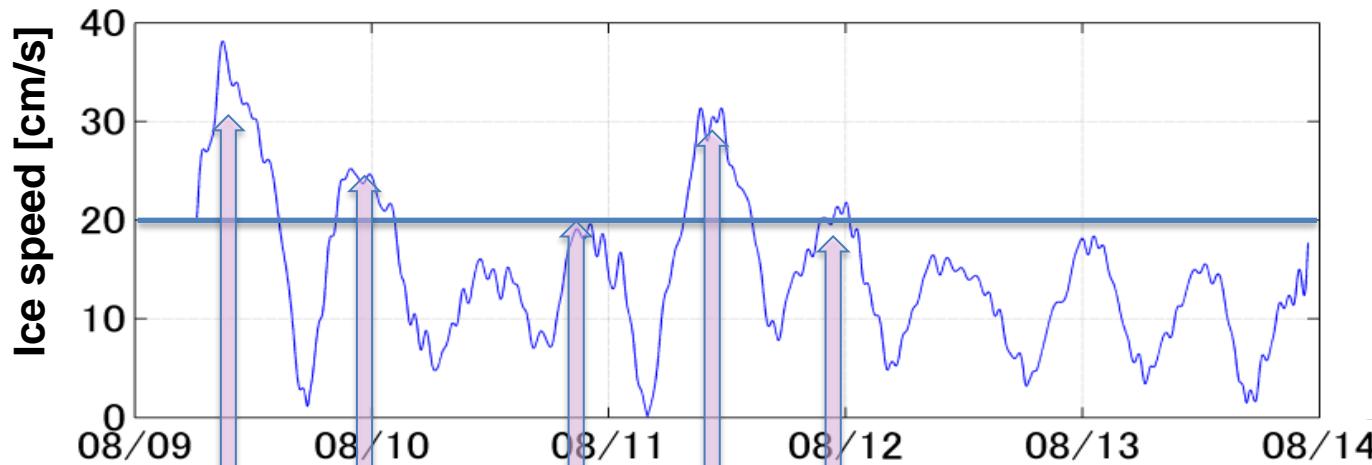


$$Ri = N^2/S^2$$

$$N^2 = -\frac{1}{\rho} \frac{\partial \rho}{\partial z}, \quad S^2 = \left(\frac{\partial u}{\partial z} \right)^2$$



Sea ice speed and Frequency of the case satisfying $Ri < 0.25$ (5-30m)

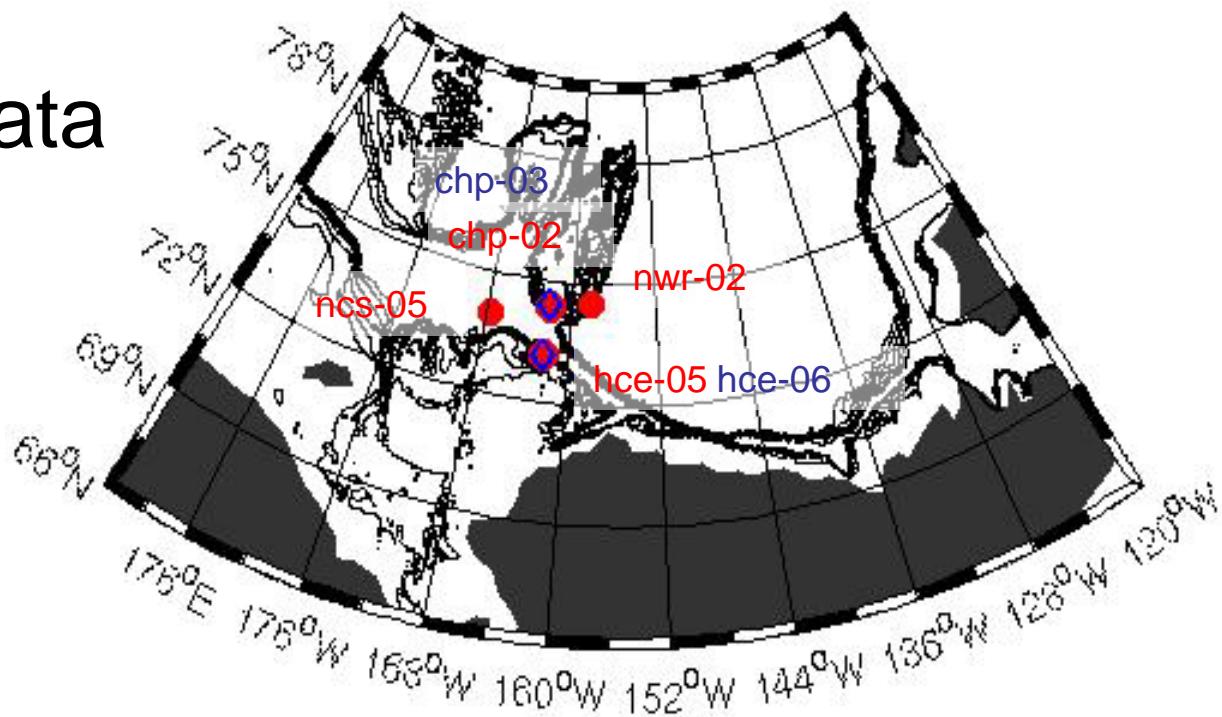


When sea ice speed is greater than 20cm/s, strong vertical mixing occurred.

The former study by Kawashima (2013) was meaningful.

Indirect evaluation of amplitude of inertial oscillation of sea ice from satellite data (AMSR-E)

Moored ADCP data



Chp-02 2002年 9月 3日 ~ 2003年 9月 17日

Chp-03 2003年 9月 3日 ~ 2004年 8月 17日

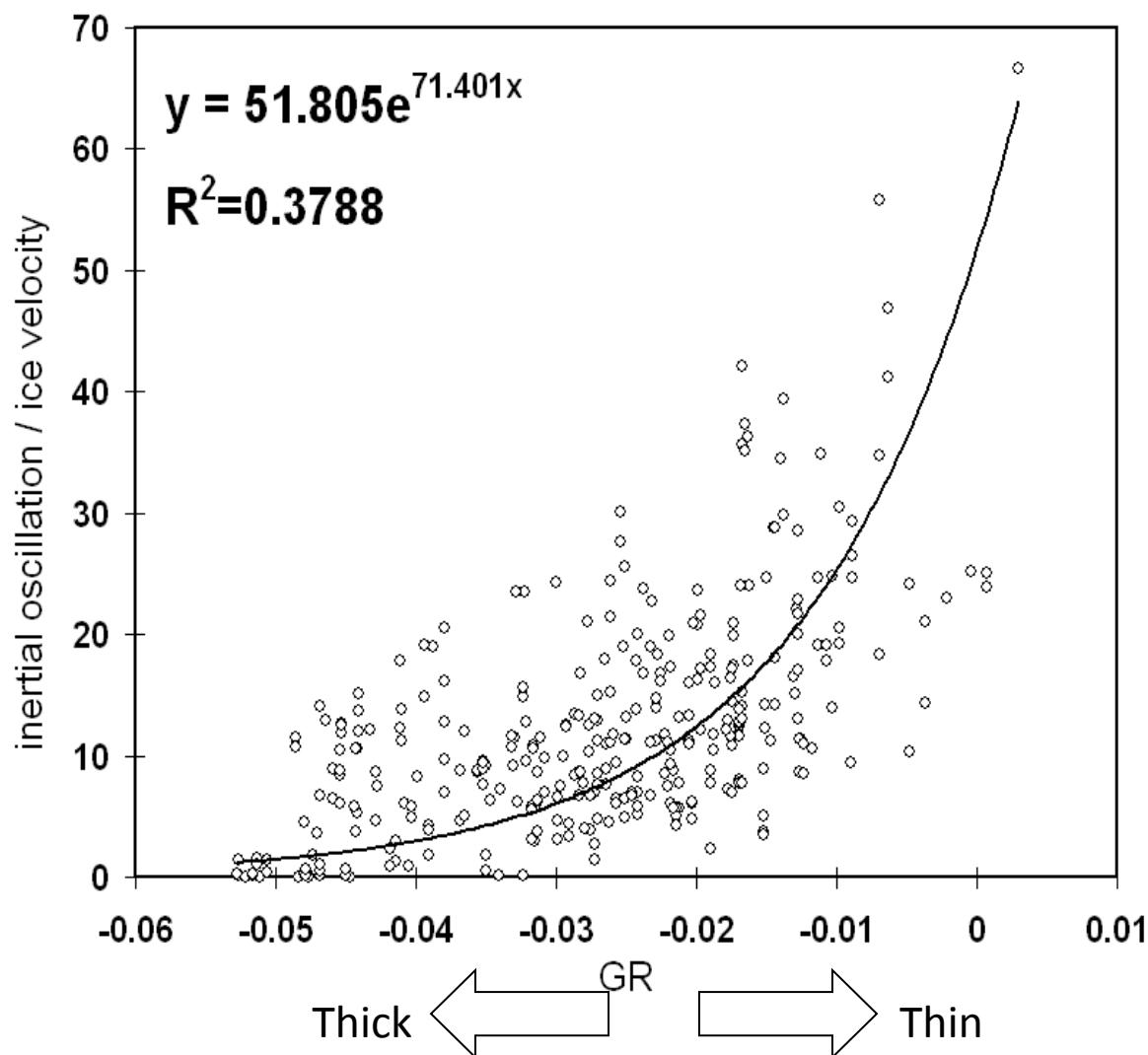
Hce-05 2005年10月 3日 ~ 2006年10月 3日

Hce-06 2006年10月 6日 ~ 2007年10月 6日

Ncs-05 2005年10月 5日 ~ 2006年10月 5日

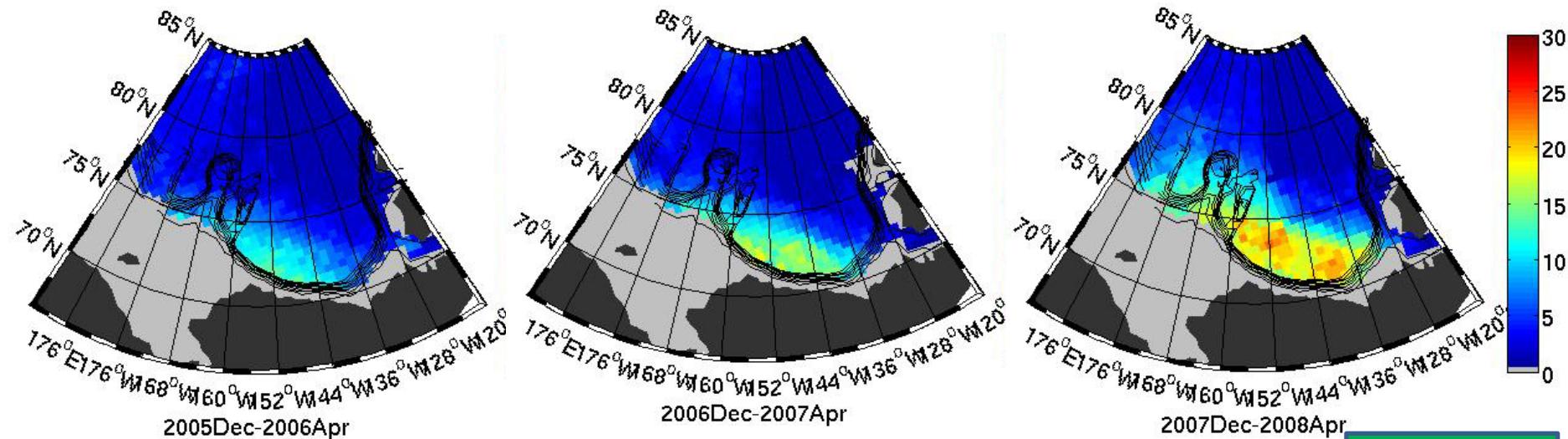
Nwr-03 2003年 8月 16日 ~ 2004年 9月 7日

[Amplitude of inertial oscillation] / [background low-passed sea ice velocity] depends on GR

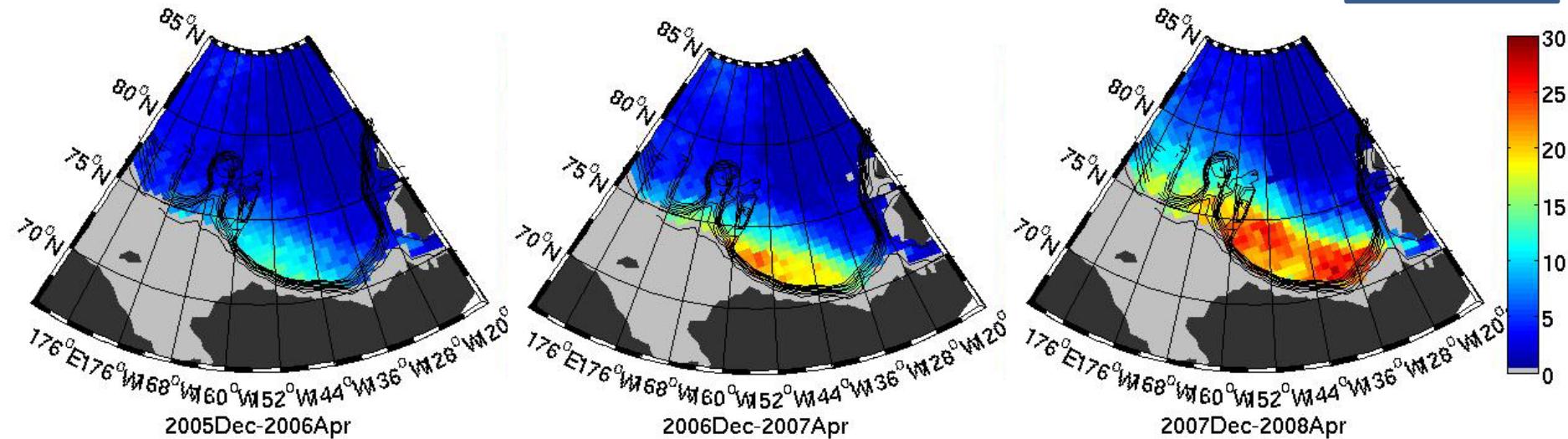


Plots for
Concentration > 95%
Wind speed > 6m/s
Ice speed > 20cm/s

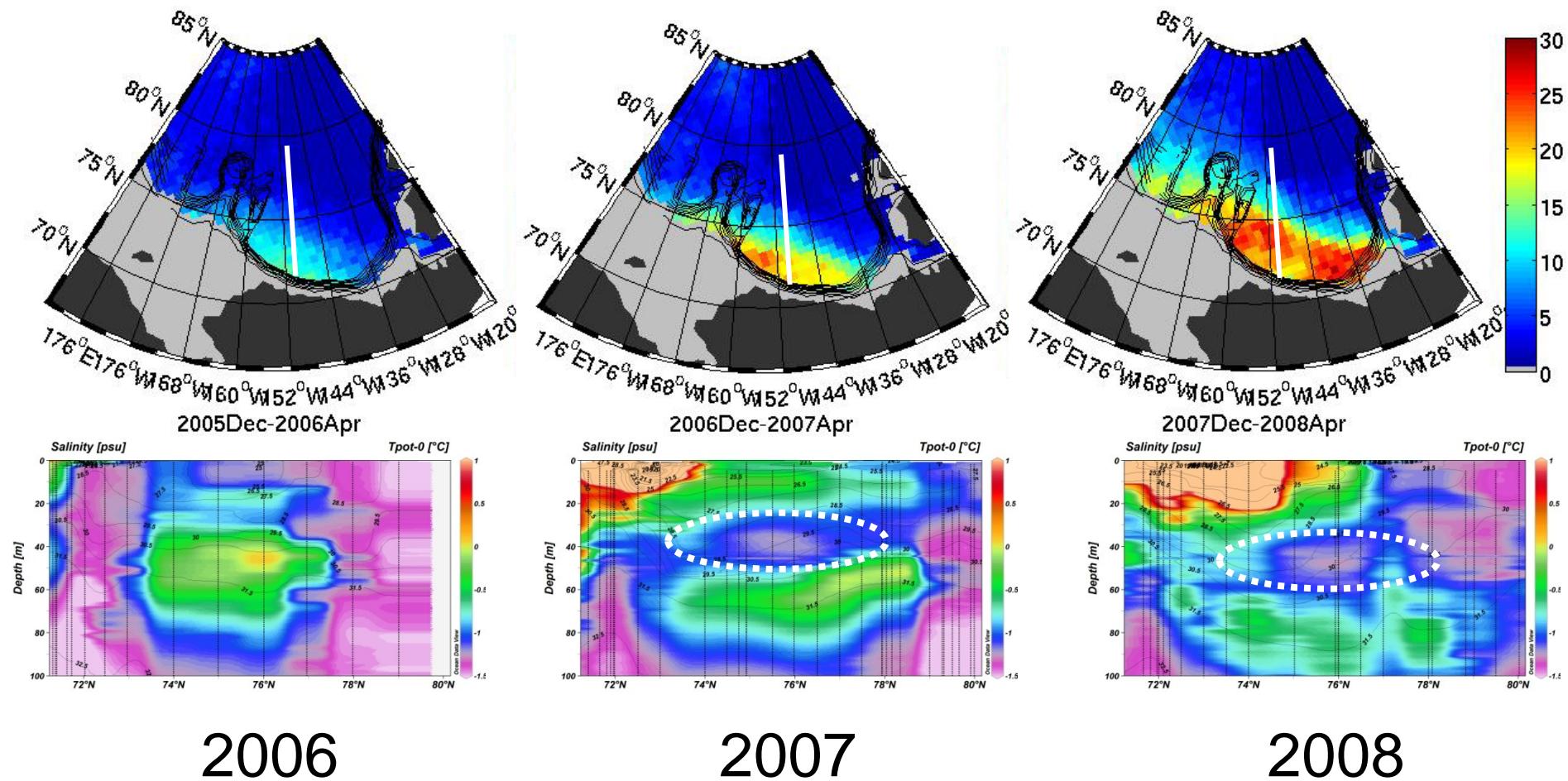
- Thickness and sea ice type control amplitude of inertial oscillation
- This regression curve enable us to estimate absolute sea ice speed including inertial oscillation

frequency (daily mean ice > 20cm/s)**frequency (daily mean ice + inertial oscillation > 20cm/s)**

Important number



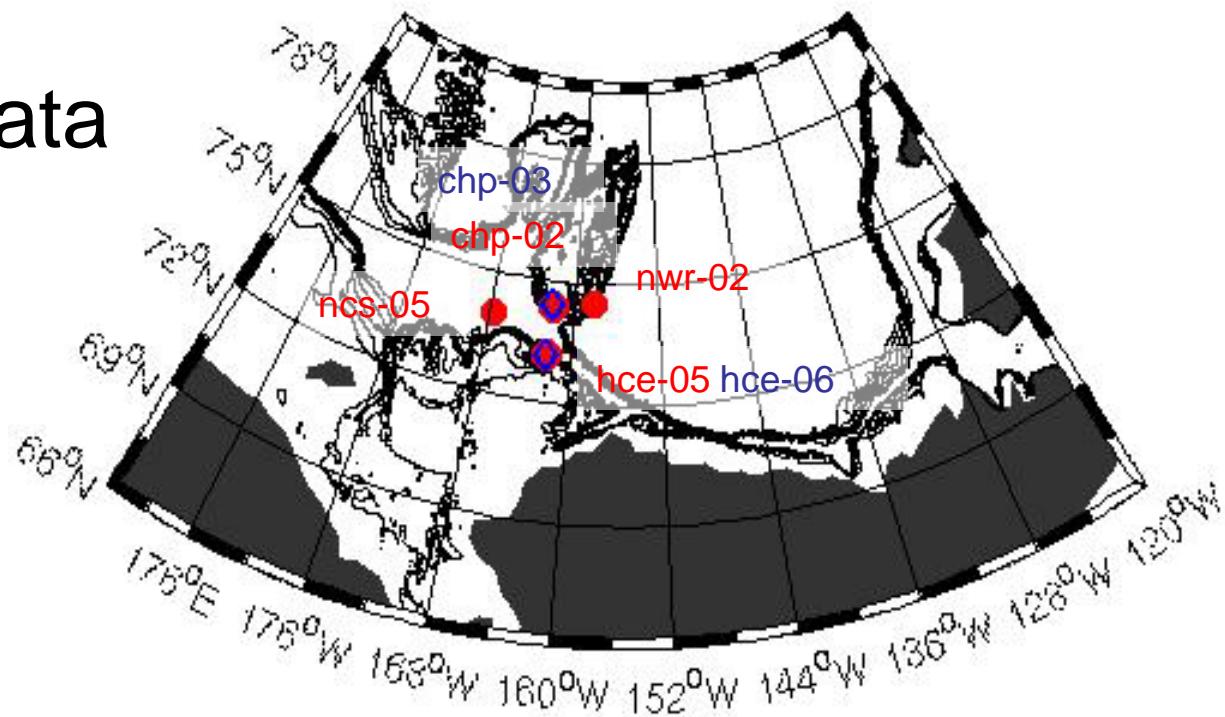
Interpretation of heat release event (=anomalous huge sea ice reduction event away from reduction trend) in 2007 and 2008



Shallow winter mixed layer (20m)
 Base depth of mixed layer is deeper than the
 general core depth of the warm Pacific
 Summer Water
 Gentle heat release

Deep winter mixed layer (50-70m)
 Base depth of mixed layer is deeper than the general core
 depth of the warm Pacific Summer Water
Active heat release \Rightarrow anomalous event-like reduction of sea ice

Moored ADCP data



Chp-02 2002年 9月 3日 ~ 2003年 9月 17日

Chp-03 2003年 9月 3日 ~ 2004年 8月 17日

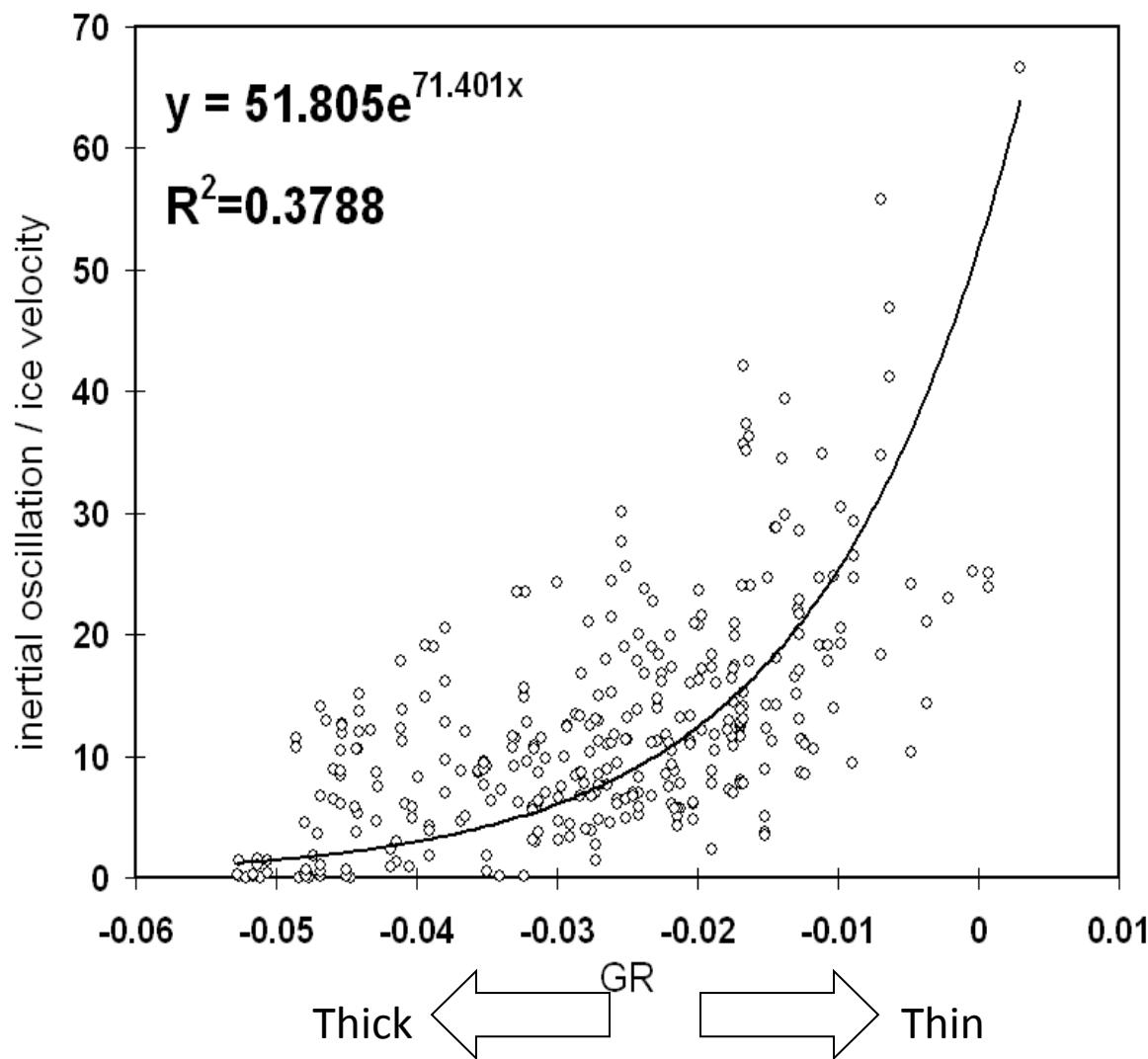
Hce-05 2005年10月 3日 ~ 2006年10月 3日

Hce-06 2006年10月 6日 ~ 2007年10月 6日

Ncs-05 2005年10月 5日 ~ 2006年10月 5日

Nwr-03 2003年 8月 16日 ~ 2004年 9月 7日

[Amplitude of inertial oscillation] / [background low-passed sea ice velocity] depends on GR

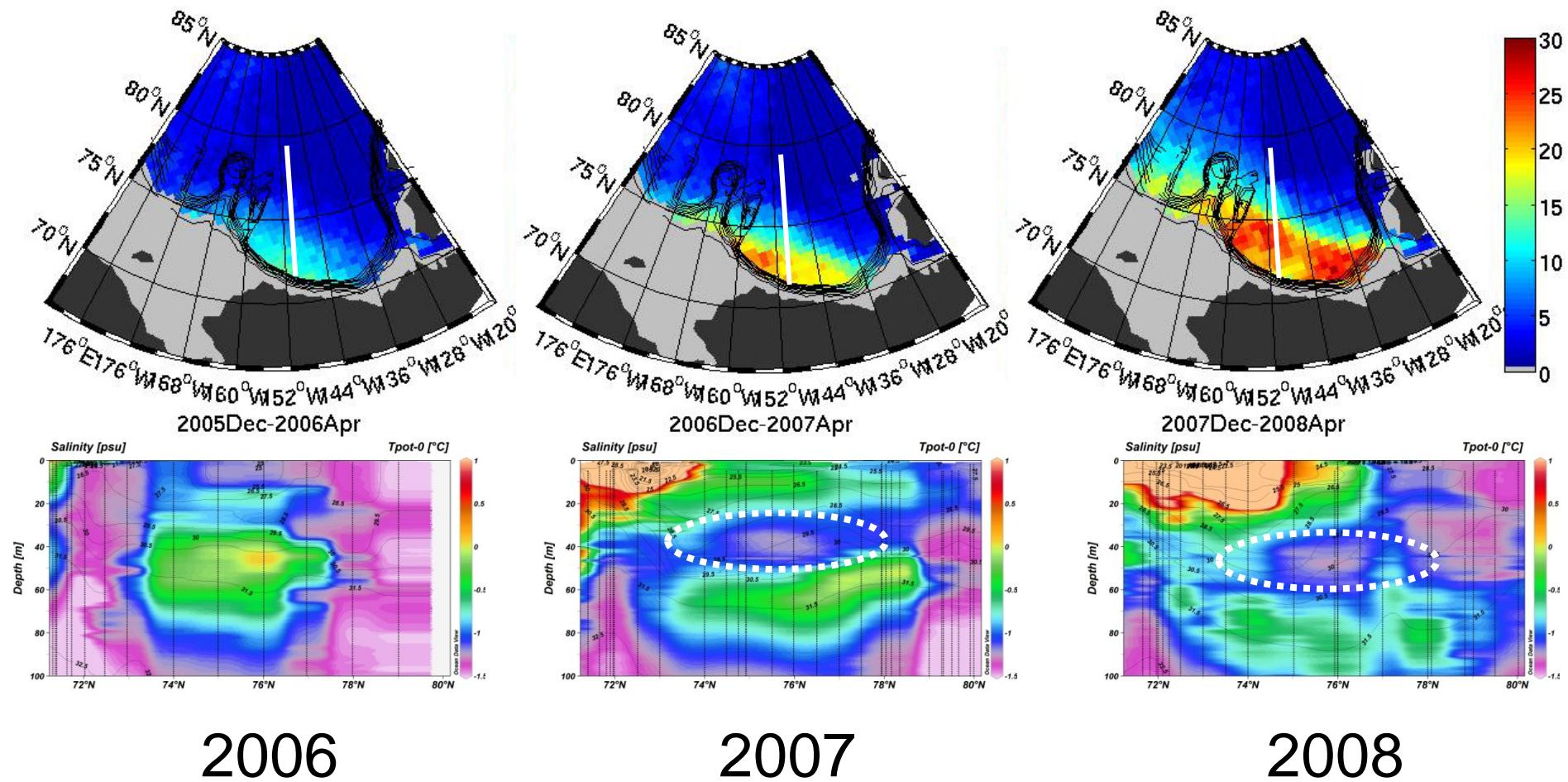


$$GR = \frac{[36V] - [18V]}{[36V] + [18V]}.$$

Plots for
Concentration > 95%
Wind speed > 6m/s
Ice speed > 20cm/s

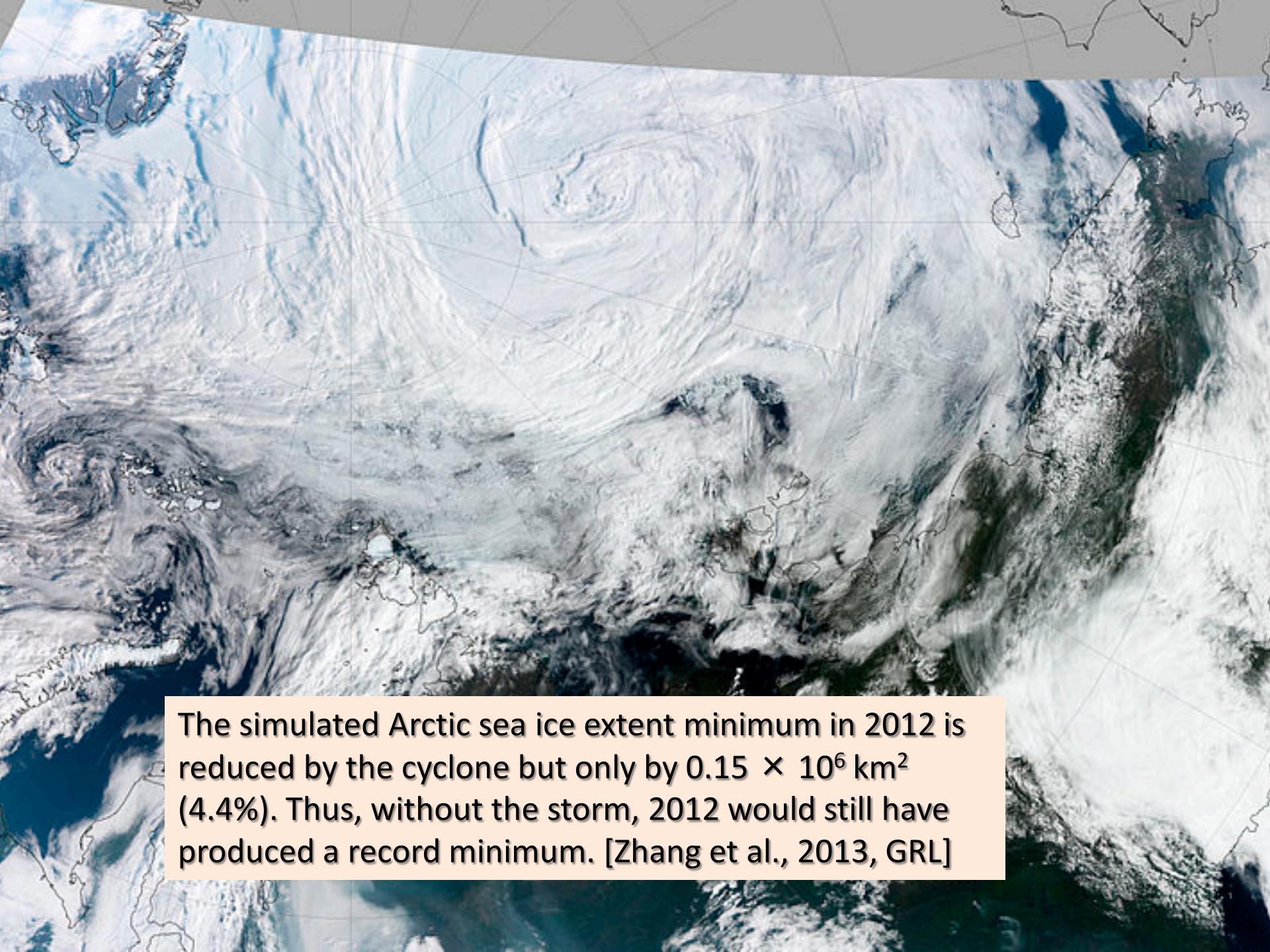
- Thickness and sea ice type control amplitude of inertial oscillation
- This regression curve enable us to estimate absolute sea ice speed including inertial oscillation

Interpretation of heat release event (=anomalous huge sea ice reduction event away from reduction trend) in 2007 and 2008



Shallow winter mixed layer (20m)
Base depth of mixed layer is deeper than the
general core depth of the warm Pacific
Summer Water
Gentle heat release

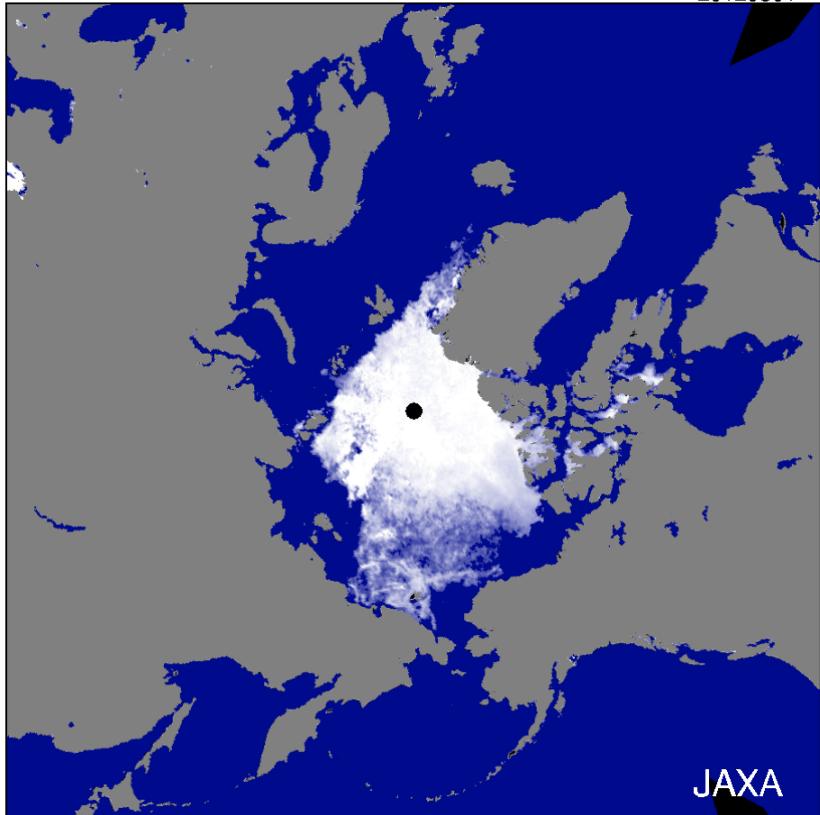
Deep winter mixed layer (50-70m)
Base depth of mixed layer is deeper than the general core
depth of the warm Pacific Summer Water
**Active heat release \Rightarrow anomalous event-like reduction of
sea ice**



The simulated Arctic sea ice extent minimum in 2012 is reduced by the cyclone but only by $0.15 \times 10^6 \text{ km}^2$ (4.4%). Thus, without the storm, 2012 would still have produced a record minimum. [Zhang et al., 2013, GRL]

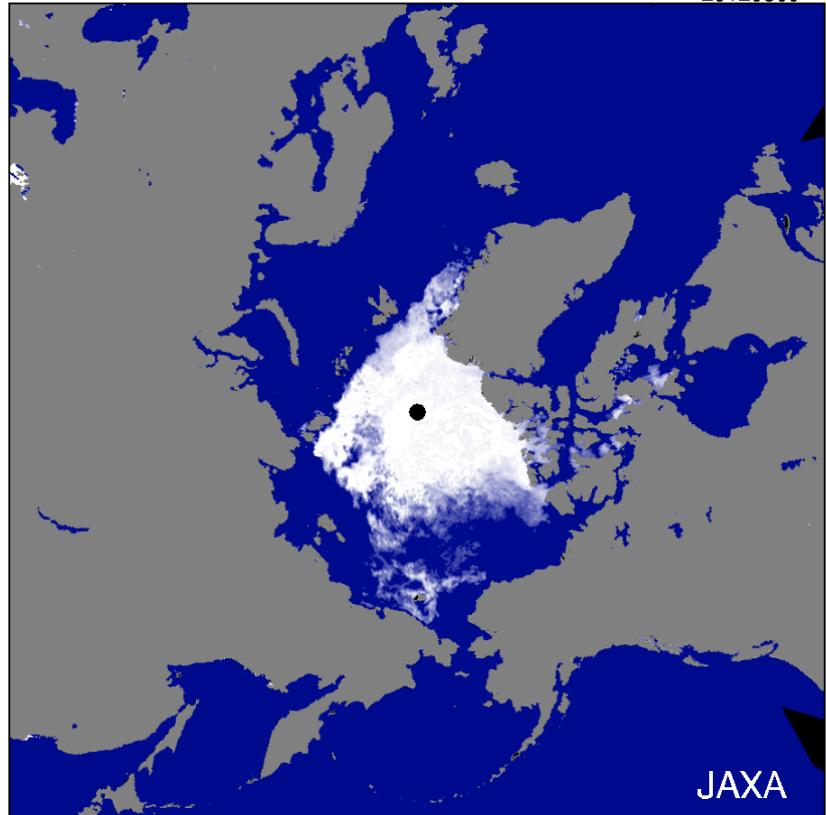
AMSR2 Sea Ice Concentration

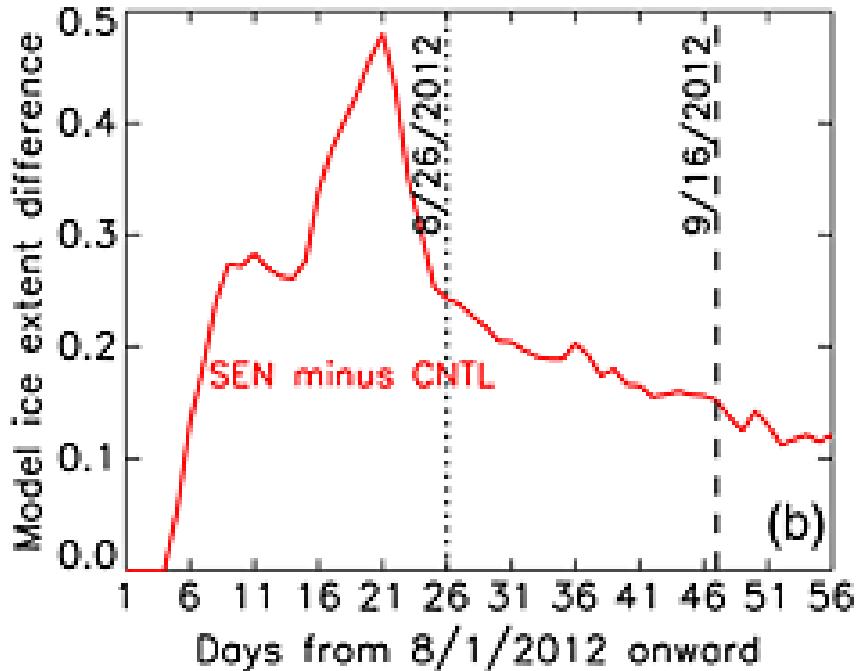
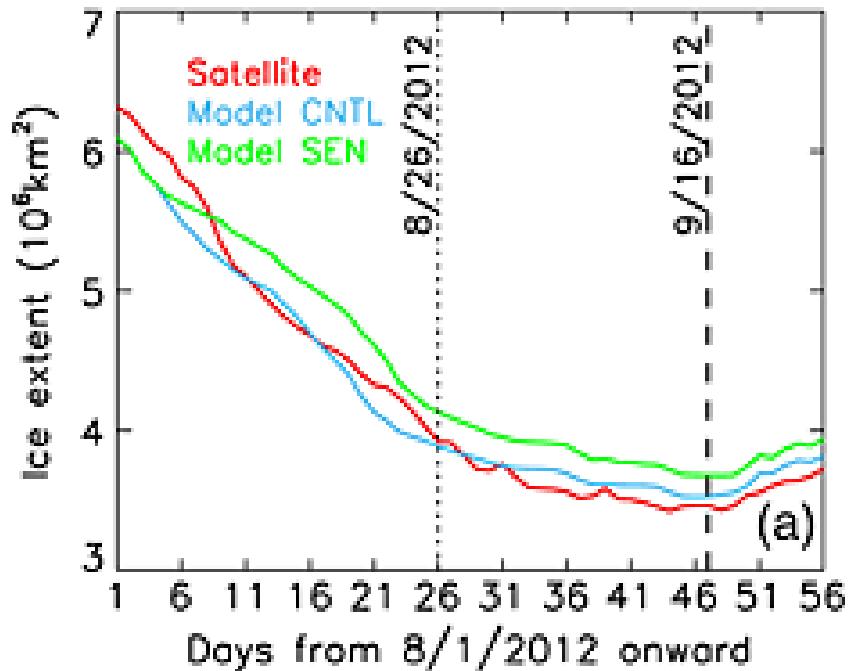
20120801



AMSR2 Sea Ice Concentration

20120809





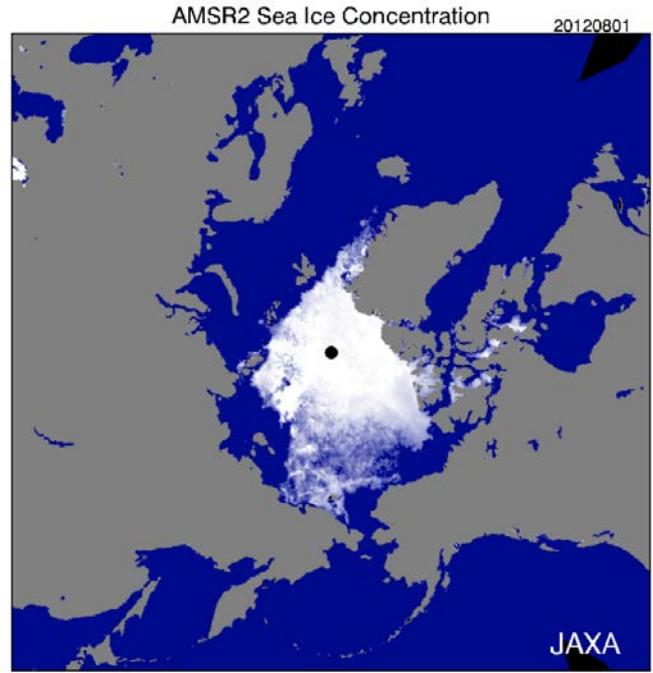
At a glance, cyclone seems to be important impression, but the influence is not so large.
Another processes is much significant to understand the sea ice decline and spatial pattern of sea ice distribution.



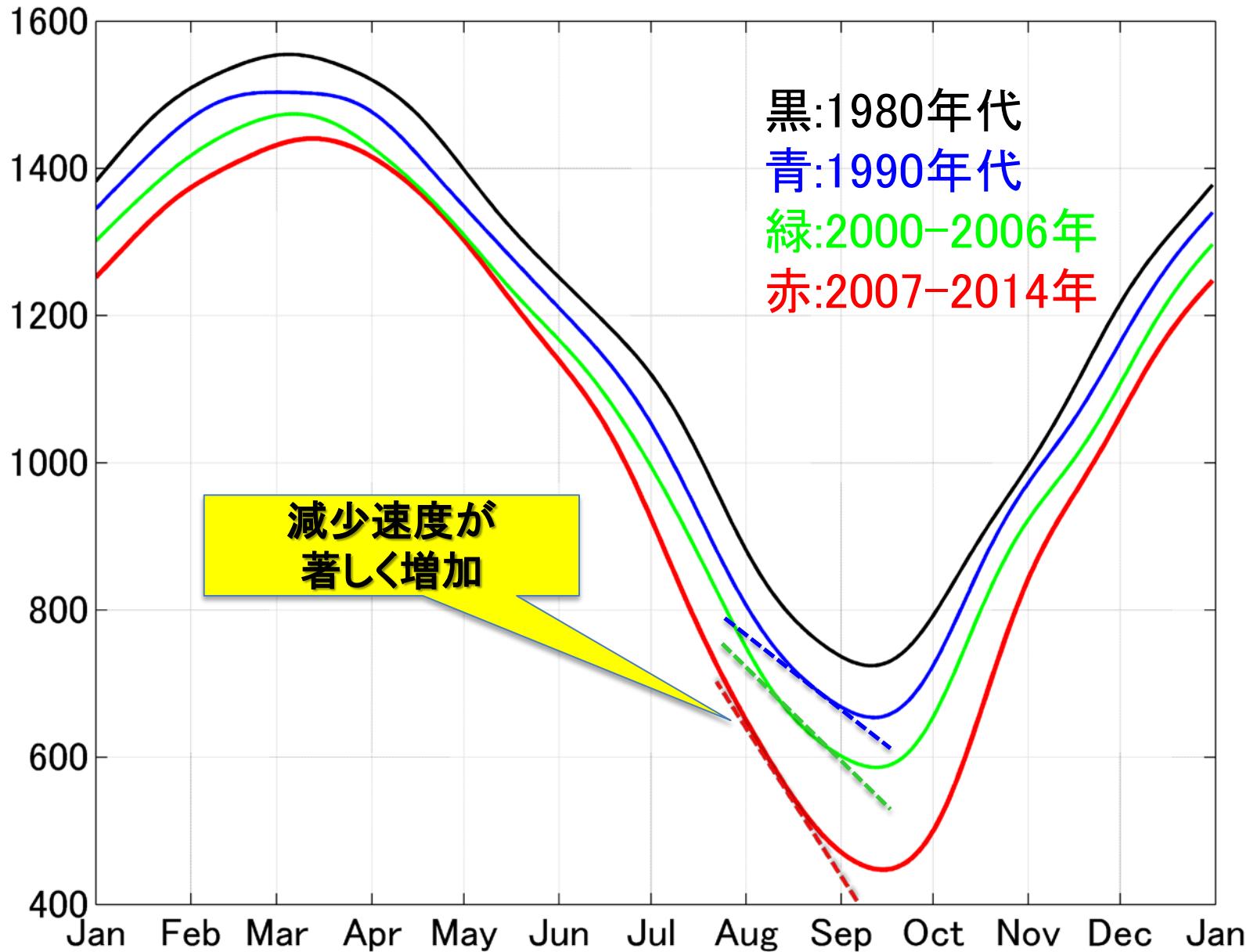
Precondition Changes in melt pond



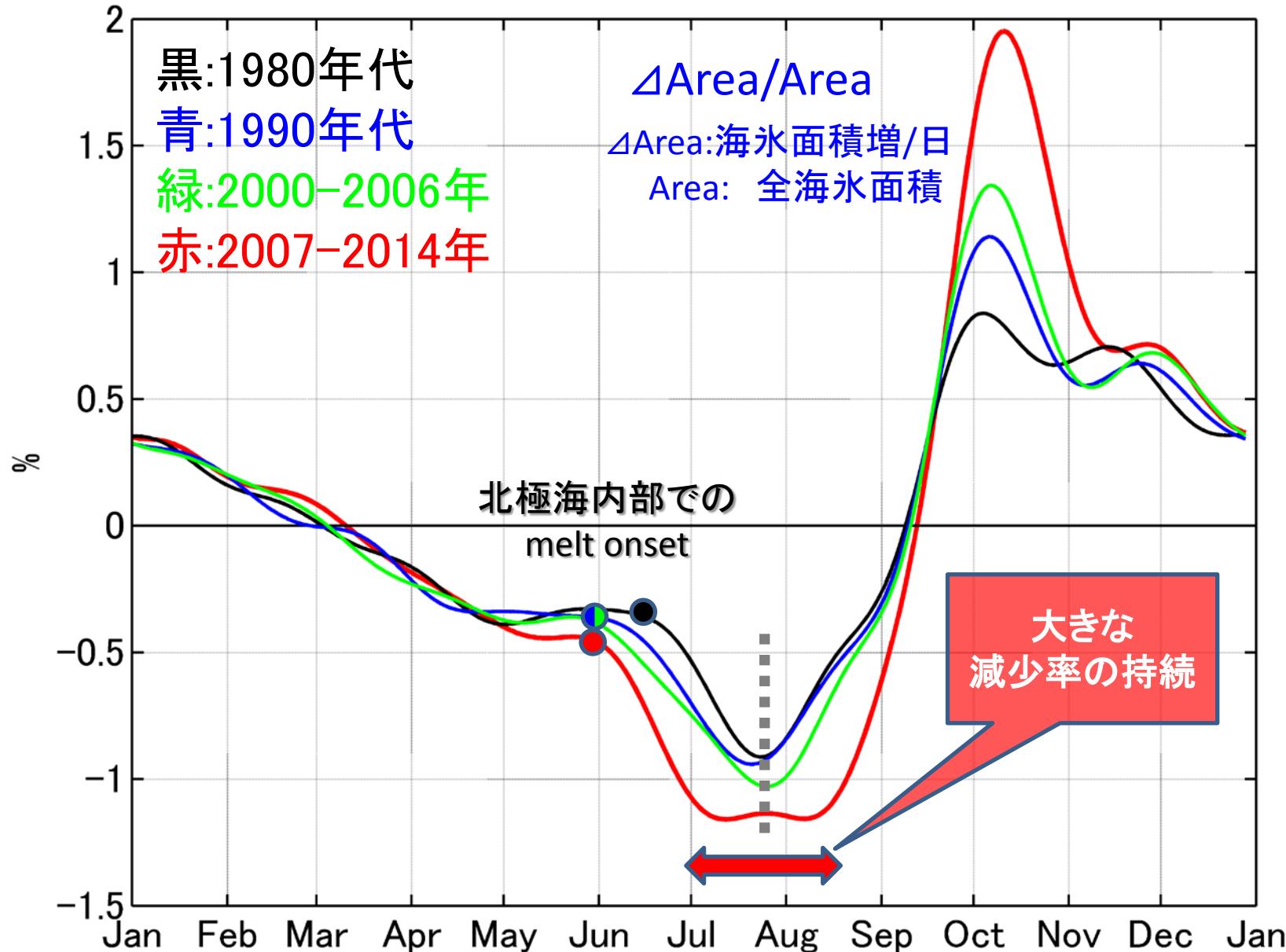
SHEBA1998



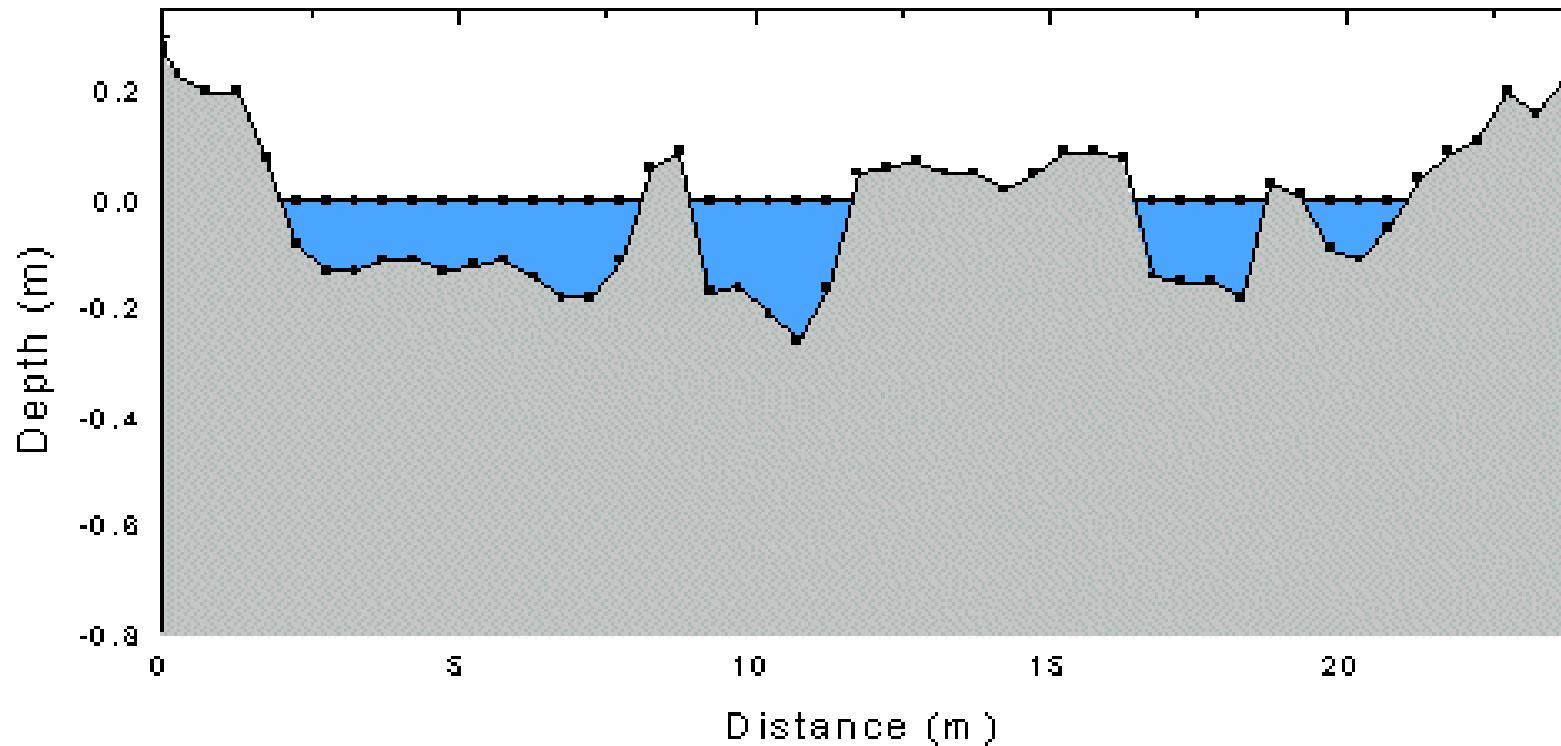
Sea Ice Extent [$\times 10^4 \text{ km}^3$]



Increasing rate of sea ice extent to total sea ice extent [%]



10 July

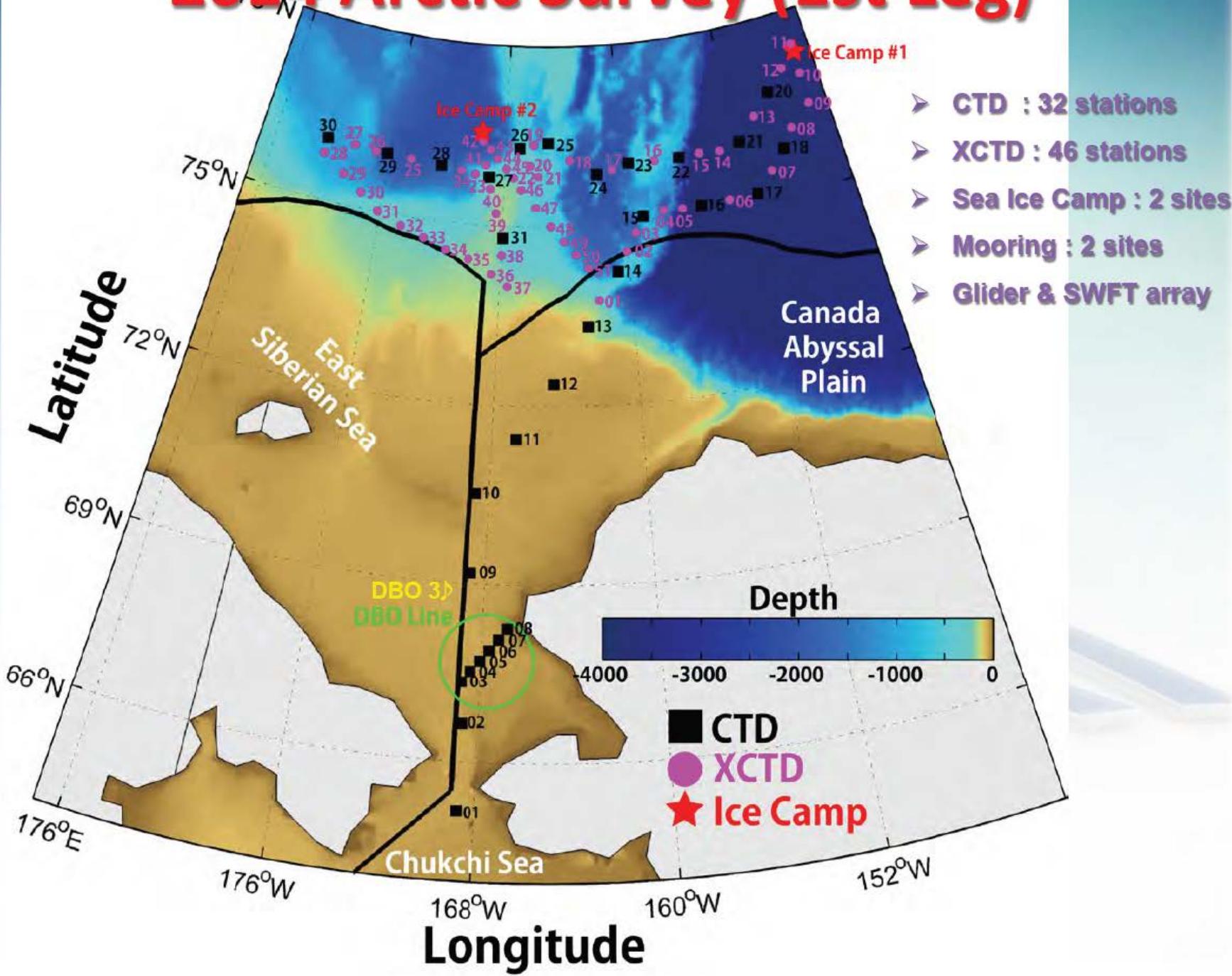


SHEBA 1998
多年氷上のmeltpond の発展

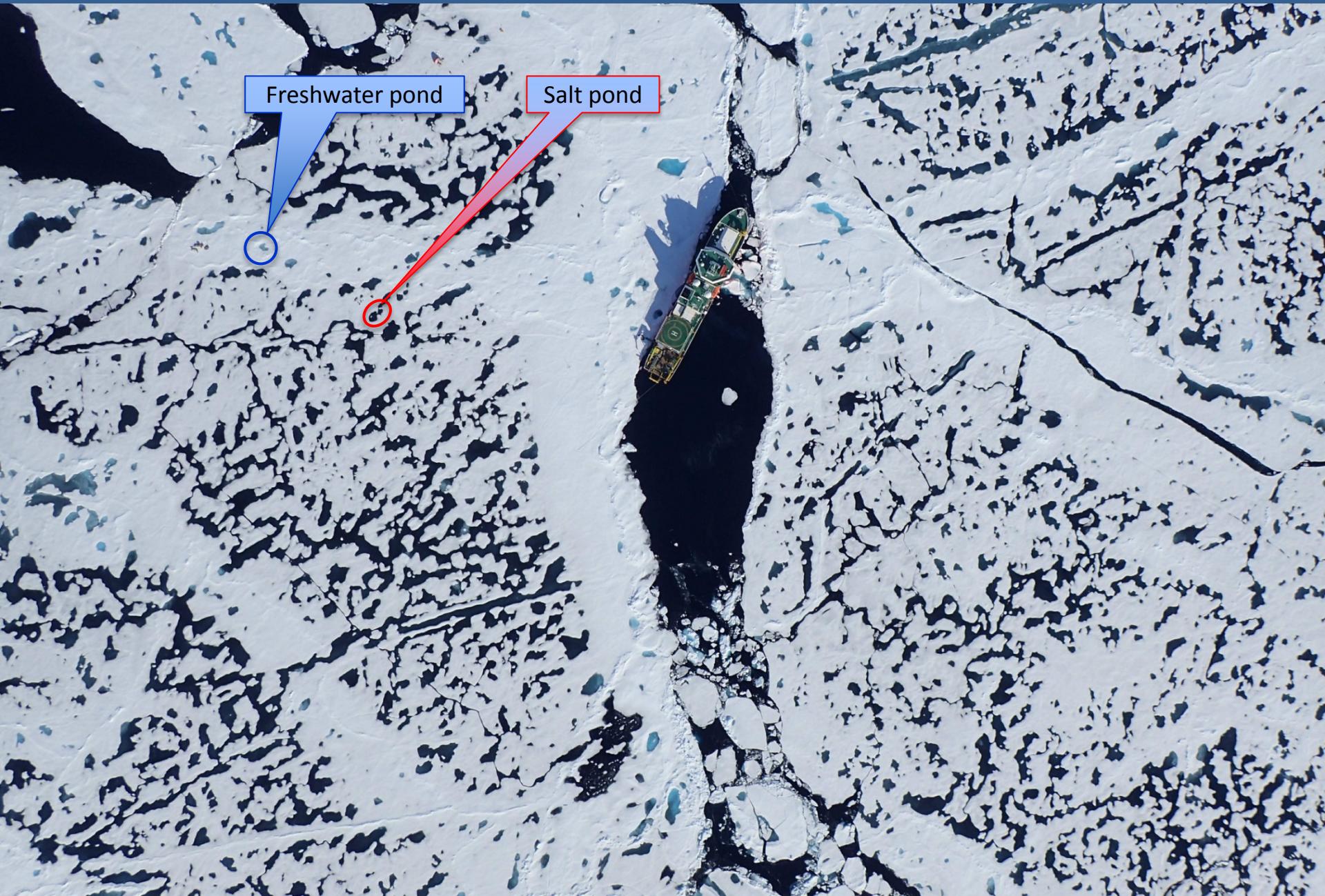
面積変化が無い
⇒ 底融解が卓越

SHEBA HP より

2014 Arctic Survey (1st Leg)



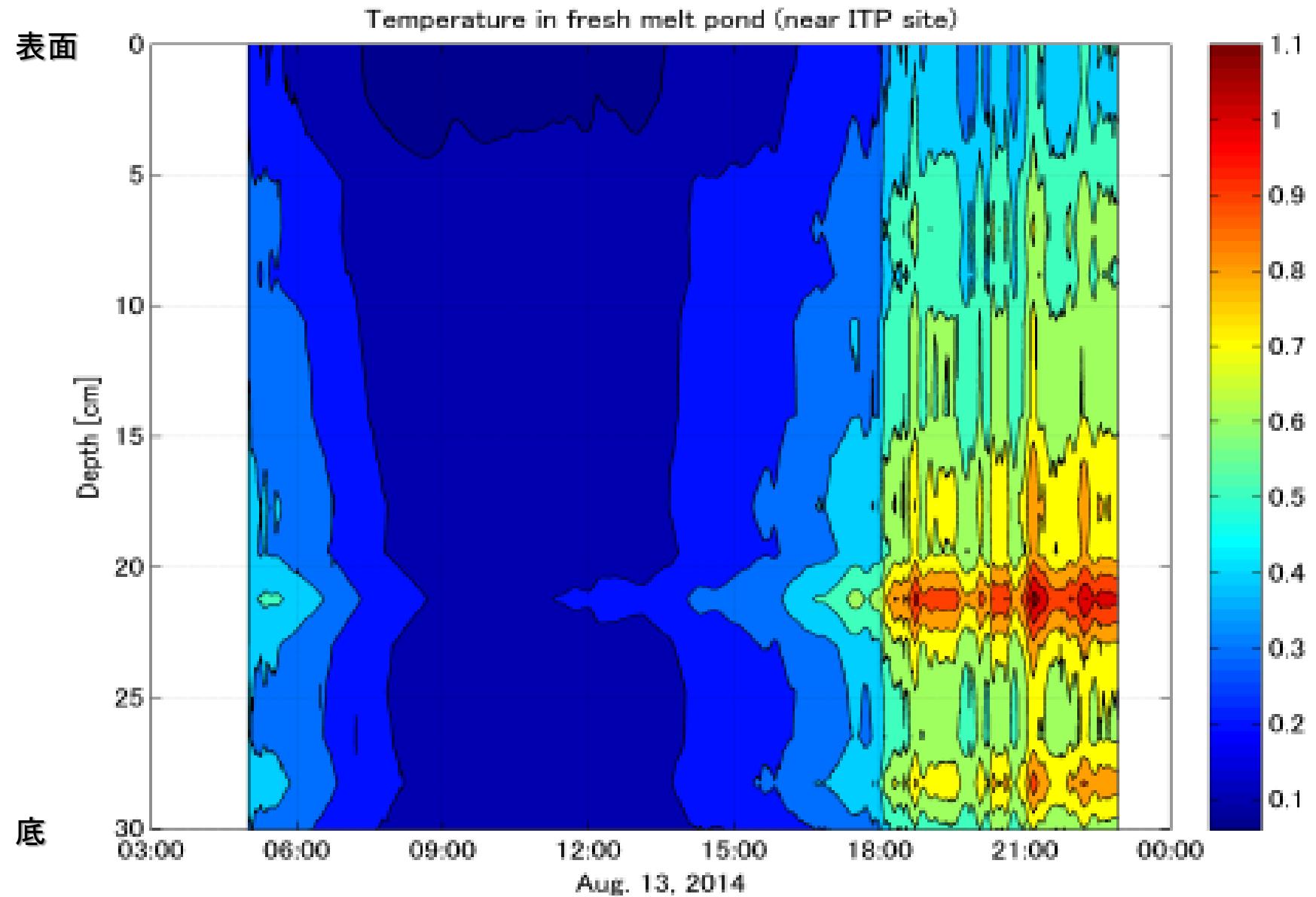
メルトポンドの再考【氷上観測】: 一年氷のメルトポンドは何故拡大、融解が速いのか？

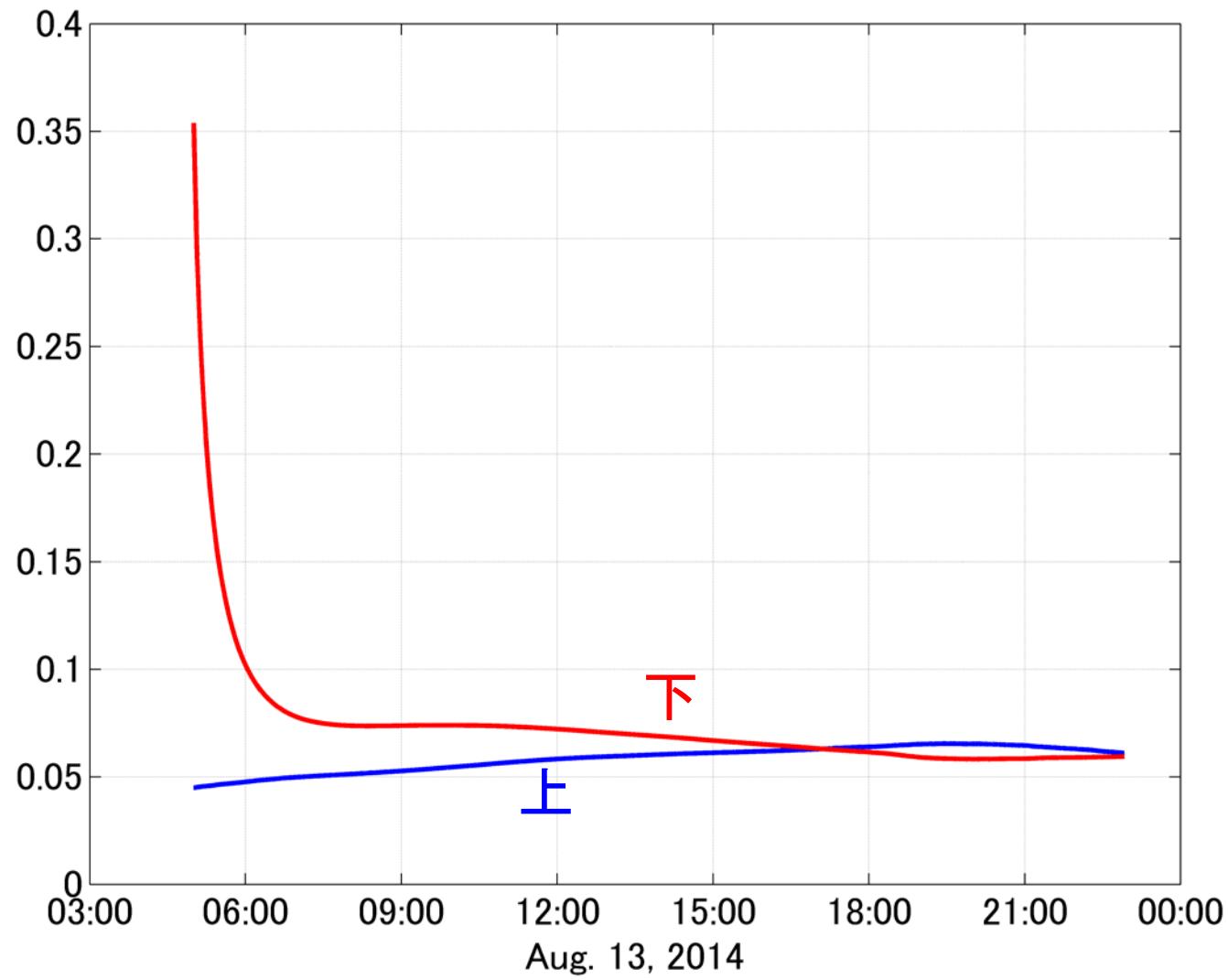




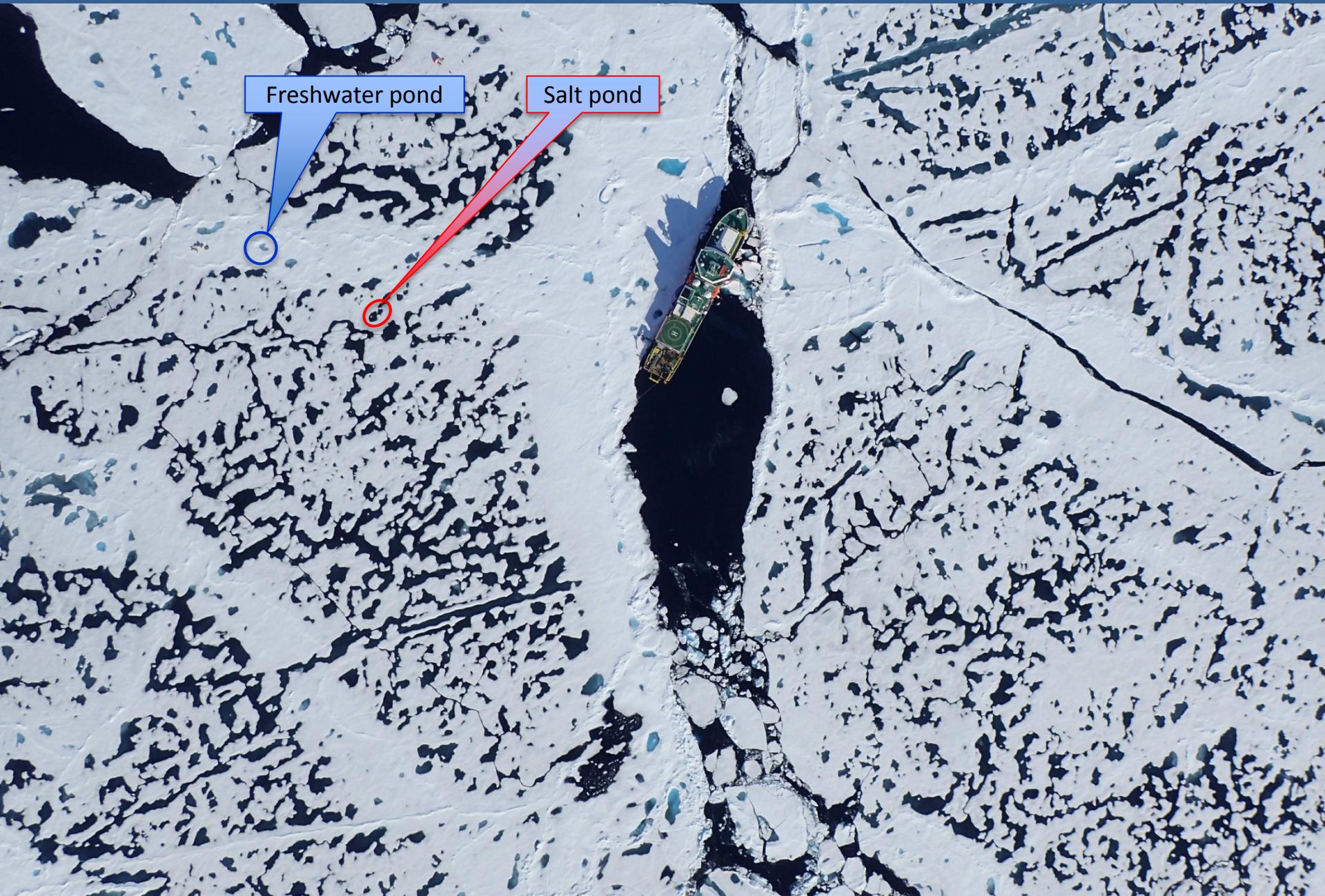


淡水ポンド





メルトポンドの再考【氷上観測】: 一年氷のメルトポンドは何故拡大、融解が速いのか？





ARAÇÁ
948

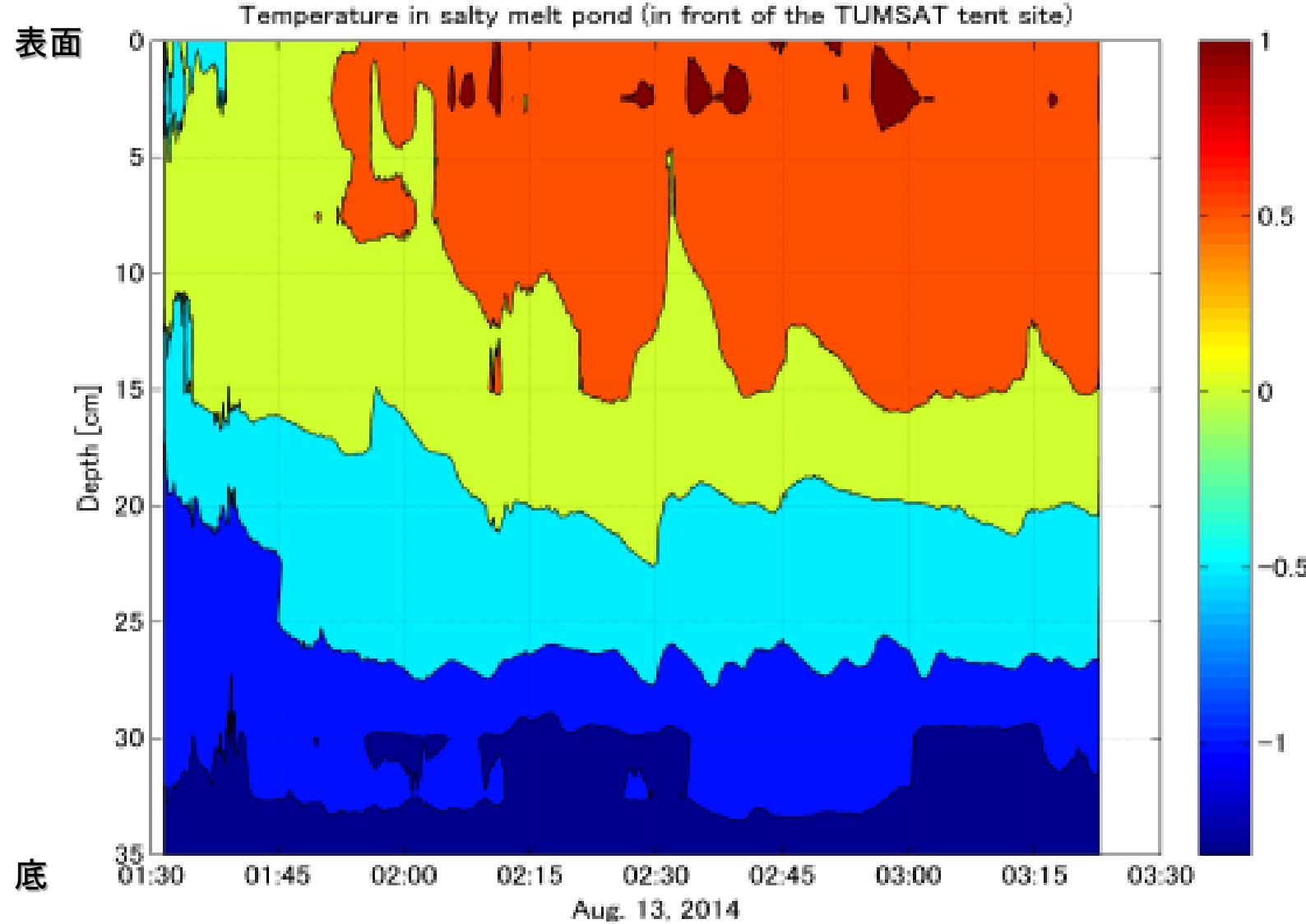
KOPRI

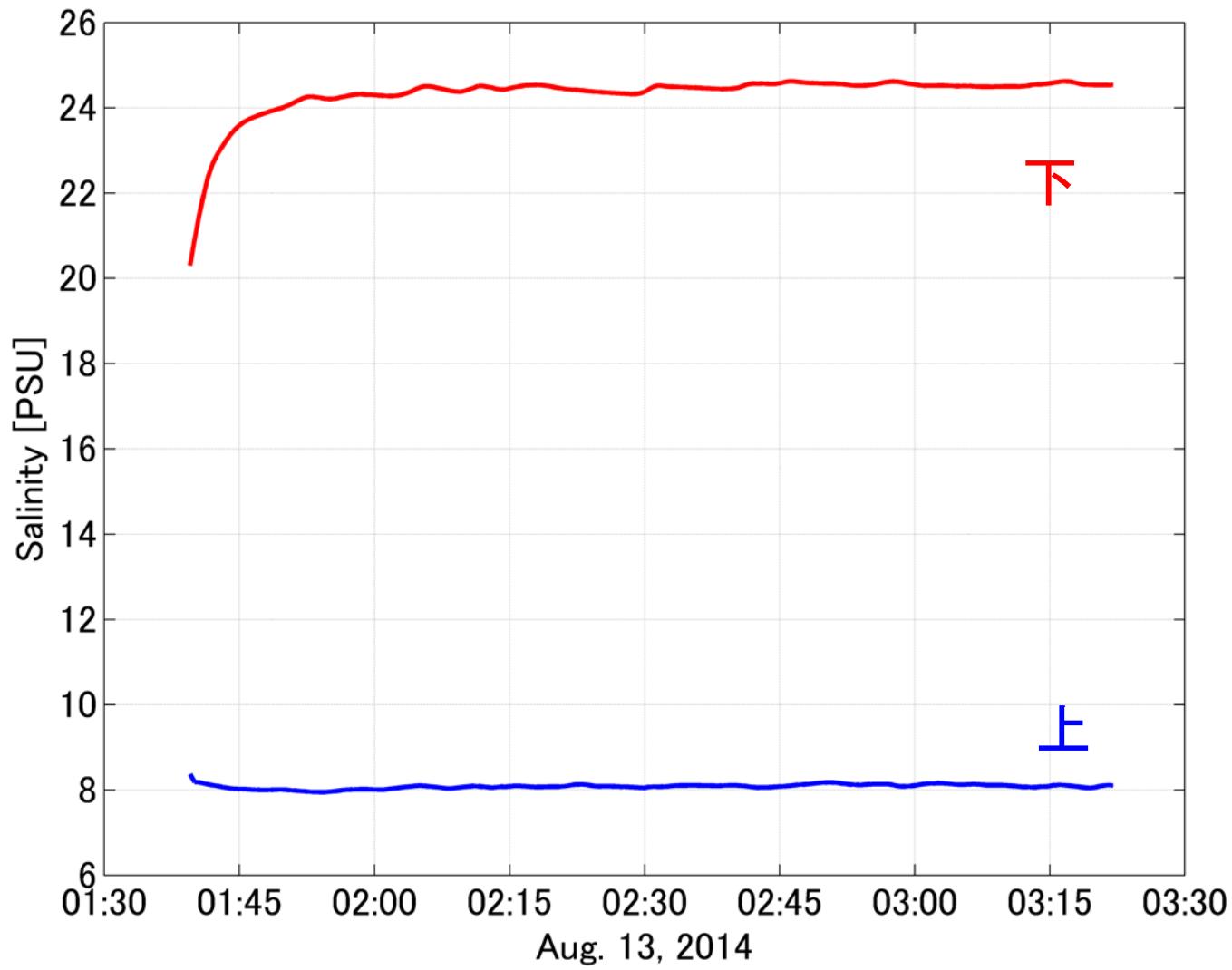


SBE
SEA-BIRD
ELECTRONICS INC.
VANE
WEATHERHOG

www.seabird.com

塩ポンド

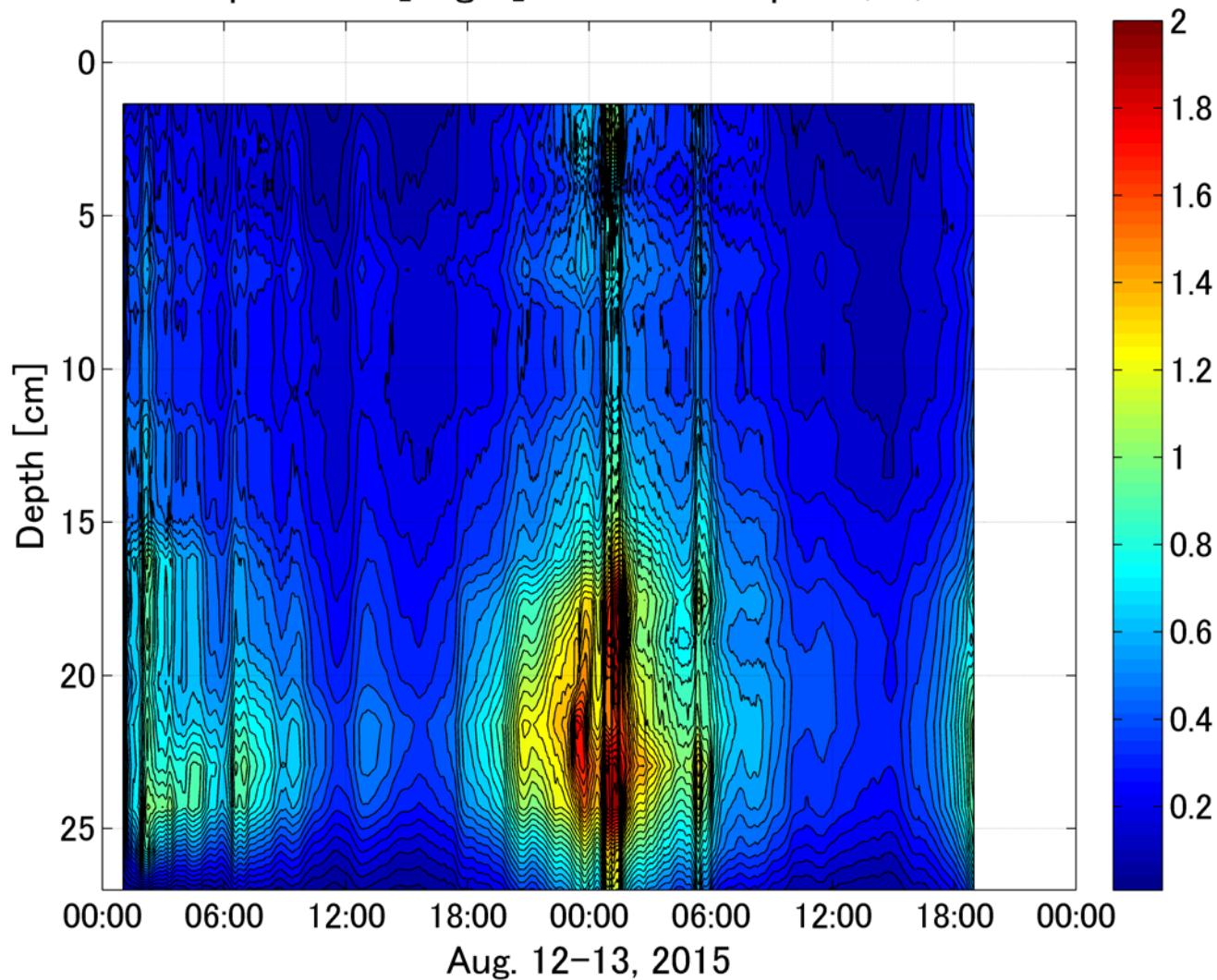






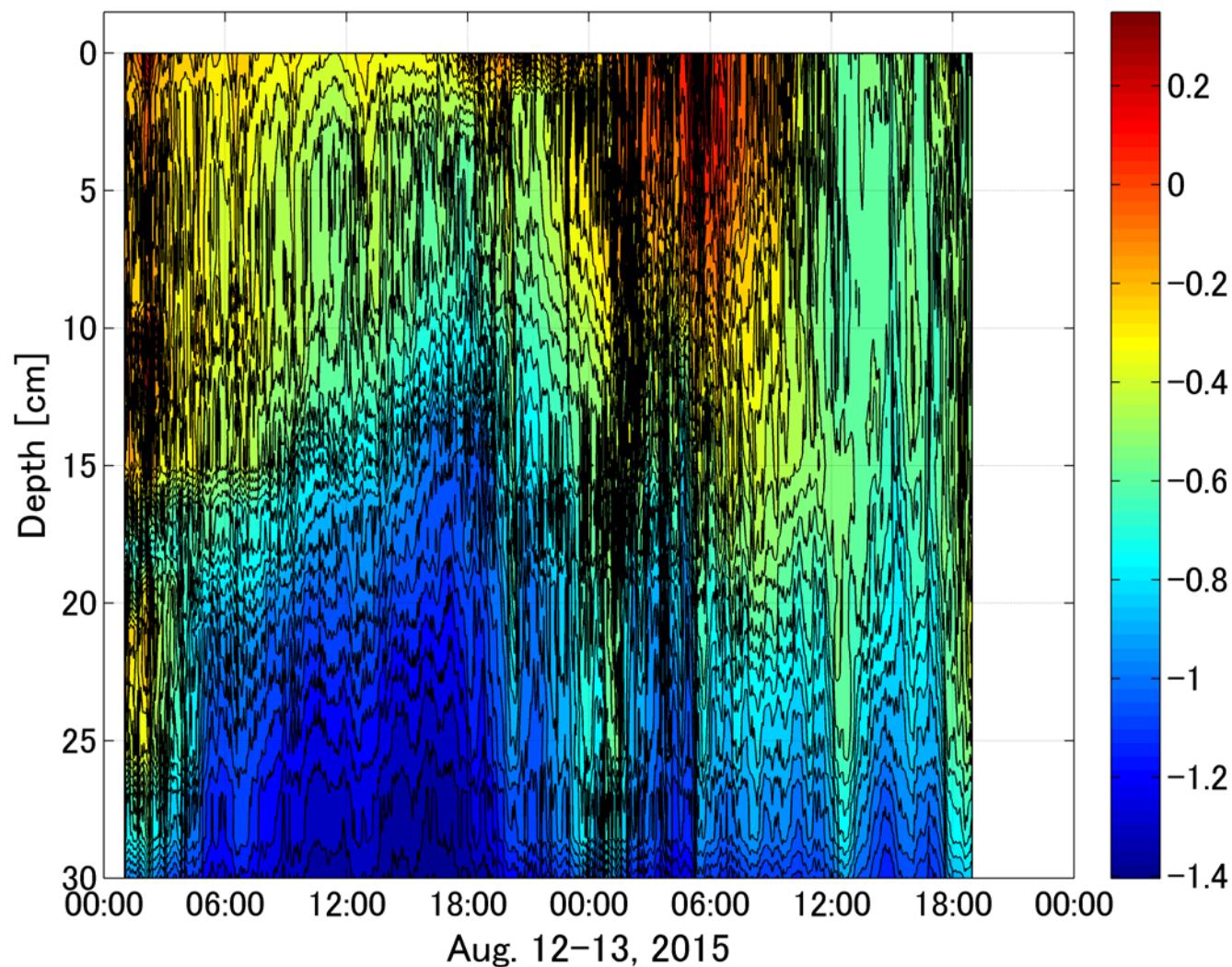


Temperature [deg.C]/ freshwater pond(#1) 2015





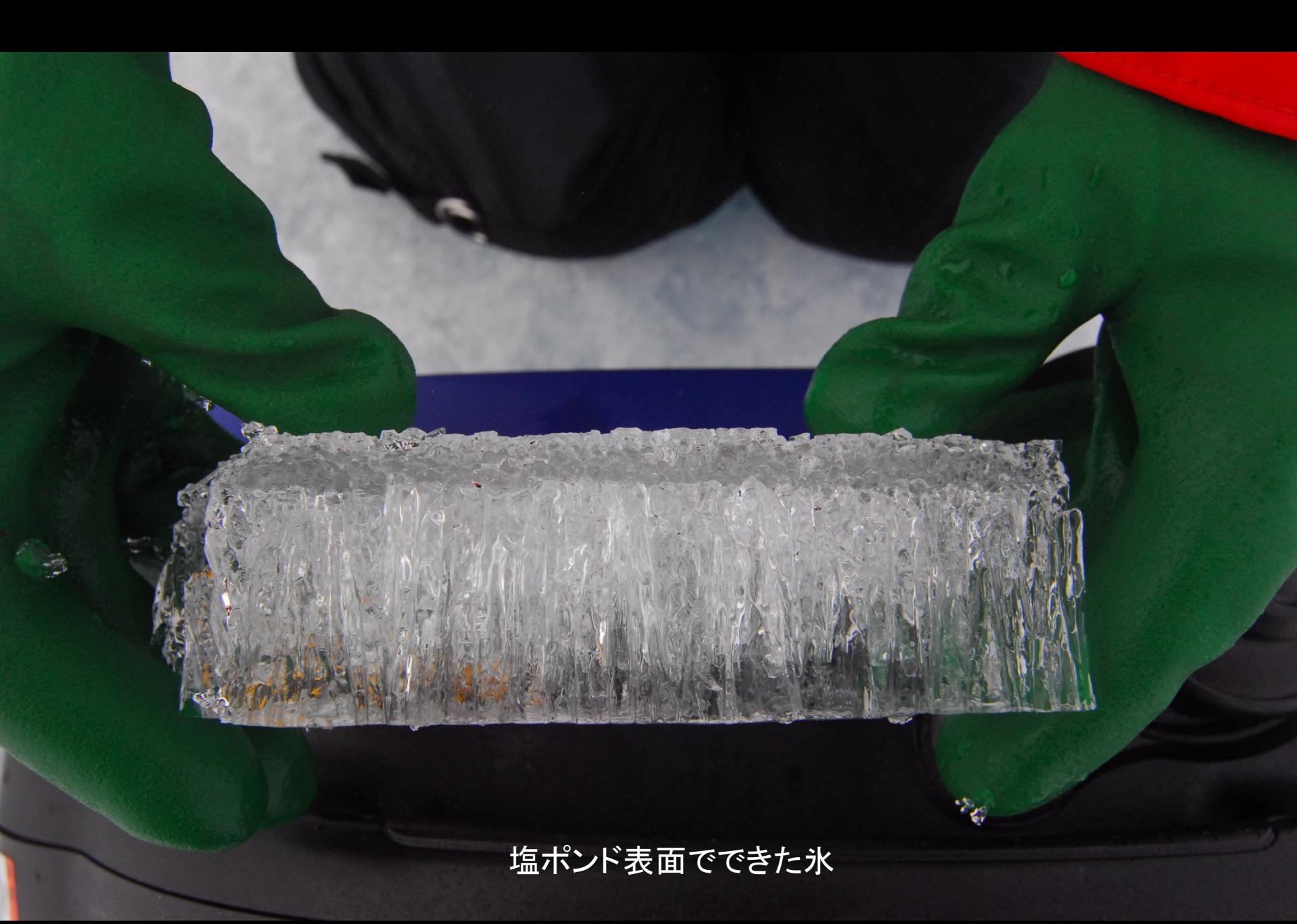
Temperature [deg.C]/ salt pond(#2) 2015





メルトポンドの中から見る
メルトポンド(一年氷)
ブラインチャンネルの名残
⇒メルトポンド内への塩水の浸透





塩ポンド表面でできた氷



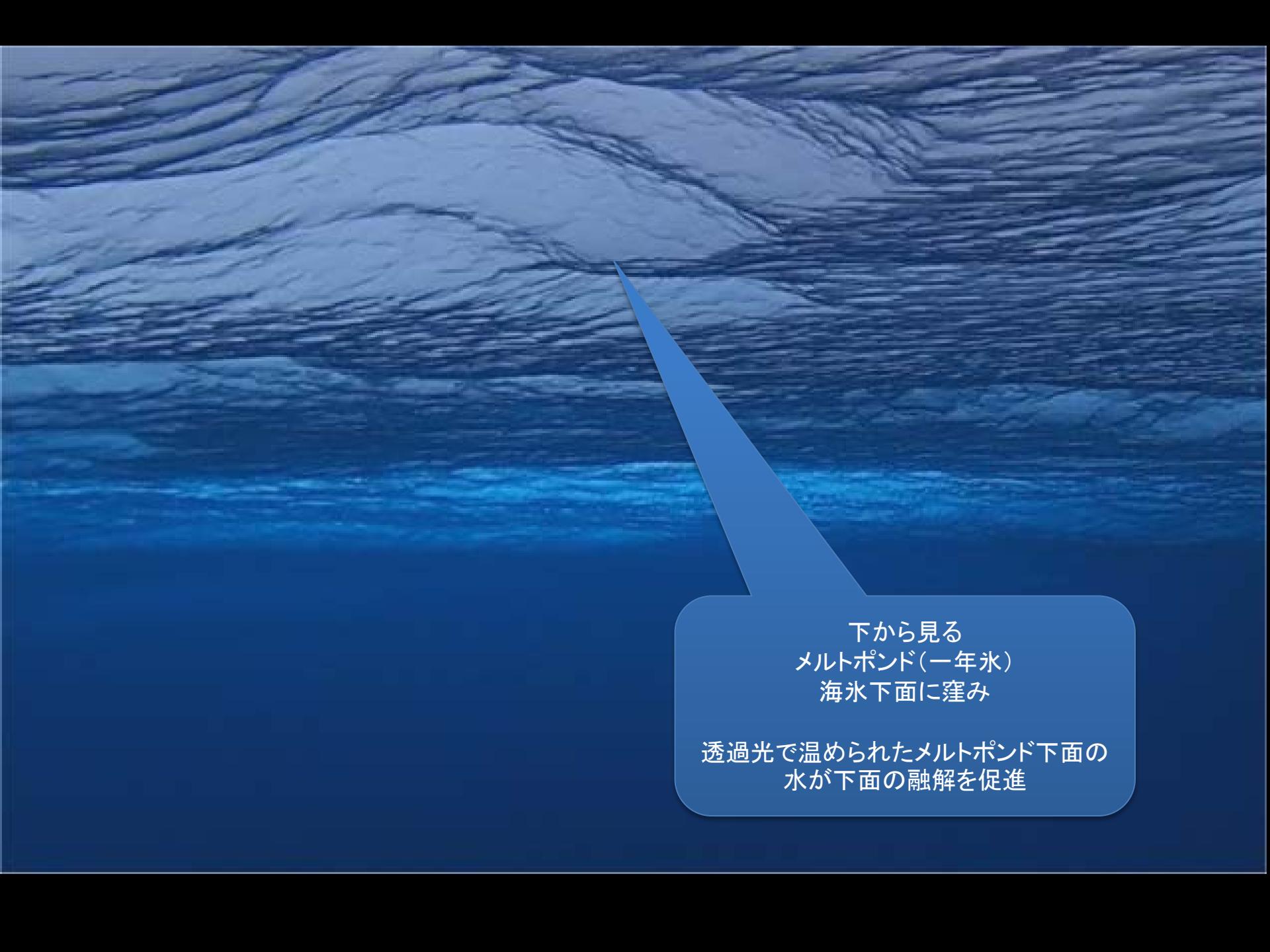








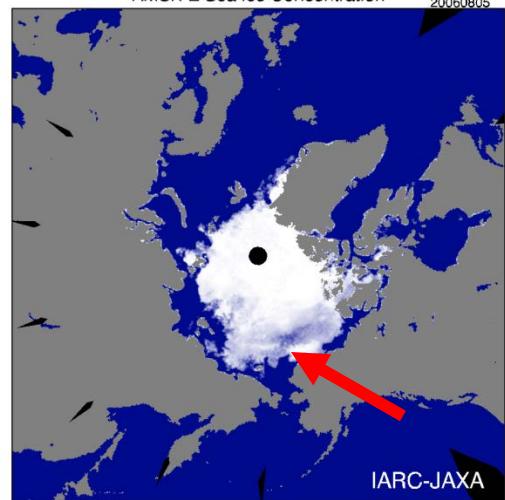


A photograph showing the underside of a multi-layered sea ice pack. The top layer is a thick, white, textured ice sheet. Below it are several thin, dark blue layers of meltponds. A large, semi-transparent blue callout box is positioned in the lower right corner, pointing towards the meltpond layers.

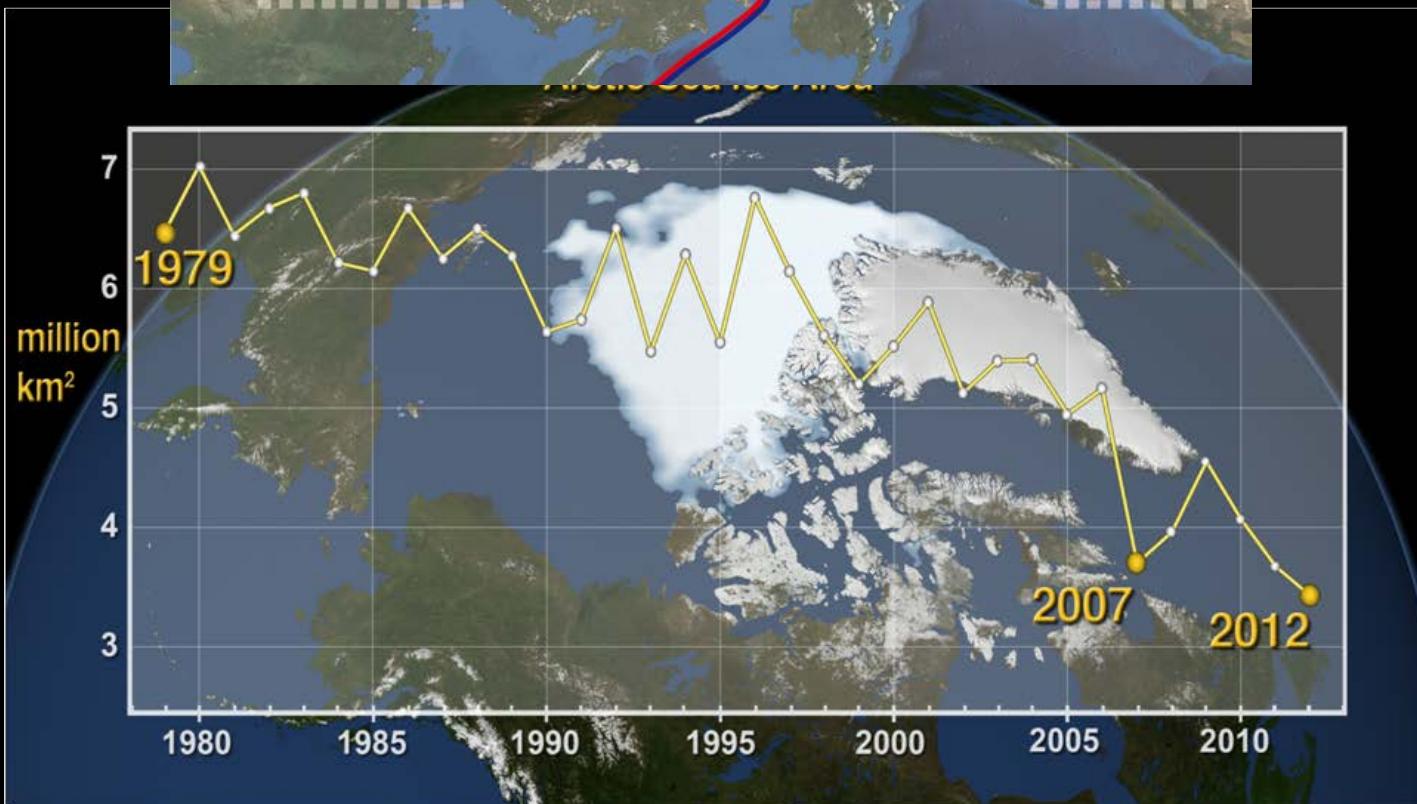
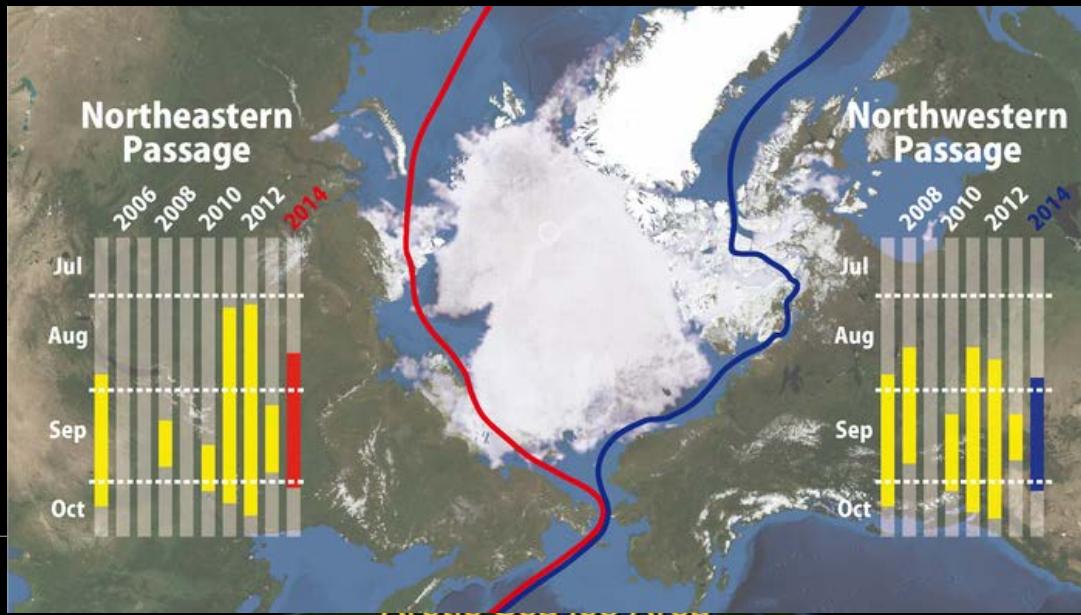
下から見る
メルトポンド(一年氷)
海氷下面に壅み

透過光で温められたメルトポンド下面の
水が下面の融解を促進

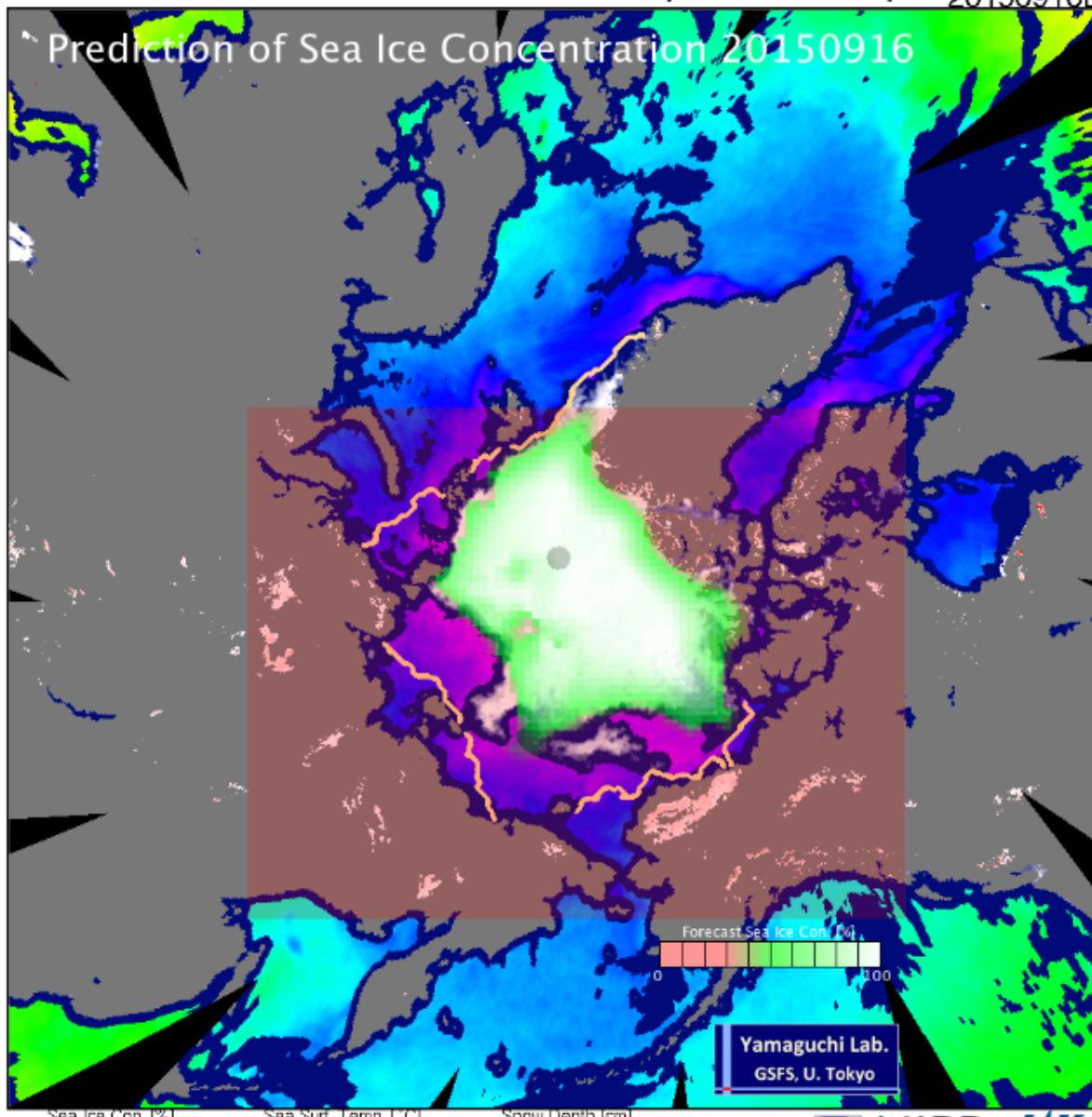
Ice Band

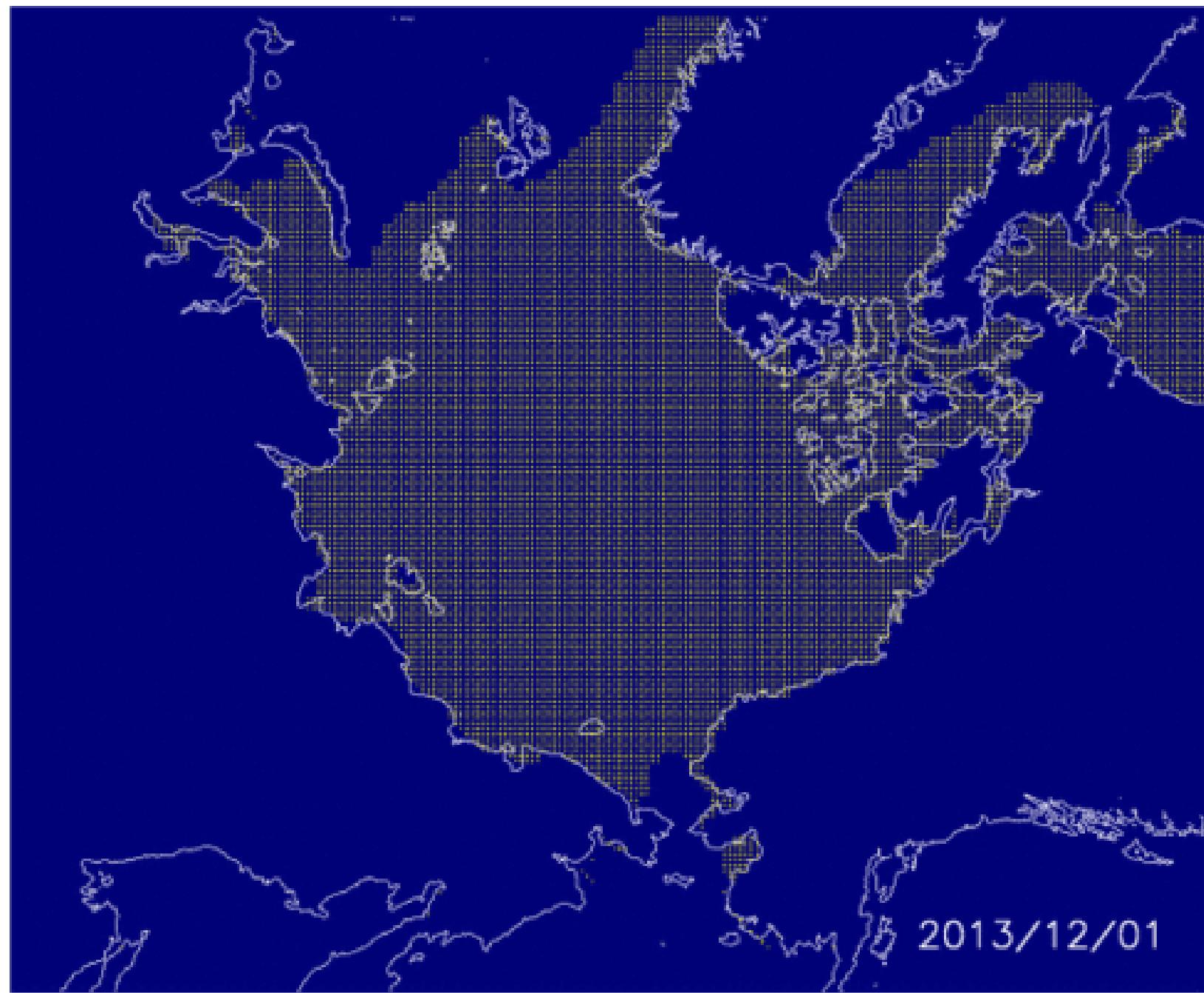


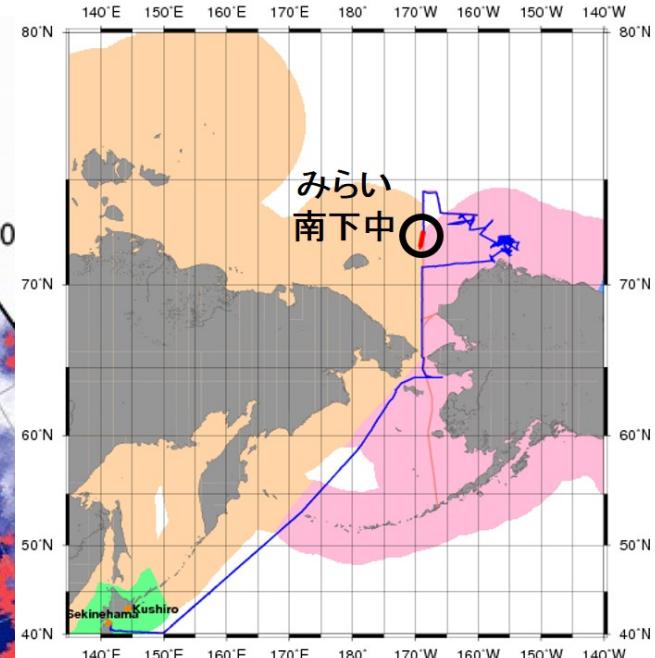
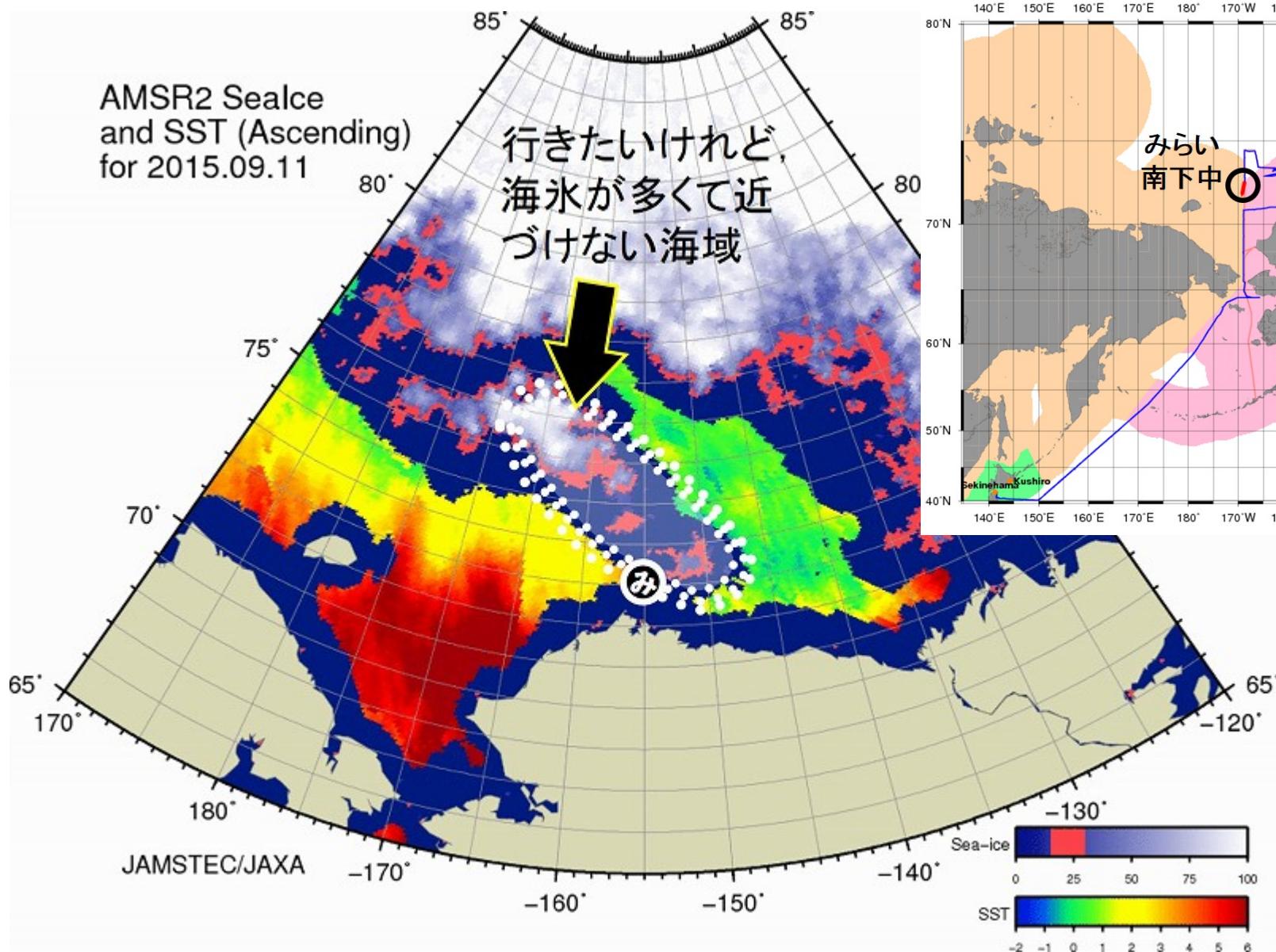
2006 northern Chukchi Sea



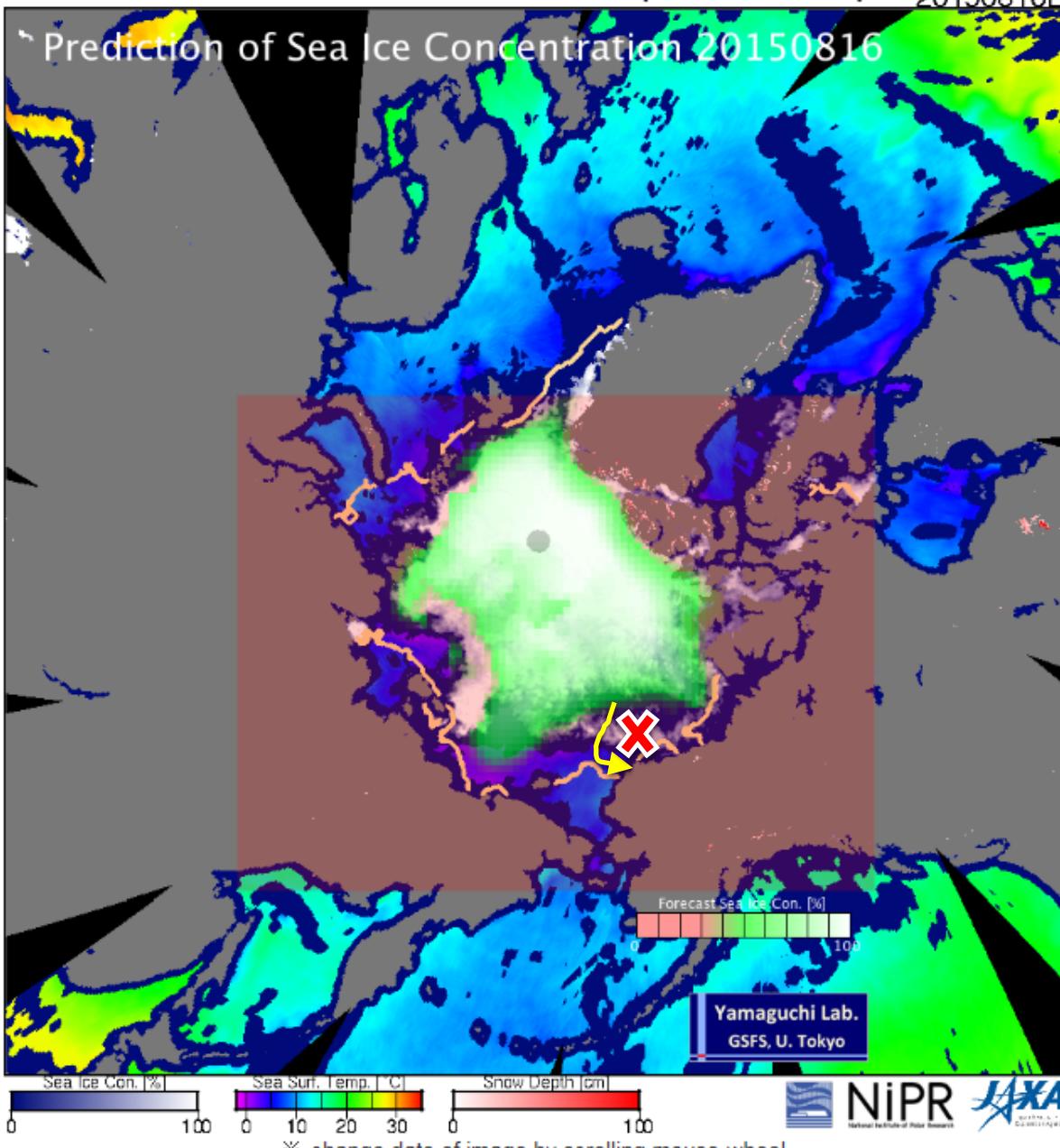
AMSR2 Sea Ice con.+Sea Surf. Temp.+Snow Depth 20150916D







AMSR2 Sea Ice con.+Sea Surf. Temp.+Snow Depth 20150816D

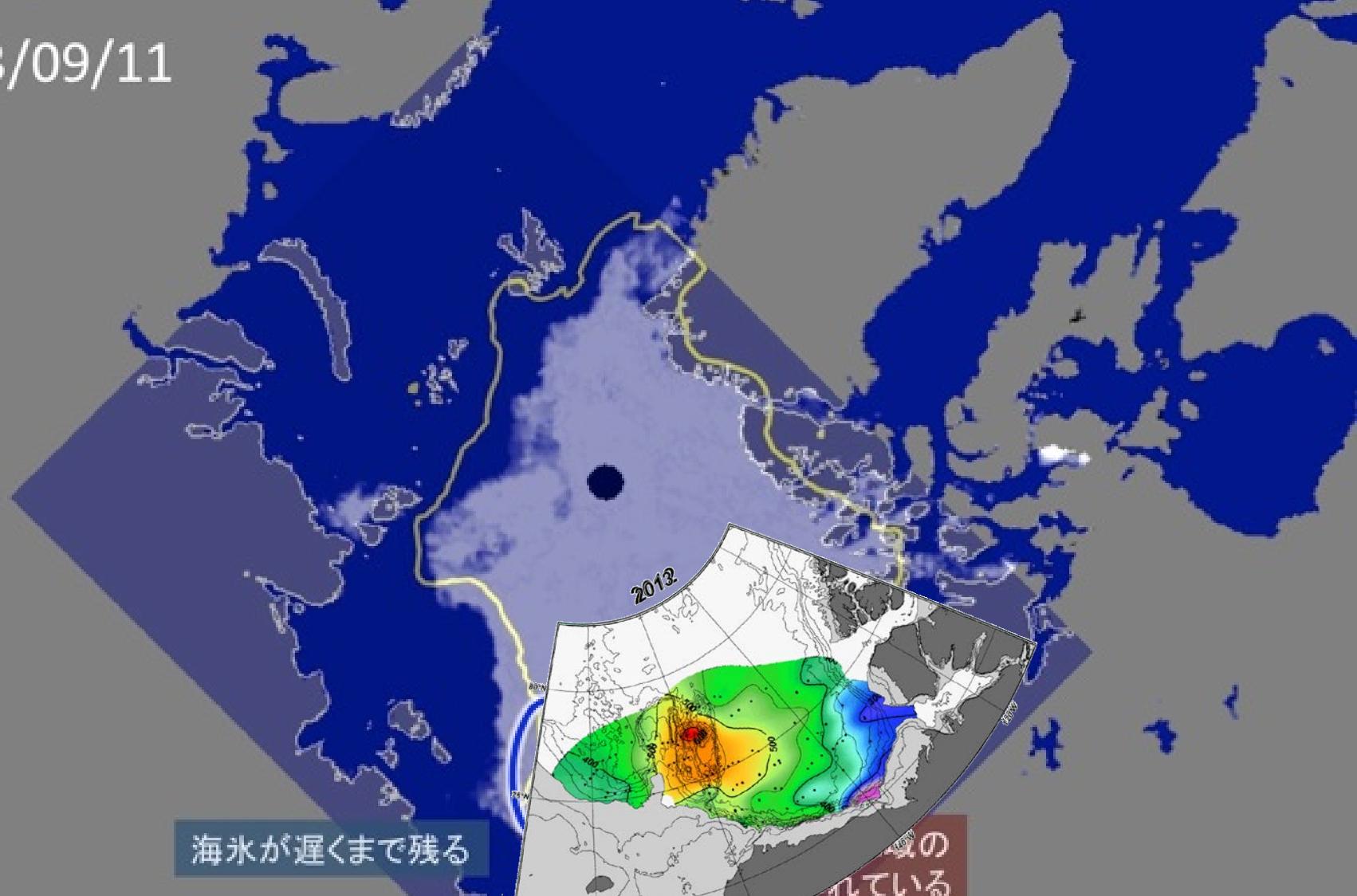


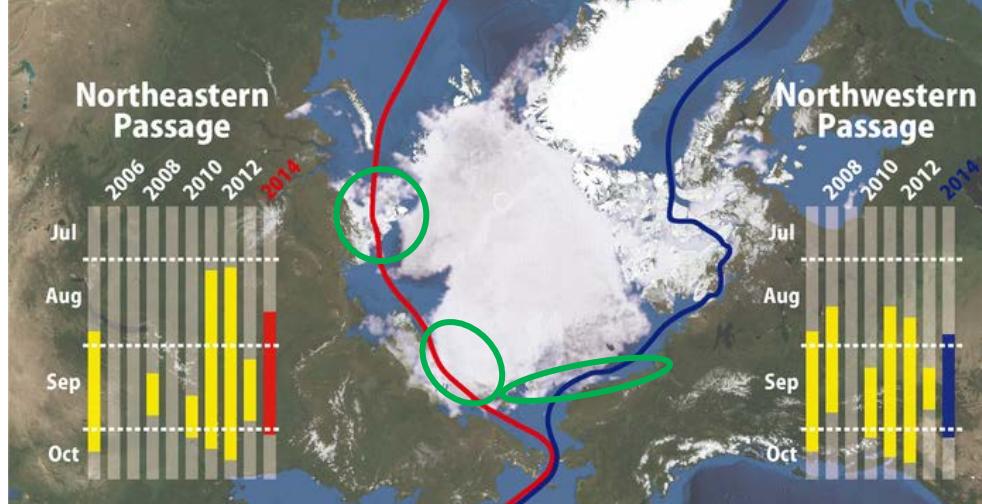
* change date of image by scrolling mouse-wheel.

* zoom/move image area by mouse click (left-drag:zoom, right-drag:move, double-click:reset)



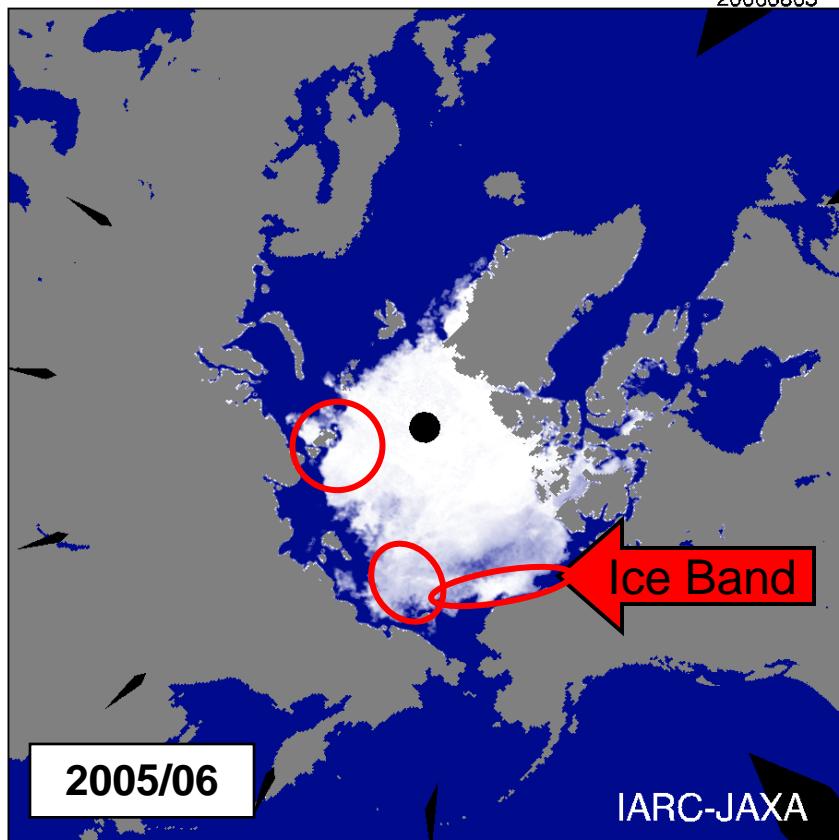
2013/09/11





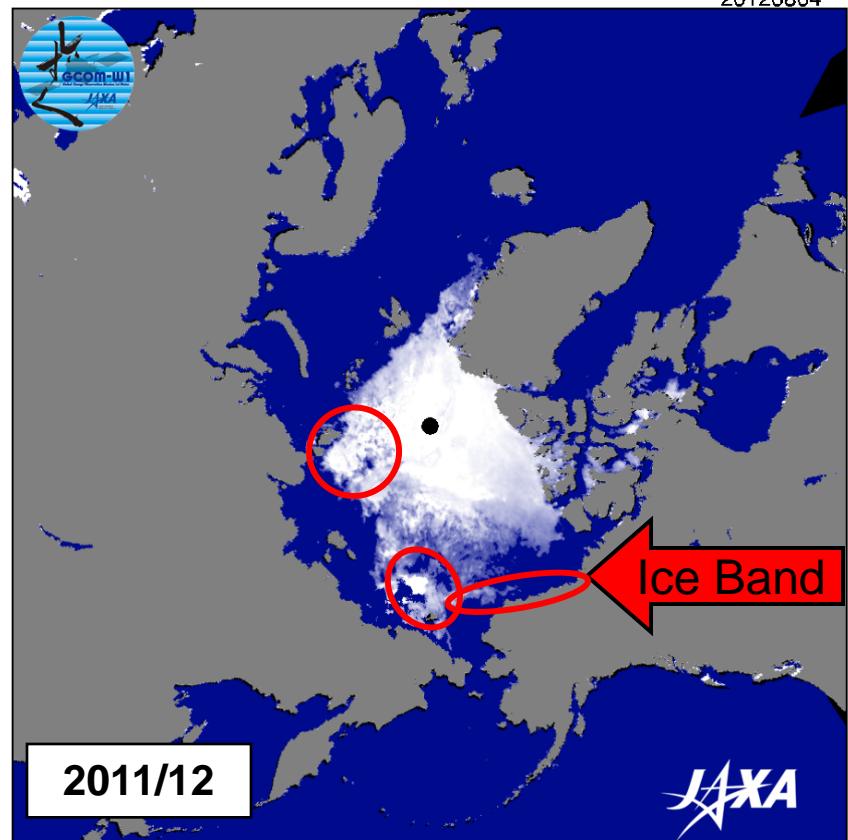
AMSR-E Sea Ice Concentration

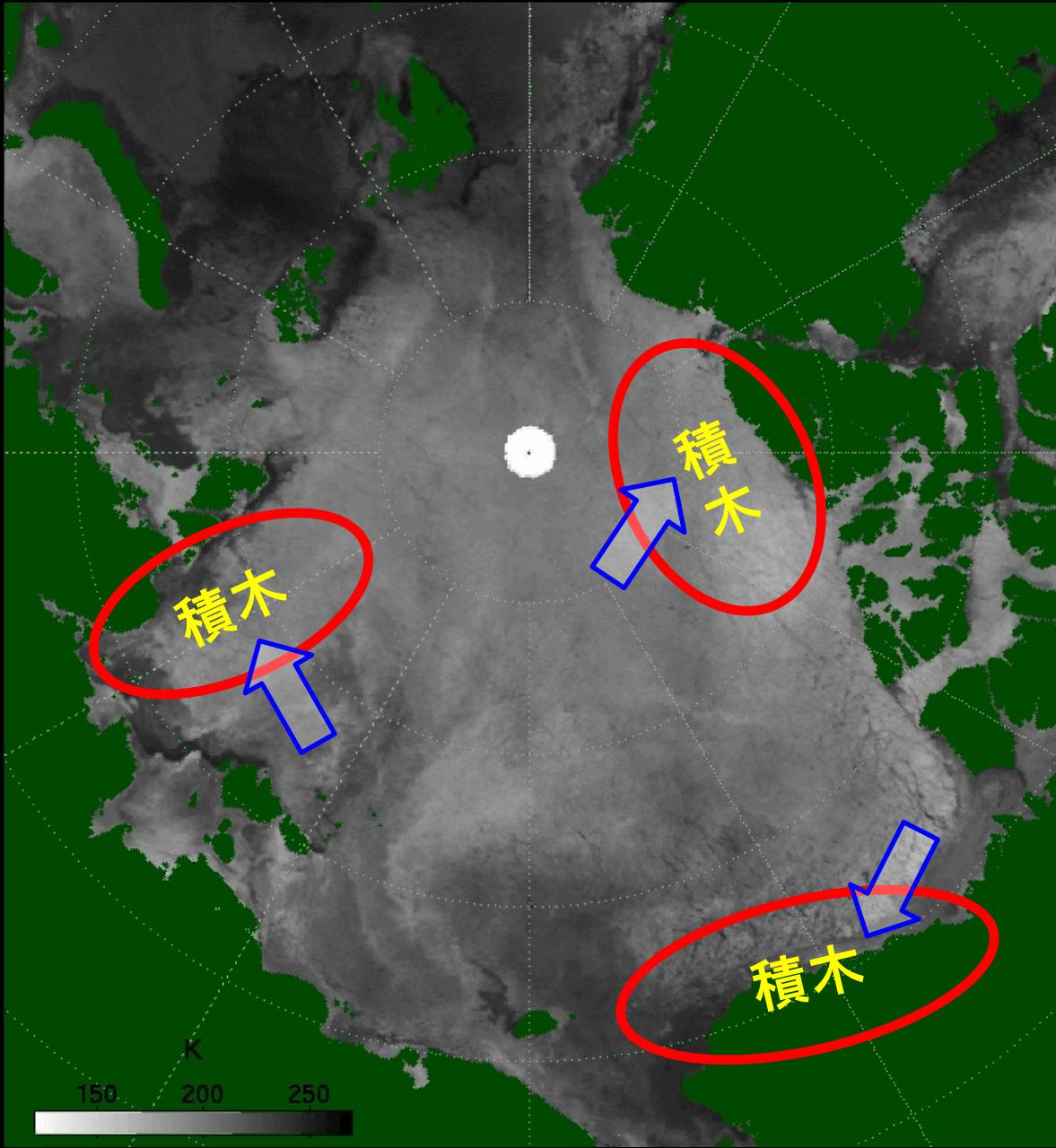
20060805



AMSR2 Sea Ice Concentration

20120804

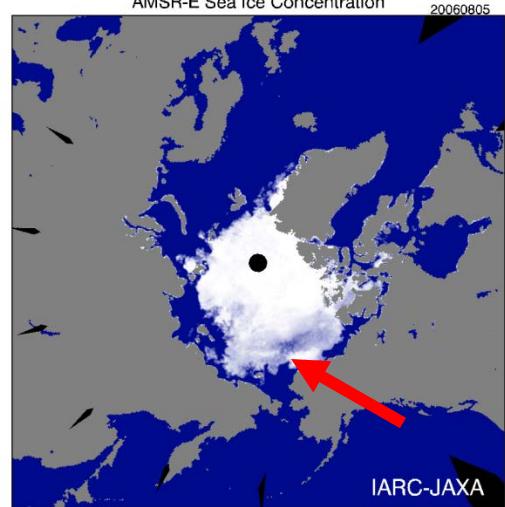




高精度かつ高空間分解能の
信頼できる海氷速度データの構築が必要
⇒島田のポスター



Application for availability of Arctic Sea routes



2006 northern Chukchi Sea

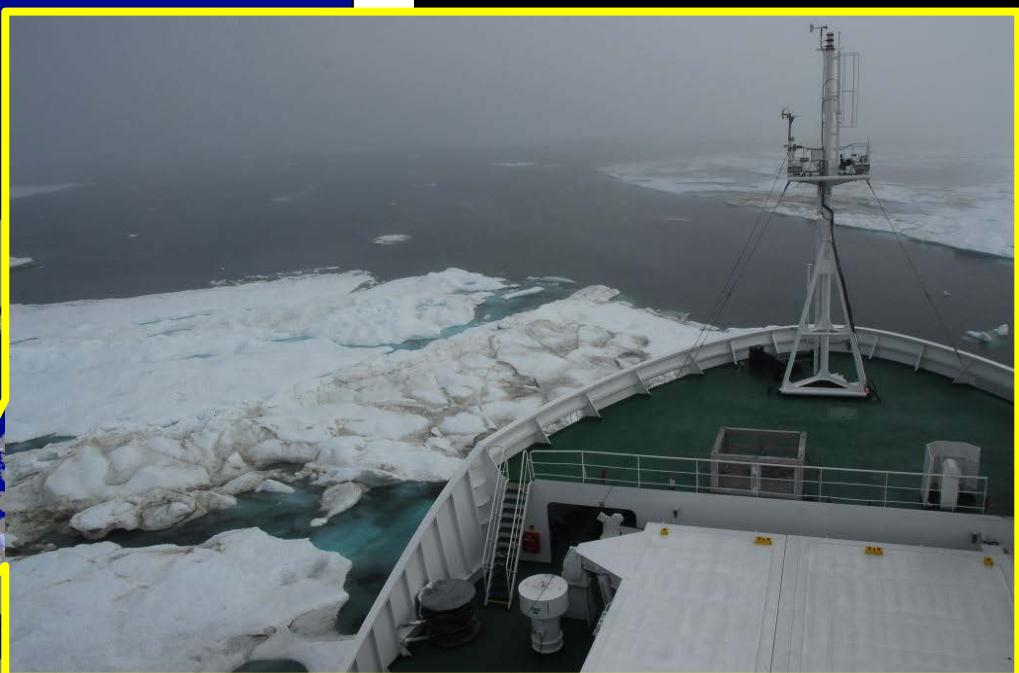
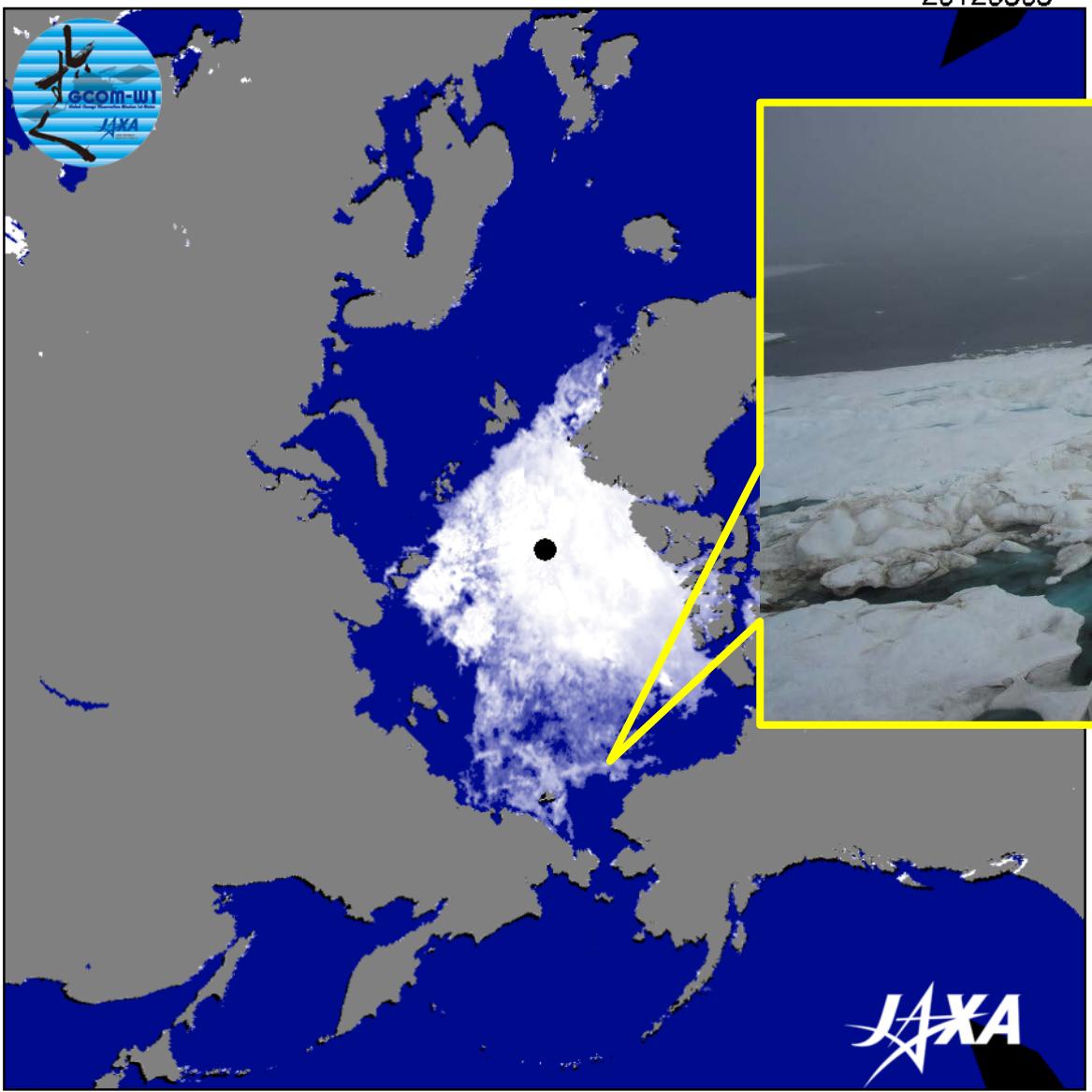


2006 northern Chukchi Sea

Photo by Koji Shimada

AMSR2 Sea Ice Concentration

20120803



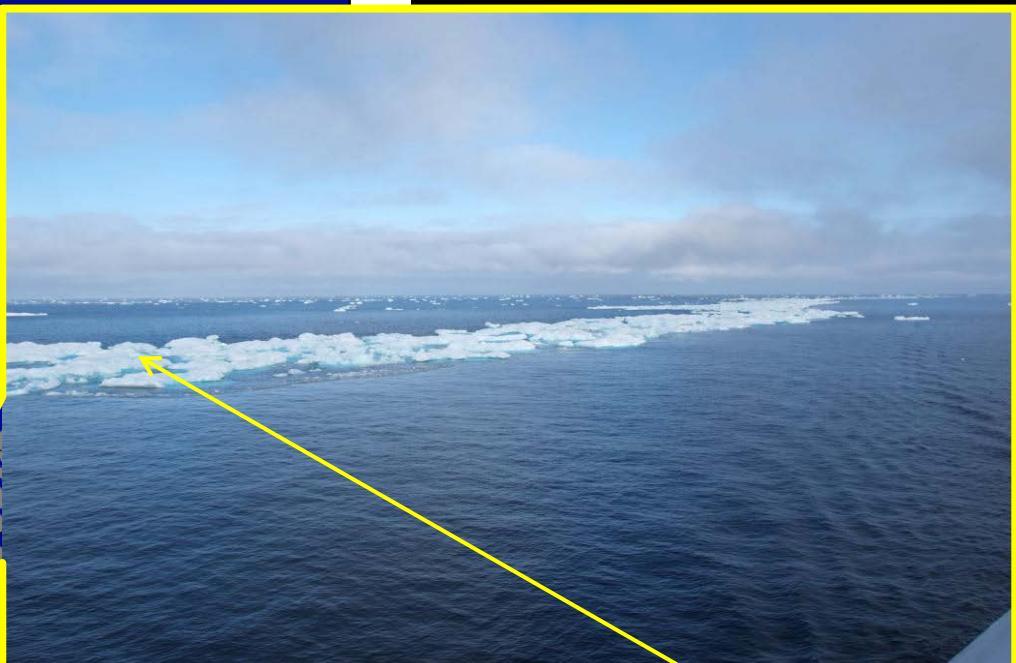
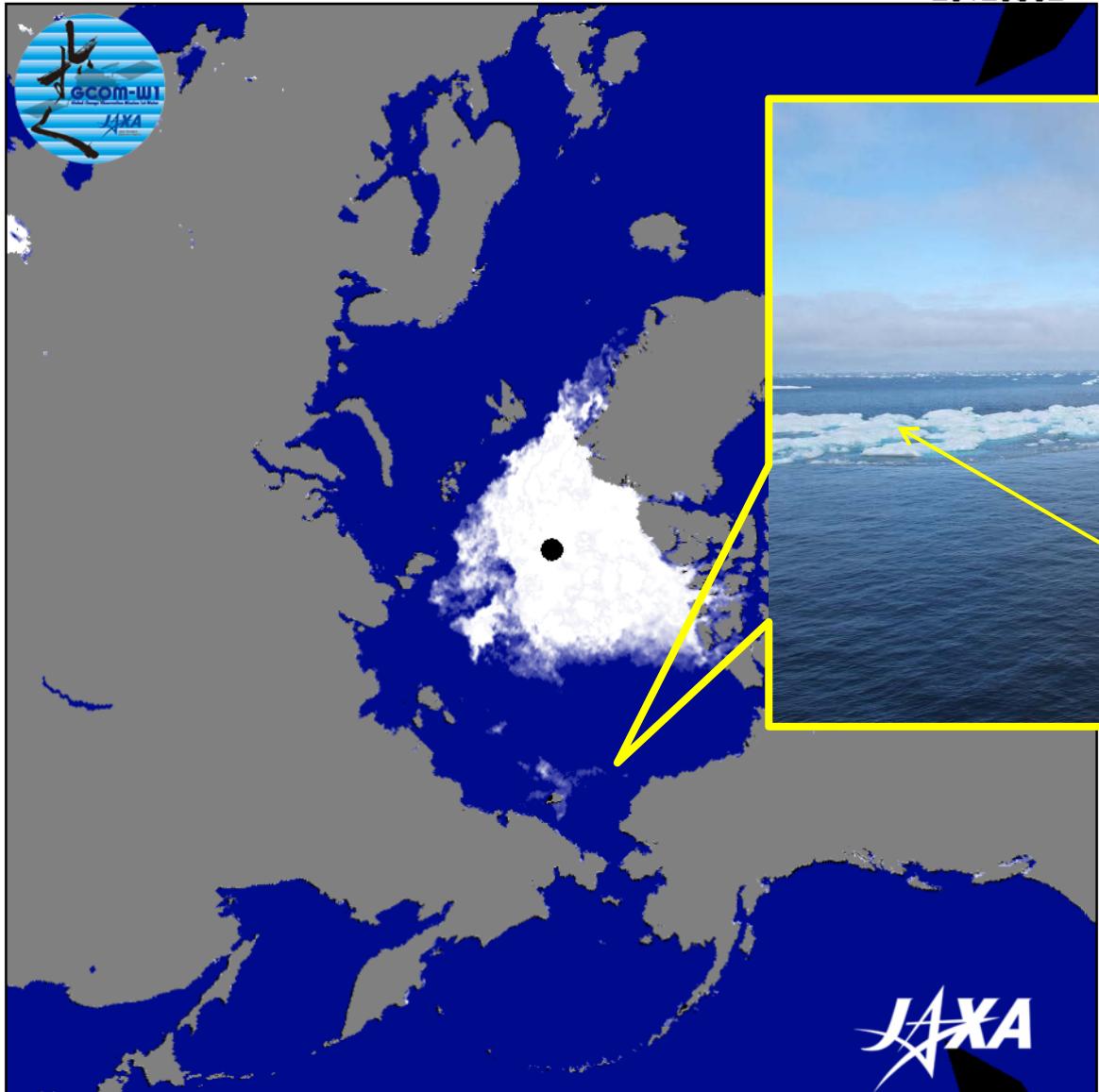
Sea ice data validation is in progress.

The value of sea ice concentration may change after the validation process in future.

JAXA

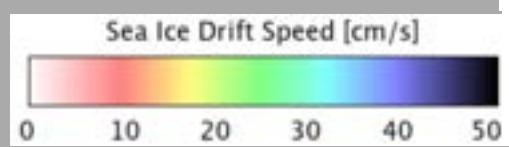
AMSR2 Sea Ice Concentration

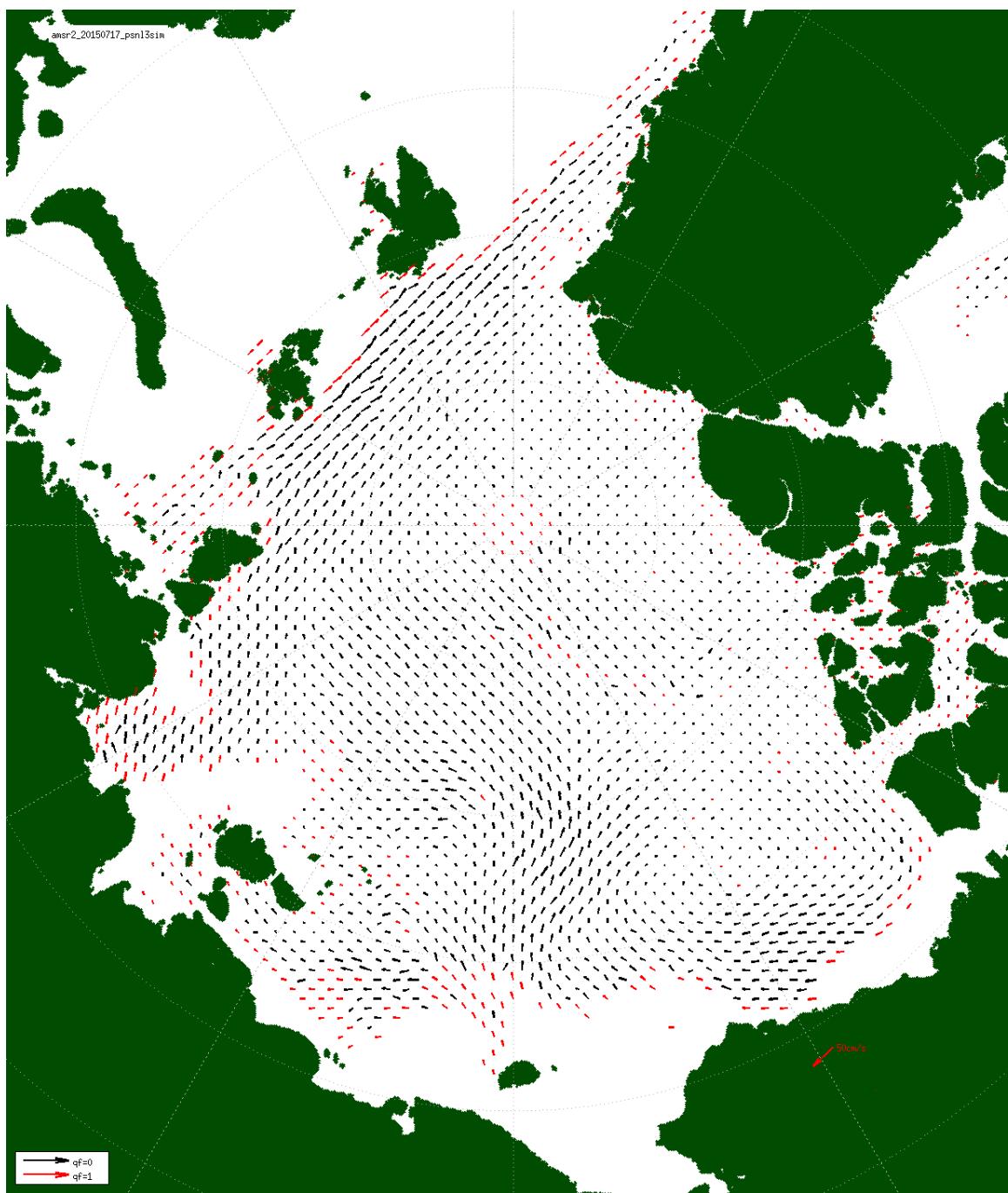
20120902



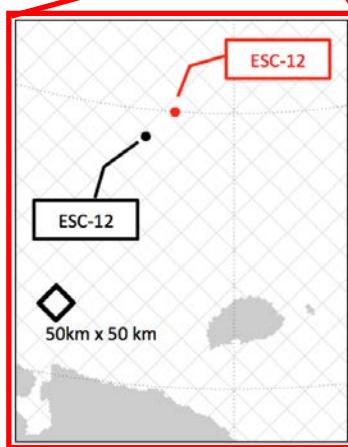
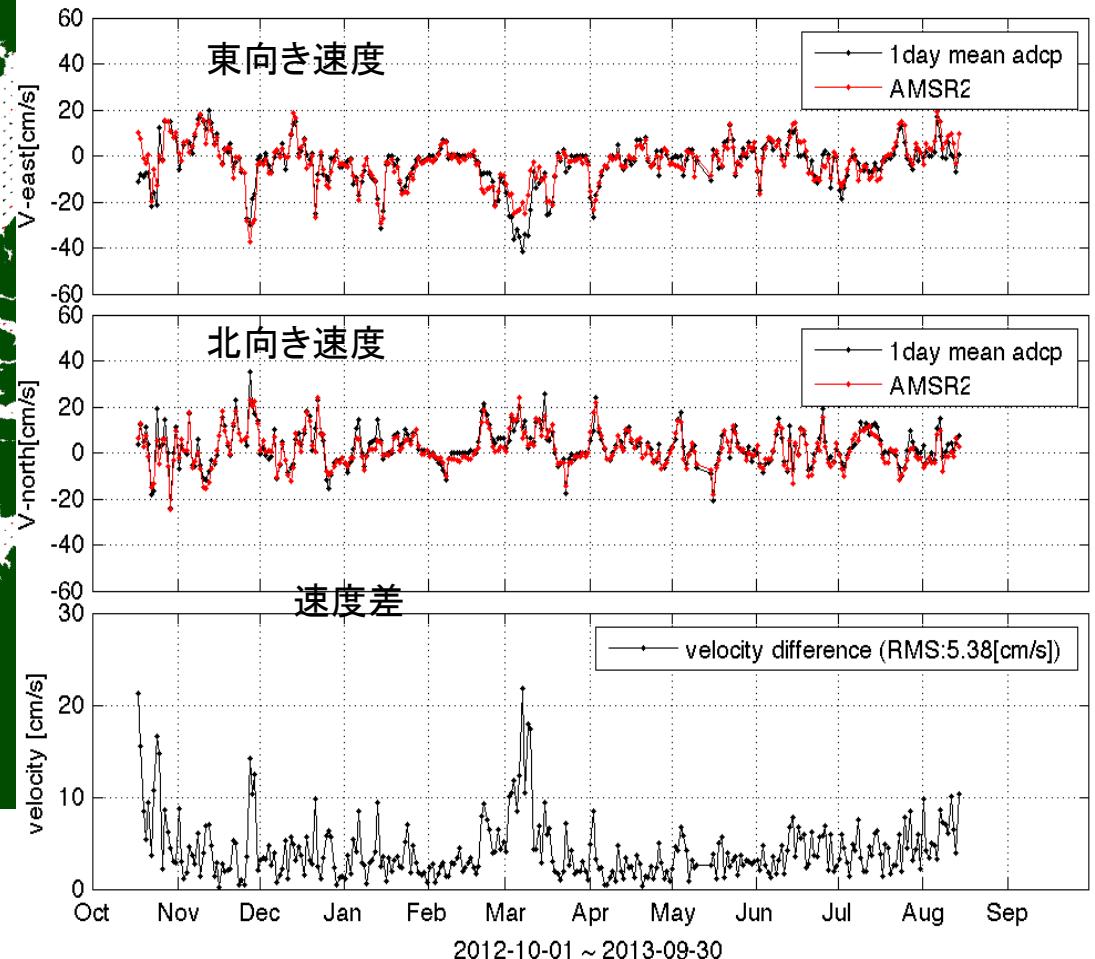
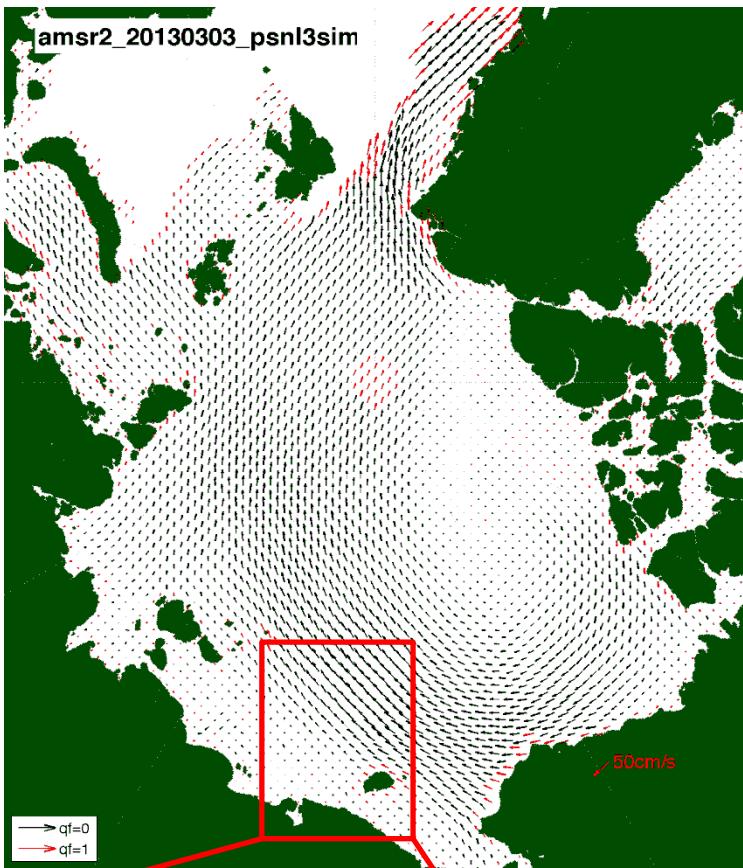
Sea ice data validation is in progress.

The value of sea ice concentration may change after the validation process in future.





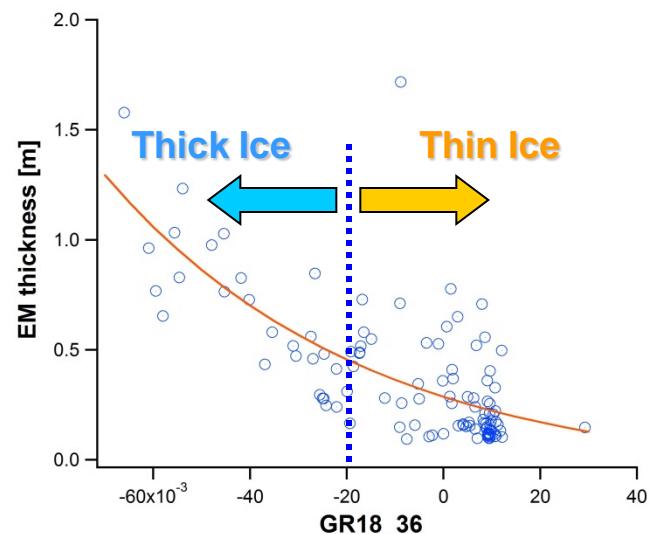
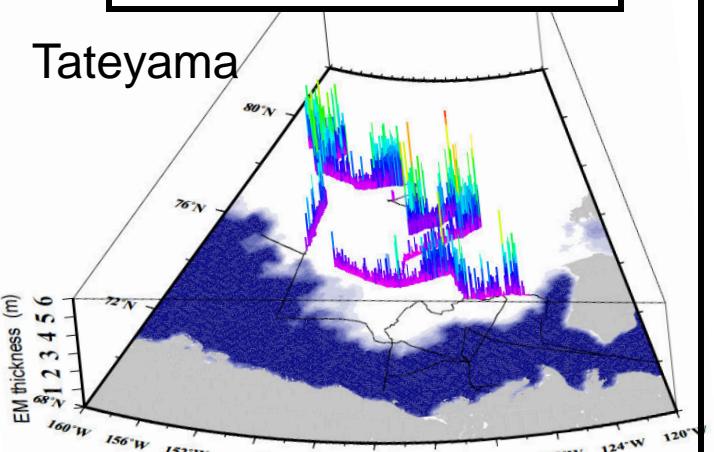
係留系観測(2012-2013) ⇒ AMSR-2ベースの海氷速度ベクトルアルゴリズム



- AMSR-2ベースの海氷速度算出アルゴリズムに利用
- 高精度リアルタイム海氷速度データ提供計画(from JAXA)

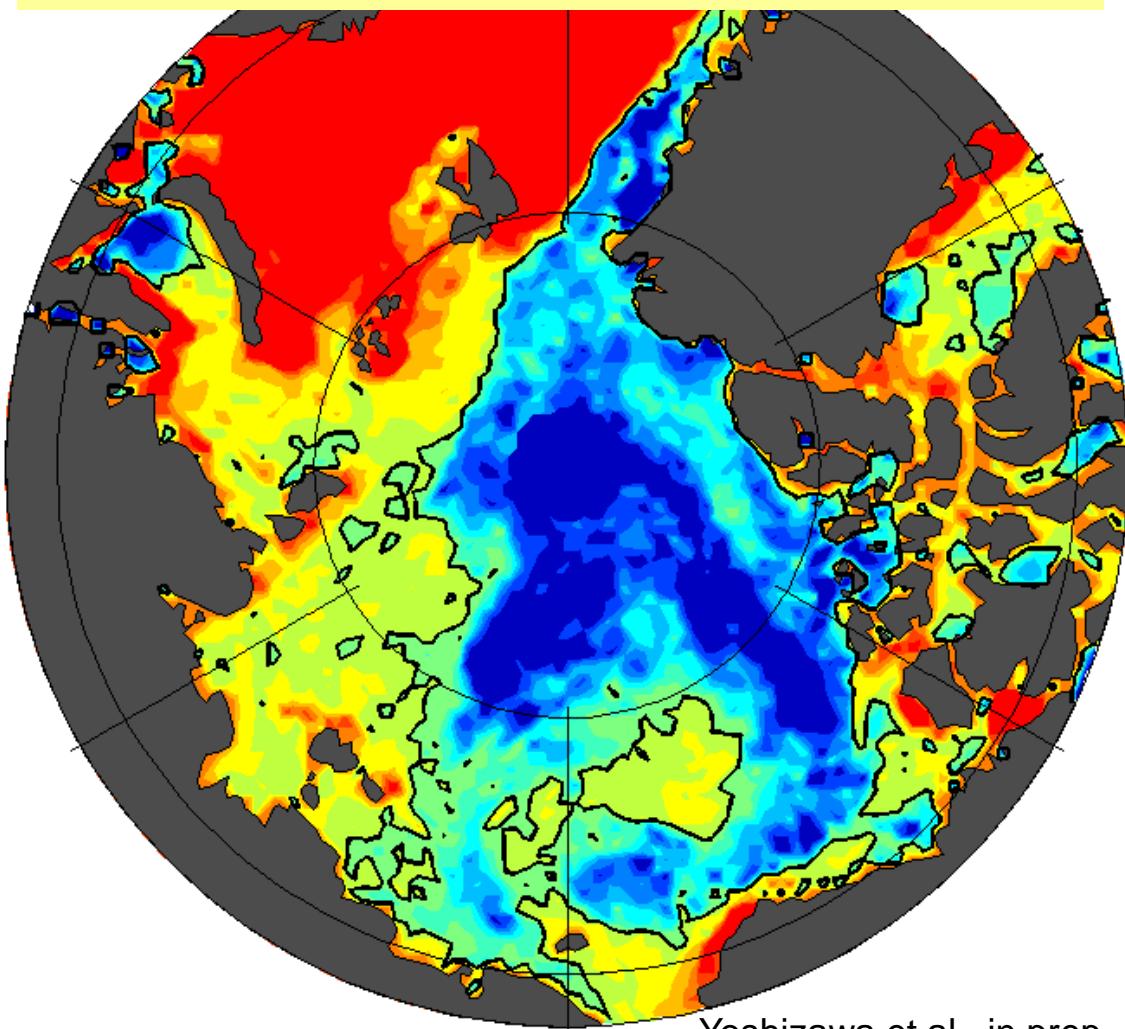
Ice Thickness Group

Tateyama



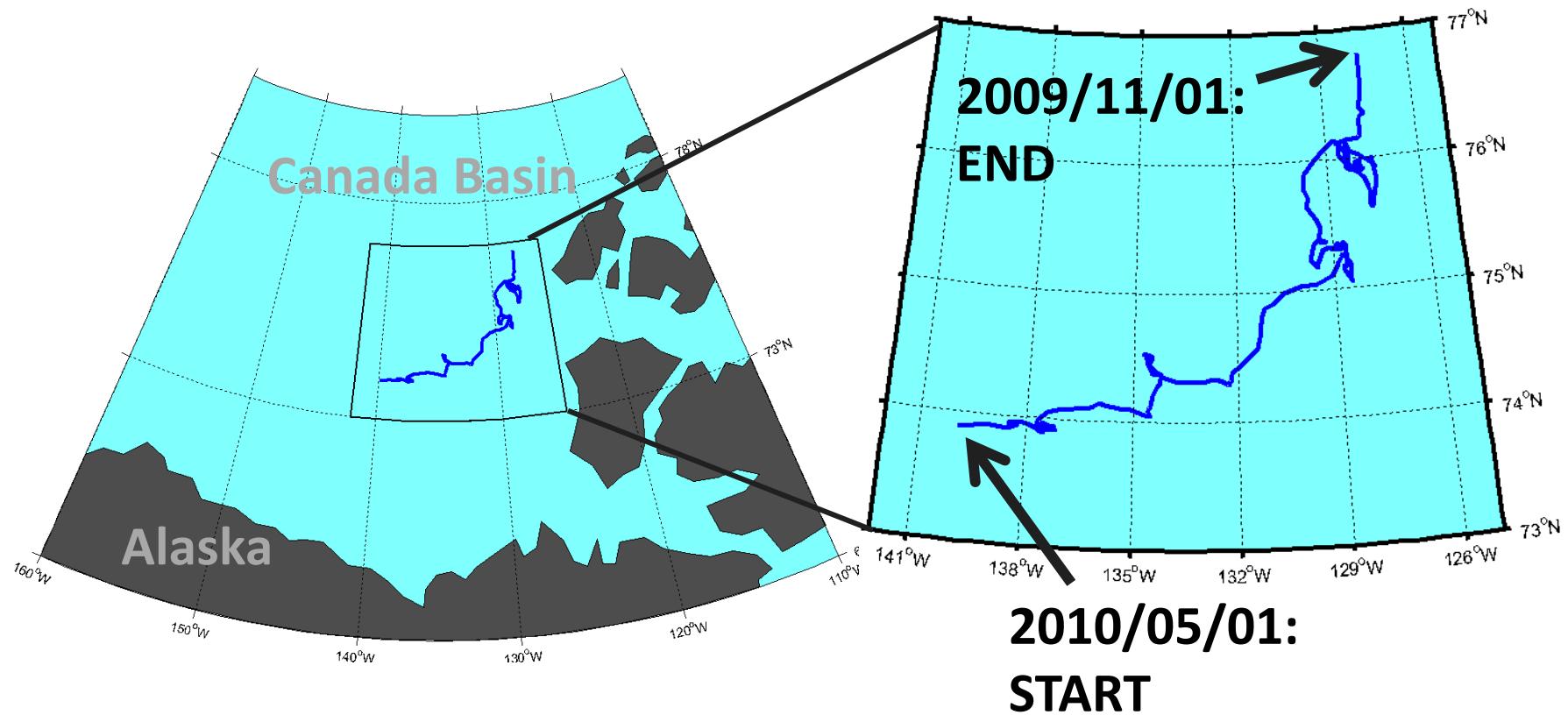
$$GR = \frac{[36V] - [18V]}{[36V] + [18V]}$$

Convergence of thick sea ice motion is important

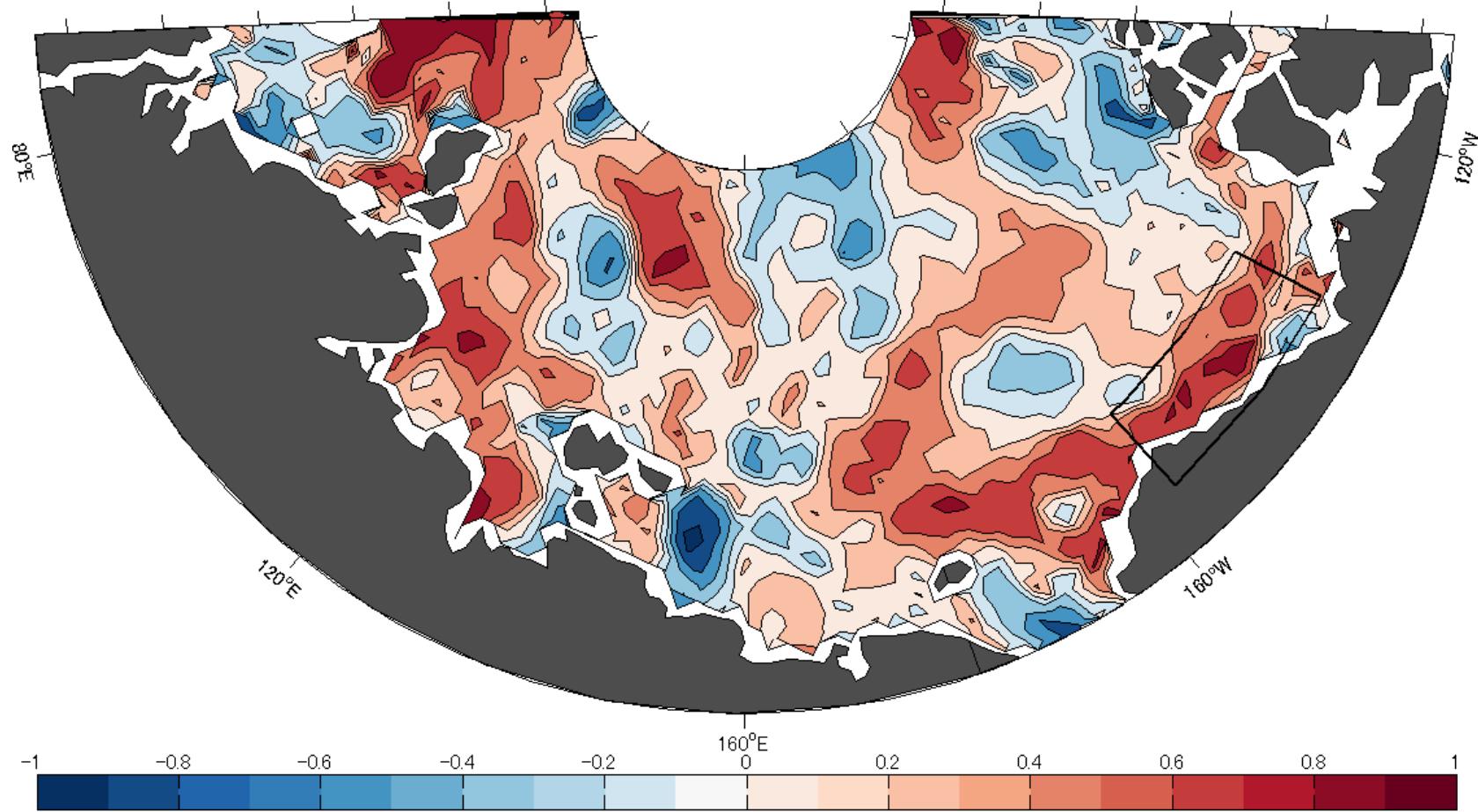


Yoshizawa et al., in prep.

Thick ice Thin ice



- Calculate convergence if GR is less than critical value (-0.02) and sea ice concentration is greater than 98%.
 ⇒ effective convergence for rafting: ECR
- Integrate ECR along drift track of sea ice from November to April.



Correlation between “integrated effective convergence of sea ice along drift track (Nov. ~ Apr.)” and “sea ice concentration in the following summer (Jun. ~ Sep.)”. Box shows a key area of the Northwest passage area (70 - 74°N, 135 - 157°W).

regression

Spring GR only (using just thickness in spring, without sea ice motion)

$$SIC_{GR} = 4.3542 \times \text{GR} + 0.2556.$$

$$(r = 0.2717)$$

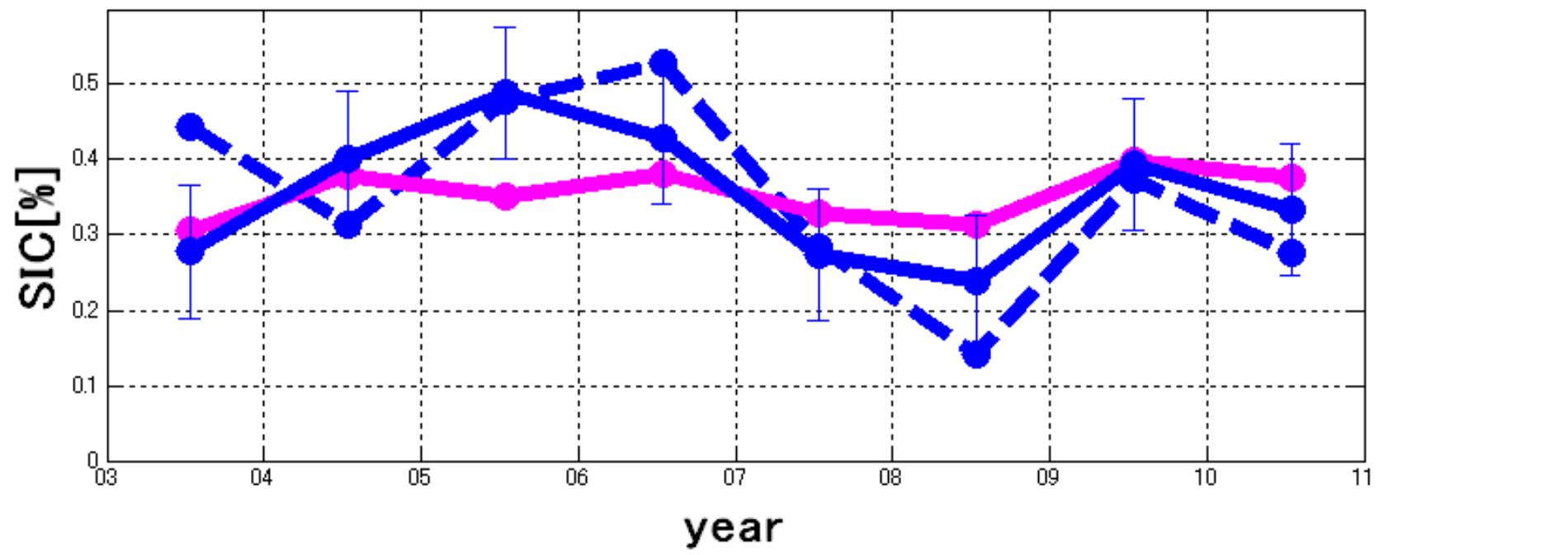
GR and integration of effective convergence for rafting (iECR)

$$SIC_{Yoshi} = 0.6924 \times \text{iECR} + 0.1547$$

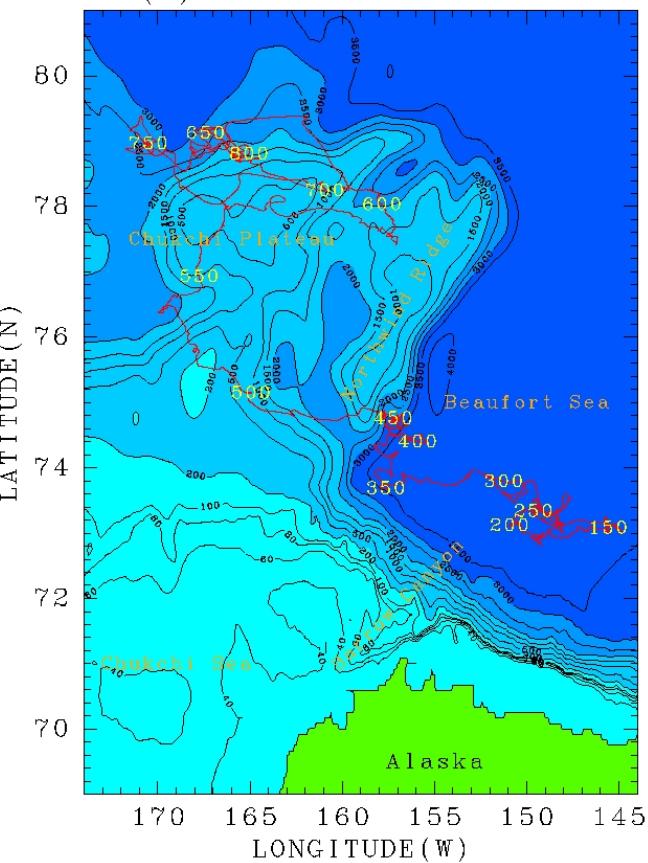
$$(r = 0.6924)$$

forecast

— SIC_{Yoshi} — SIC_{GR} - - - observed

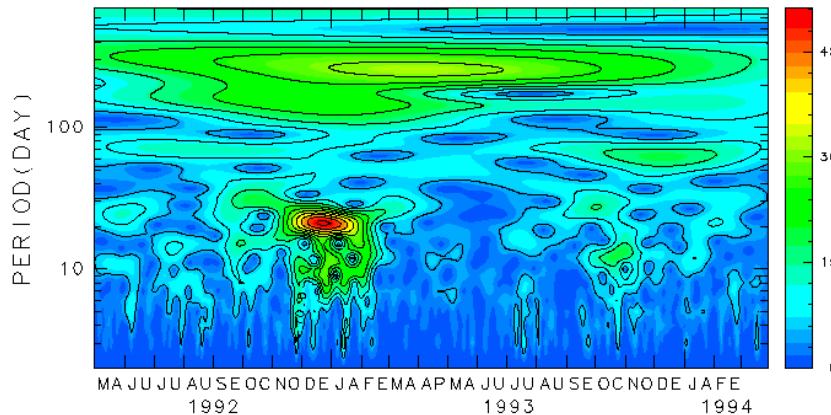


(a) DRIFT TRACK OF IOEB

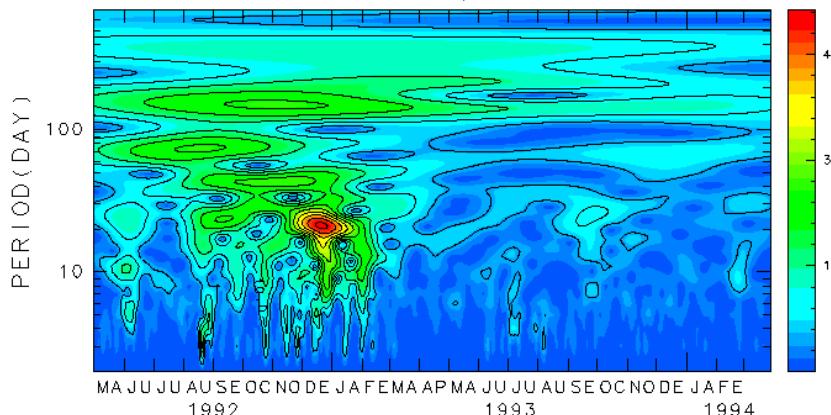


$T=5.4d$: $R=15\text{km}$, $U=20\text{cm/s}$
 $T=22d$: $R=30\text{km}$, $U=10\text{cm/s}$

ioeb1acp1ue



ioeb1acp3ue



ioeb1acp5ue

